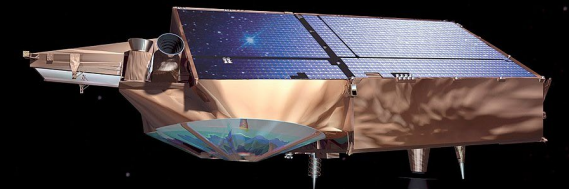


# The 4DGreenland project

Quantifying Greenland ice sheet hydrology components from remote sensing

*Louise Sandberg Sørensen & the 4DGreenland team*



- Funded by ESA through the POLAR+ programme, Theme 2.
- The project runs Sept 2020- Sept 2022. Nine partners.

## Objectives:

- Generate novel Earth Observation-derived datasets characterizing the different components of the hydrological system.
- Perform thorough validation of all derived products
- To perform an integrated pan-Greenland scientific analysis and study of the hydrological system.
- Advance our understanding of the Greenland hydrology and its impacts on the Greenland and Arctic environment.



# The 4DGreenland team



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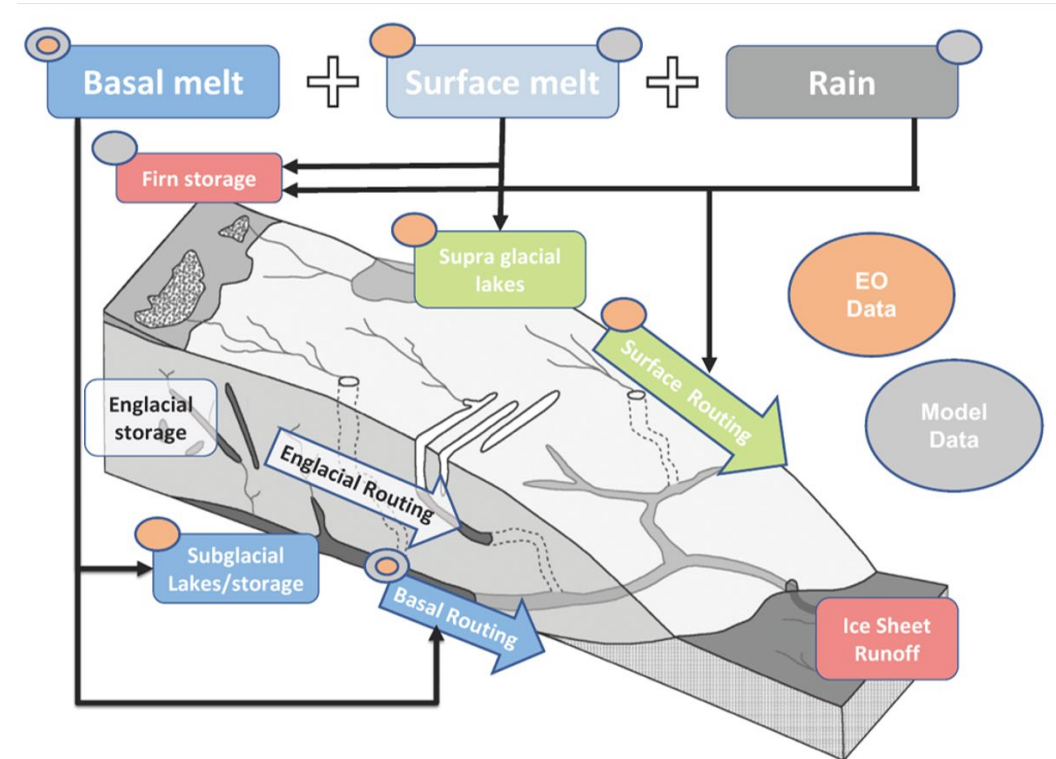
Nanna Karlsson, Anne Solgaard, Sofia Ribeiro

GEUS



Andrew Shepherd

- Activity 1 : **Surface melt processes**
- Activity 2 : **Supraglacial storage and drainage**
- Activity 3 : **Subglacial melt, drainage and lakes**
- Activity 4 : **Integrated Greenland hydrology assessment**



