

living planet | BONN

23–27 May
2022

Impact of varying snow and ice properties of northern lakes on backscatter and brightness temperature measurements from altimetry missions: Results from the ESA LIAM project

C. Duguay^{1,2}, G. Picard³, E. Zakharova^{4,5}, J. Murfitt¹, M. Restano⁶, and J. Benveniste⁷

¹ University of Waterloo, Waterloo, Canada

² H2O Geomatics, Waterloo, Canada

³ Université Grenoble Alpes, Grenoble, France

⁴ EOLA, Toulouse, France

⁵ Water Problem Institute of RAS, Moscow, Russia

⁶ SERCO c/o ESA-ESRIN, Frascati, Italy

⁷ ESA-ESRIN, Frascati, Italy

Global Climate Observing System (GCOS) Requirements

ECV Lake Ice Thickness	
Measurement uncertainty	1-2 cm
Stability	N/A
Spatial resolution	100 m
Temporal resolution	Monthly

Revised/refined for
2022 GCOS
Implementation Plan
(currently under Public Review)

GCOS-200, 2016. The Global Observing System for Climate: Implementation Needs.

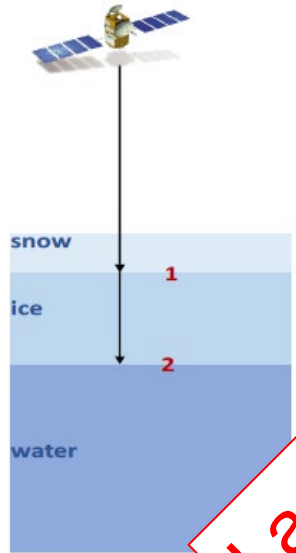
Climate Community Requirements

Parameter	Application	Horizontal Resolution	Observing Cycle	Precision	Accuracy	Stability
Lake Ice Thickness	Trend monitoring	< 200 m	Weekly	5 cm	5 cm	
	Seasonal / decadal forecasting	10 / 25 km	Weekly	10 cm	10 cm	
	Reanalyses	10 / 25 km	Daily	10 cm	10 cm	

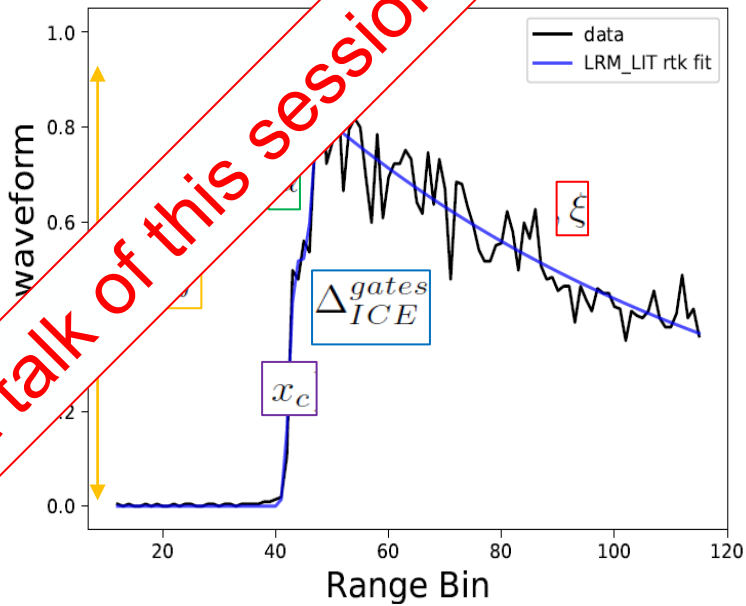
Table adapted from: CMUG CCI+, 2020. *Deliverable 1.1: Climate Community Requirements*, v2.2, 9 November 2020.

ESTIMATION OF LAKE ICE THICKNESS FROM ALTIMETRY MISSIONS

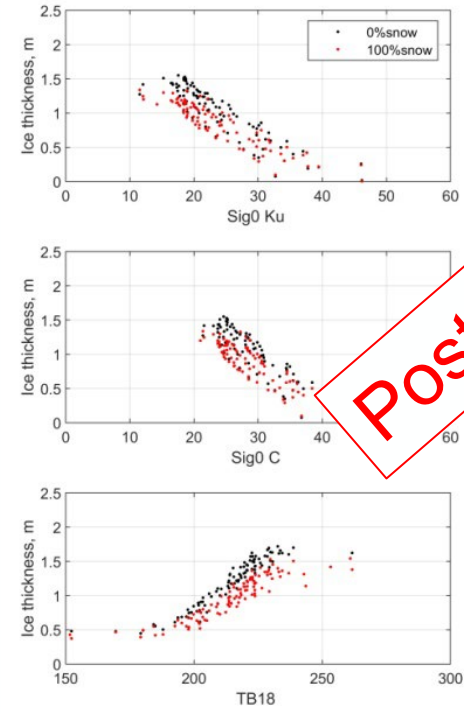
Physical/Analytical Approach¹



Last talk of this session!