*Background image adapted from Cuffey and Paterson (2010)*

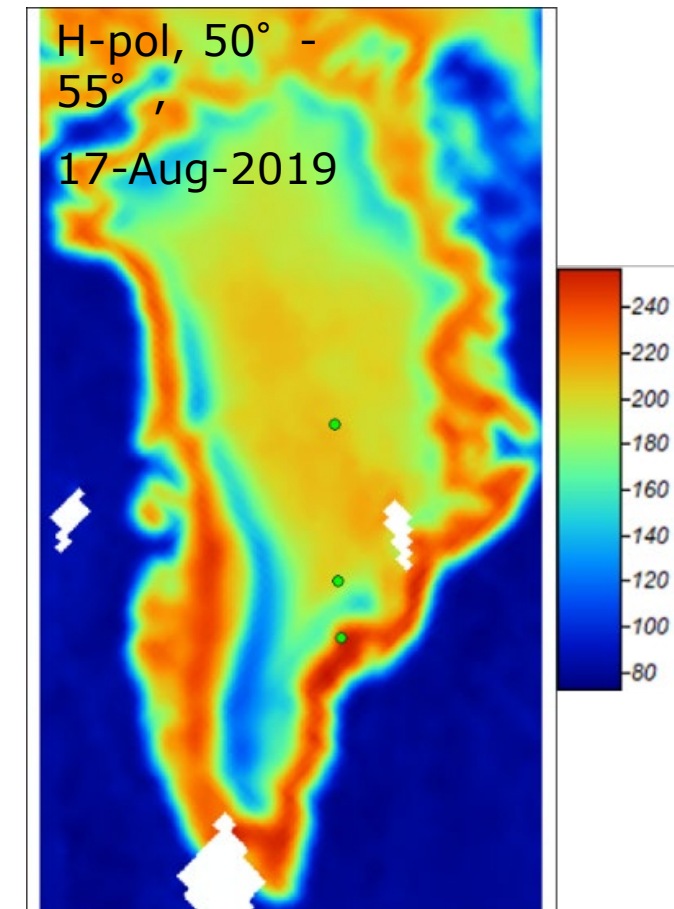
## Primary goal

- Develop, test & implement algorithm for generating maps of surface melt extent and melt phase from:
  - High to medium resolution C-band backscatter measurements (S1, ASCAT)
  - Lower resolution PMW data (SMOS, SMAP, SSMIS, AMSR2)
  - Investigate possibility to use the developed methods in a synergistic way
- Value added products: onset & end date per year, melt duration, melt stage
- Assessment of algorithms with in-situ meteorologic data (AWS) and Regional Climate Models (RCMs)

Sentinel-1 TOPS IW



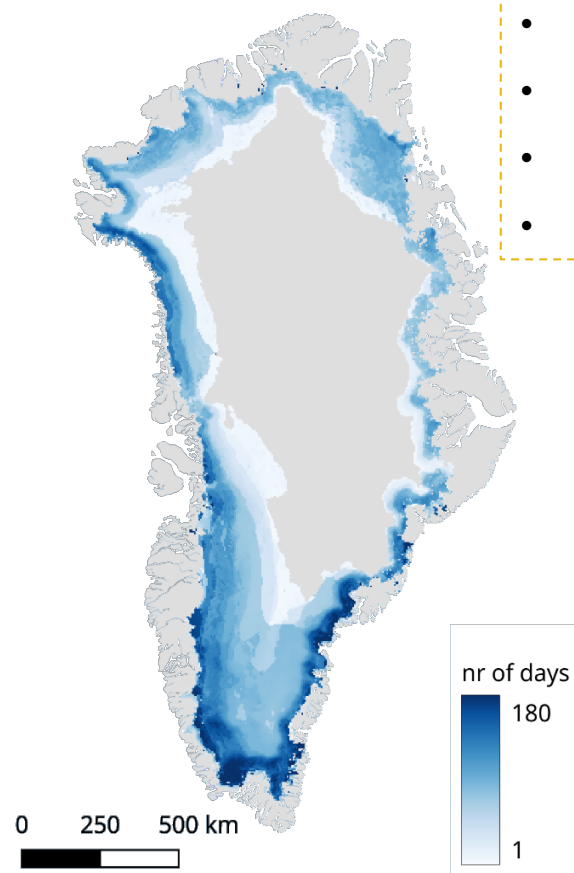
SMOS



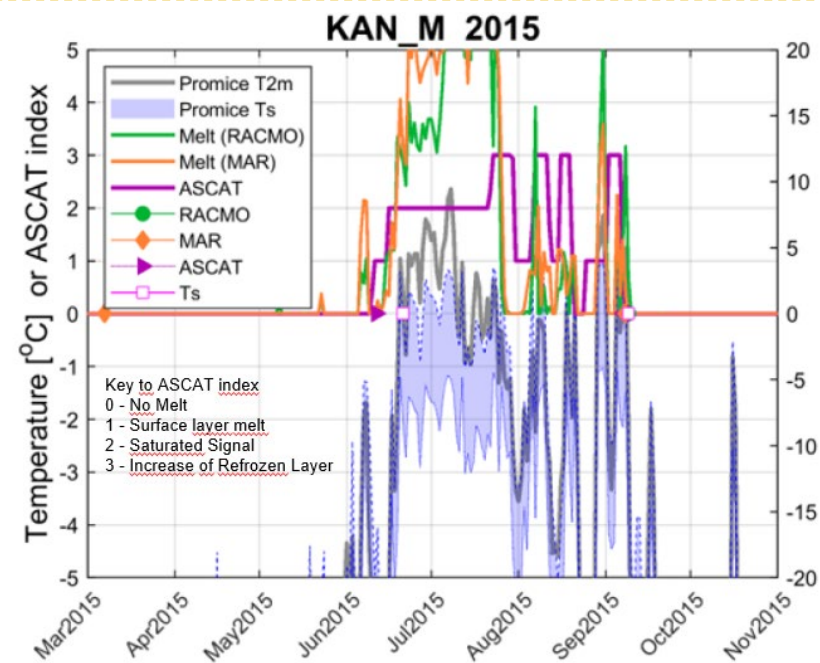
2012-05-01



No melt  
 Increase of wet snow layer  
 Wet snow layer (saturated signal)  
 Increase of refrozen layer above wet snow



- Identification of surface melt and refreezing by ASCAT
- Combined with Sentinel-1 backscatter maps (co&cross pol)
- Method supported by backscatter modelling
- Good agreement with RCMs and AWS data



ASCAT data:  
SIR product BYU - MERS

Daily 5-km ASCAT\* melt maps 2012

Melt Duration 2020

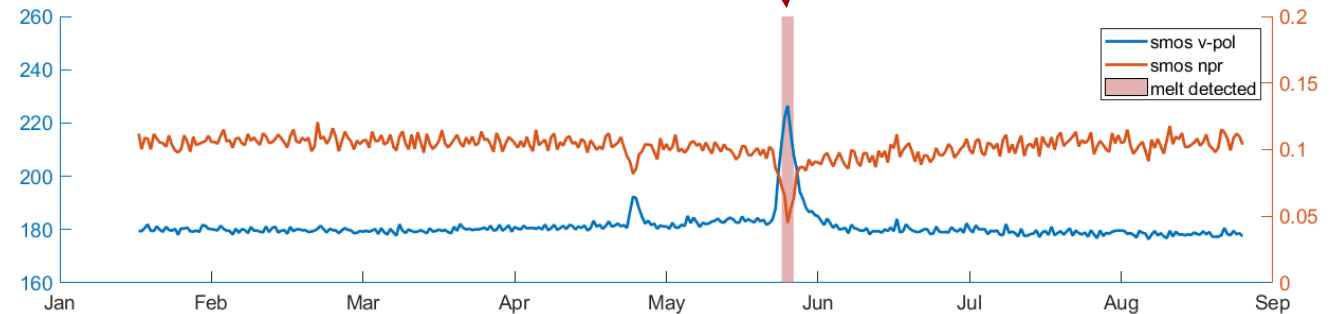
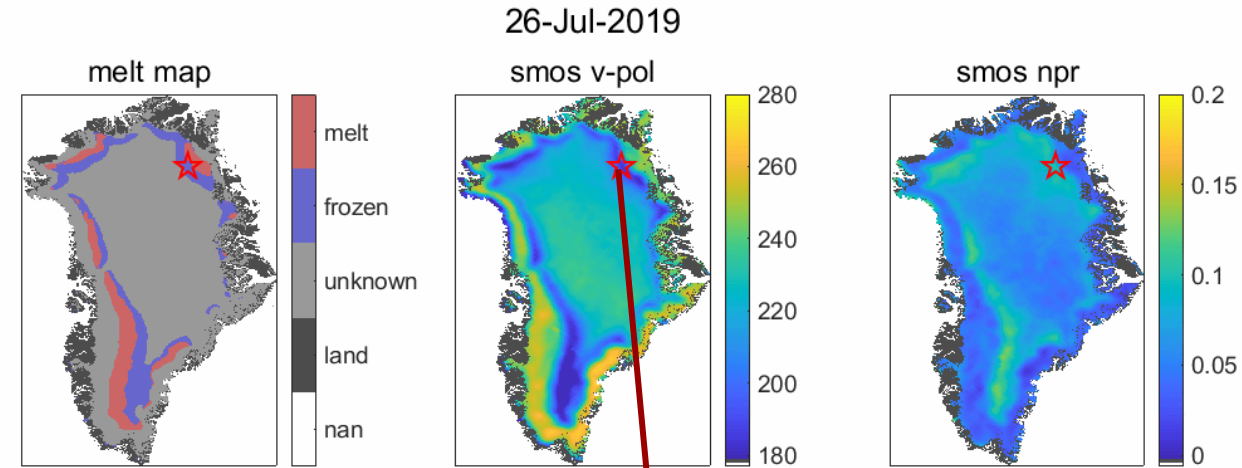
AWS/RCM Intercomparison

## L-band PWM melt detection to support high resolution SAR method

- L-band – different response to the melt event, potential information from deeper layers. Less sensitive to structural changes after re-freeze.
- Brightness temperature and NPR signals for melt detection

$$NPR = \frac{T_{b,V} - T_{b,H}}{T_{b,V} + T_{b,H}}$$

- Method developed originally for SMAP (Mousavi et al., 2021) - sigma above winter mean.

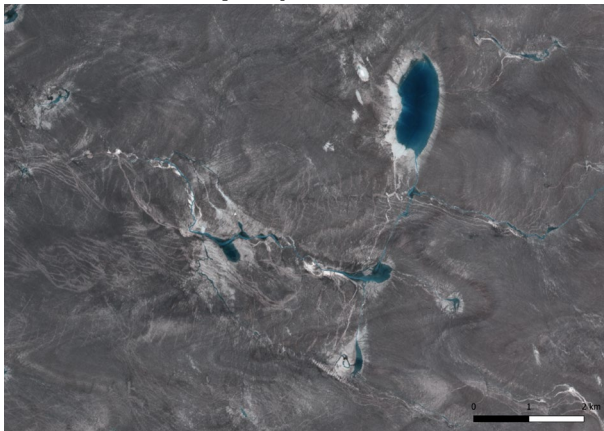


*Melt detection using SMOS Level 3TB, 2019*

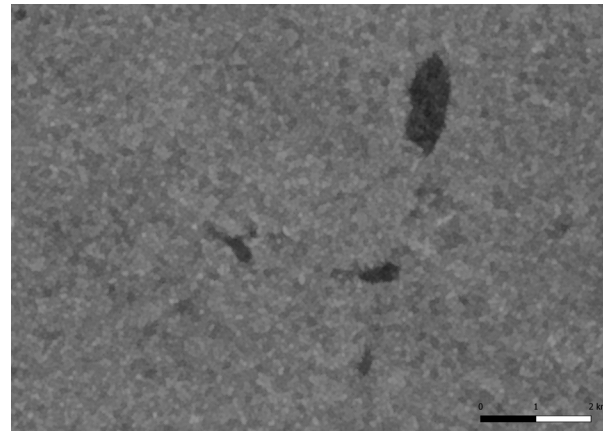
## Proof-of-concept studies within this work package include:

Exploring SAR imagery to map supraglacial lakes using dynamic thresholding and deep learning methods.

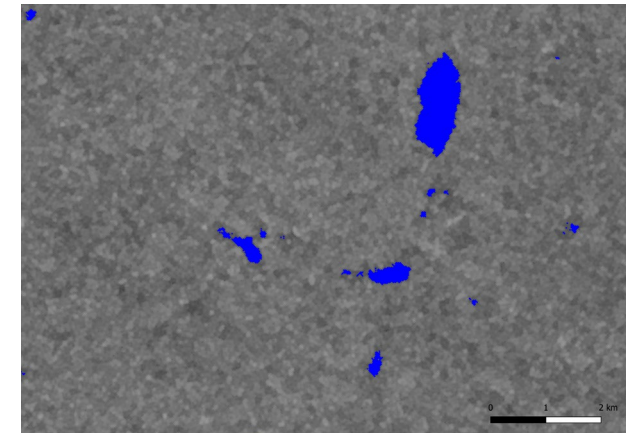
- SAR imagery is advantageous compared to optical datasets for two main reasons: 1) SAR is not limited by cloud cover unlike optical datasets; and 2) it is possible to monitor supraglacial lakes through winter with SAR.
- The proof-of-concept study has shown large lakes are detected and delineated with accuracy from SAR but small channels (<50 m wide) were not identified. There is an opportunity for further method development to ensure methods are mature enough to roll out over large (ice sheet basin) spatial scales.



Sentinel-2 optical image 25<sup>th</sup> July 2019



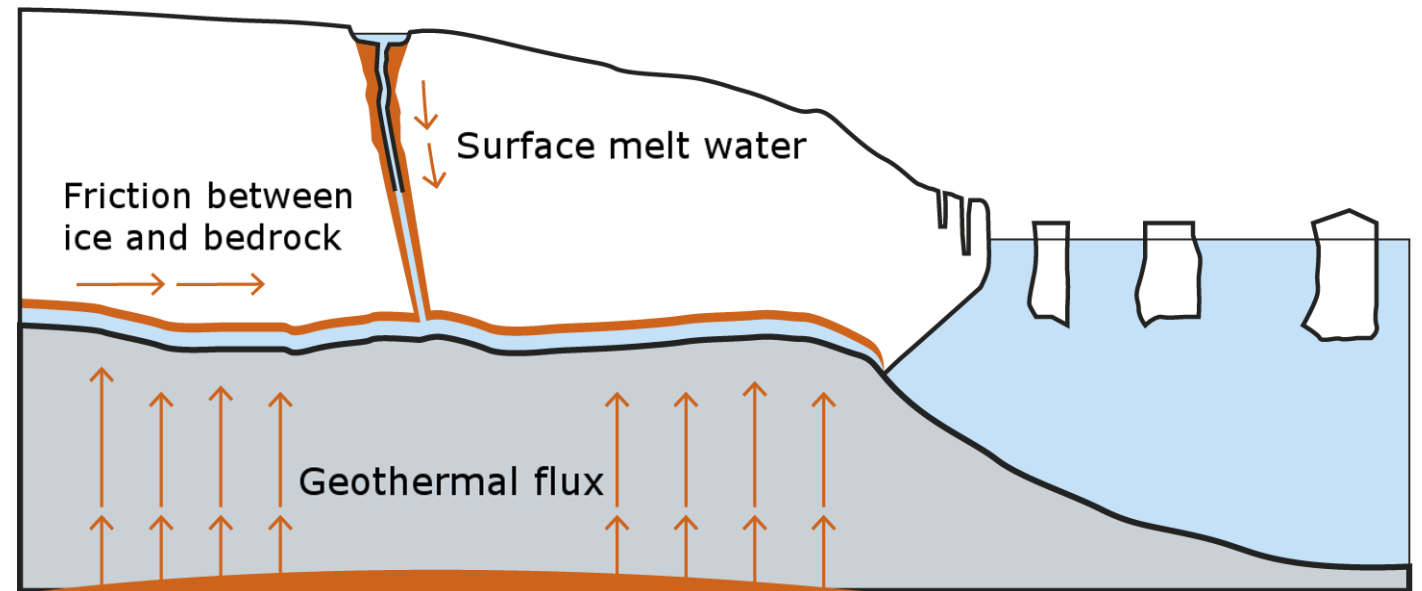
Sentinel-1 SAR image HV polarisation



Lake/no lake classification derived from dynamic threshold overlay on SAR image.



- Melt at the bed is very difficult to measure directly.
- Melt is generated by three heat sources:
  - Geothermal flux
  - Friction between ice and bedrock
  - Heat from surface melt water



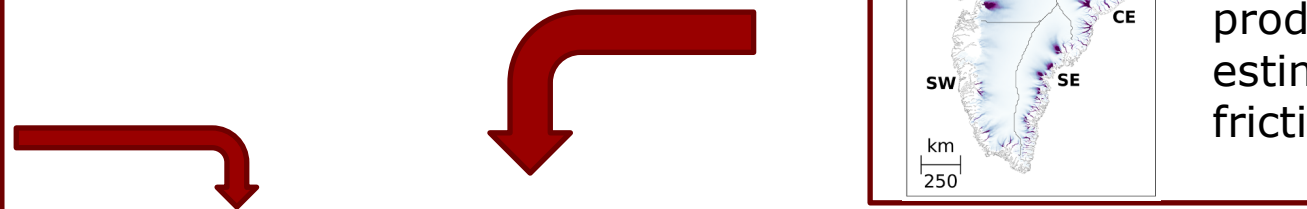
# Subglacial melt

Geothermal heat flux (mW per m<sup>2</sup>)

Geothermal flux:  
The average of three heat maps:  
Shapiro & Ritzwoller (2014), Fox Maule et al. (2009) and Martos et al. (2018)

Surface velocity (meter per year)

Elmer/Ice model (Gillet-Chaulet et al., 2012) tuned to assimilate observed surface velocities produces an estimate of friction heat.



Basal melt rate (meter per year)

Geothermal: 0 m per year

Friction: 0.01 m per year

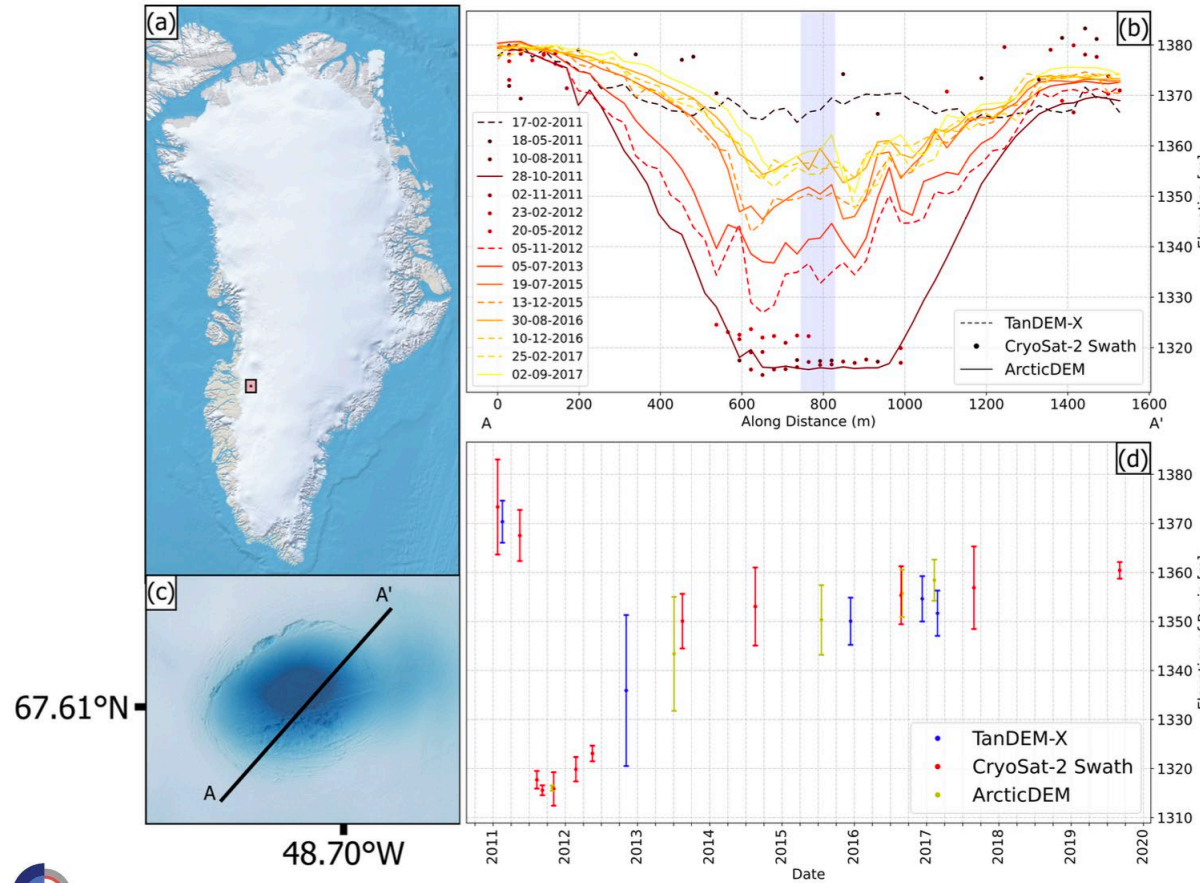
Surface melt water: 2000 m a.s.l.

Heat from surface meltwater (10<sup>6</sup> J)

Viscous heat from surface meltwater (Mankoff et al., 2017; 2020) is routed through the subglacial system.

# Active Subglaciale lakes

- Goal: Better constrain subglacial lake activity by CryoSat-2 swath processed data and TanDEM-X DEMs.



Four collapse basins found in Greenland are associated with drainage of subglacial lakes (Bowling et al., 2019).

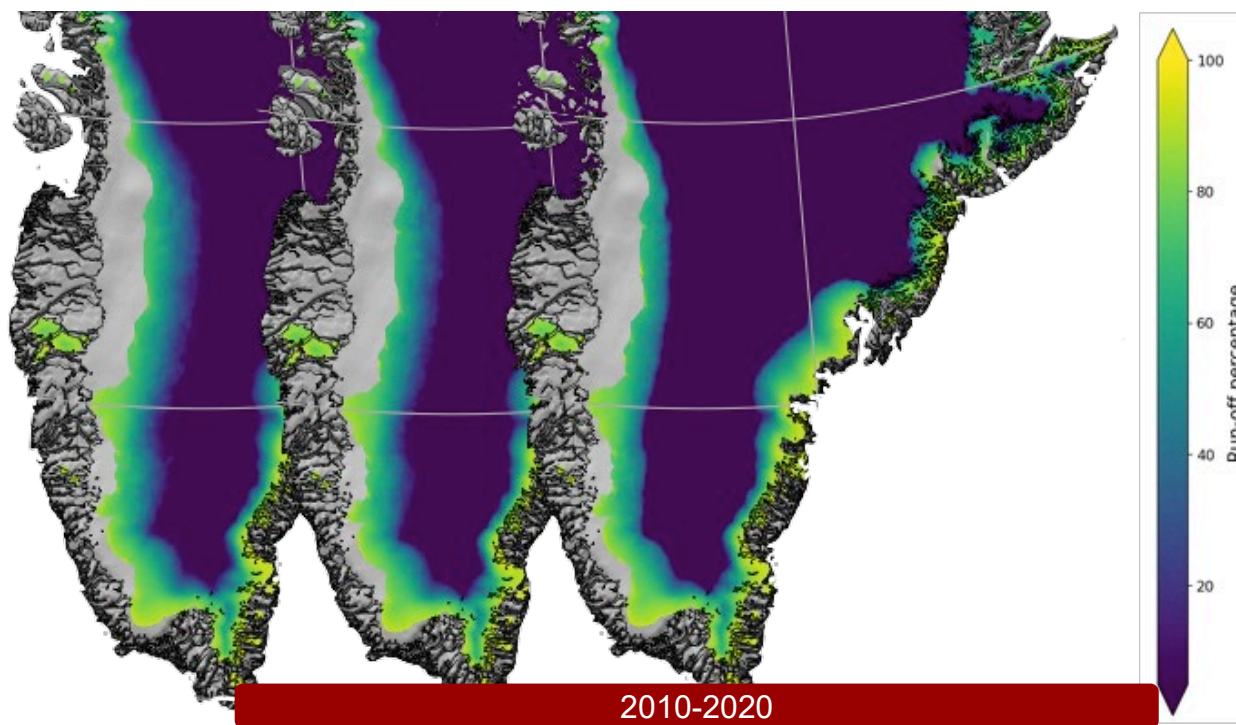
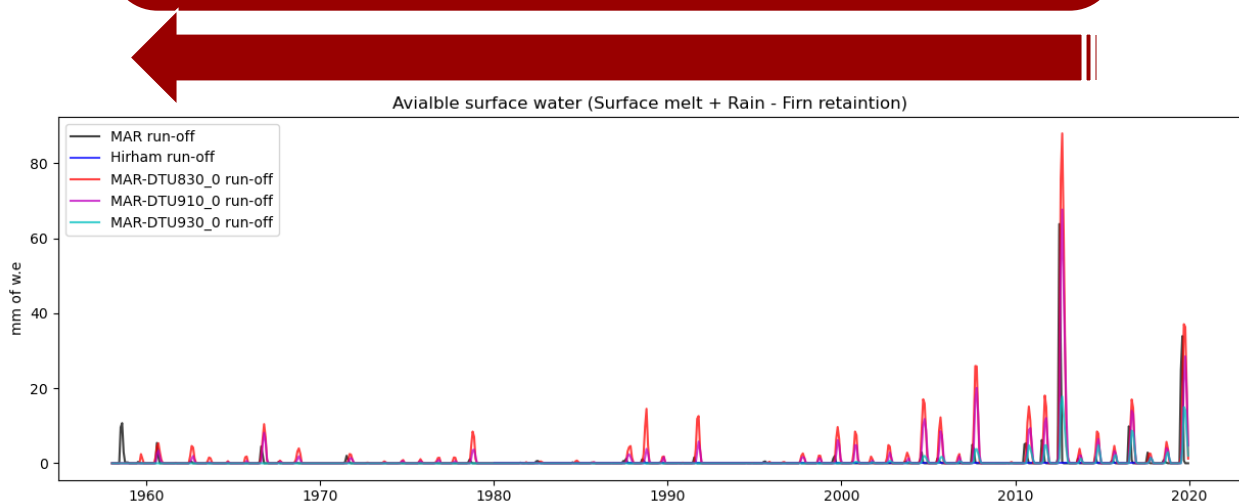
The addition of CryoSat-2 and TanDEM-X data helps to constrain the discharge event and the re-filling rate

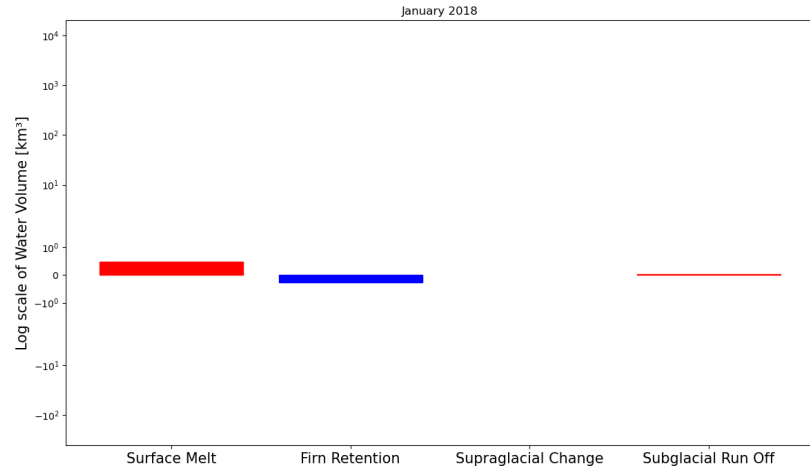
The improved DTU-firn model now provides updated estimates of meltwater retention and delayed.

Results depend on RCM used and percolation limit assumptions.

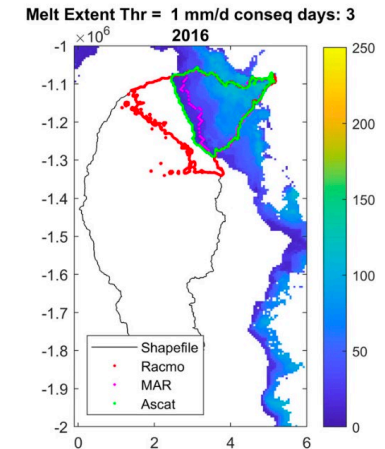
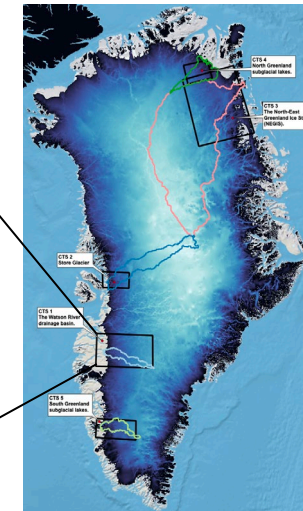
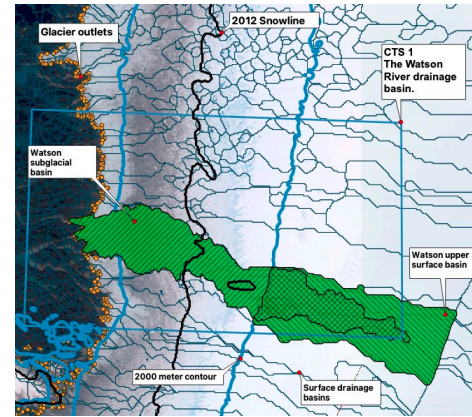
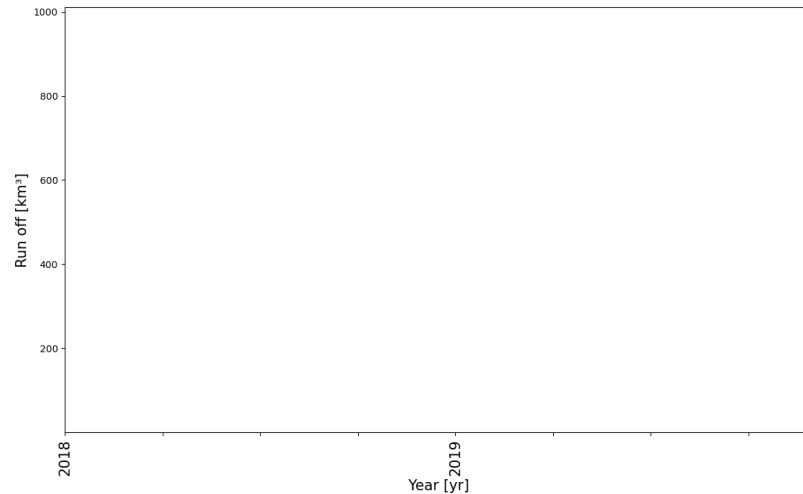
Run-off for the Watson area in time:  
 HIRHAM : 0.0%    MAR: 2.5%  
 DTU<sub>830</sub> : 9.0%    DTU<sub>910</sub> : 6.0%    DTU<sub>930</sub> : 1.8%

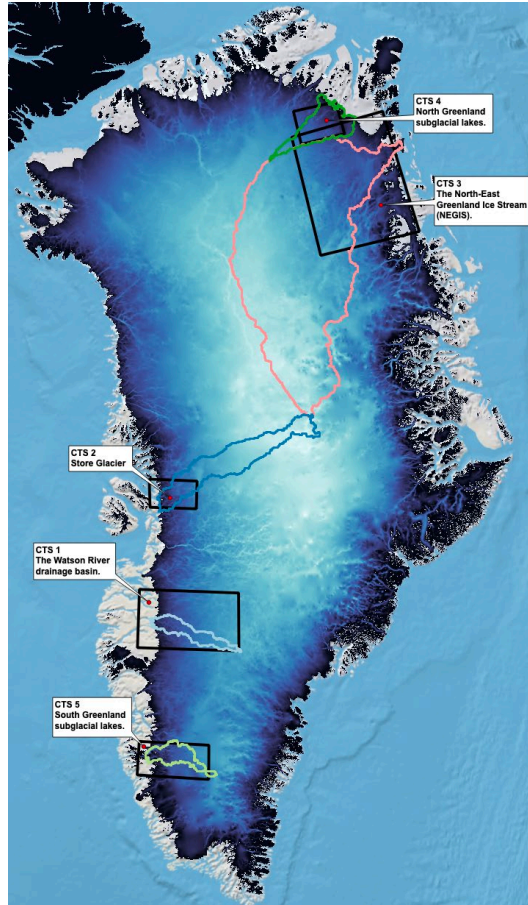
**Percolation limit**  
 830 kg/m<sup>3</sup>    910 kg/m<sup>3</sup>    > 917 kg/m<sup>3</sup>





- As a first attempt to make an integrated assessment we have focused on the Watson drainage basin in West Greenland 2018-2019.
- Combining all products to estimate monthly runoff:
  - Surface melt volume
  - Melt retained in the snow
  - Change in the meltwater volume in supraglacial lakes
  - Subglacial runoff





In June our experimental datasets will be made available.  
They are provided for our four test regions.  
Final data products ready by end of project (September 2022)

- **Activity 1** : Surface melt processes
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**Thank you for the attention.**

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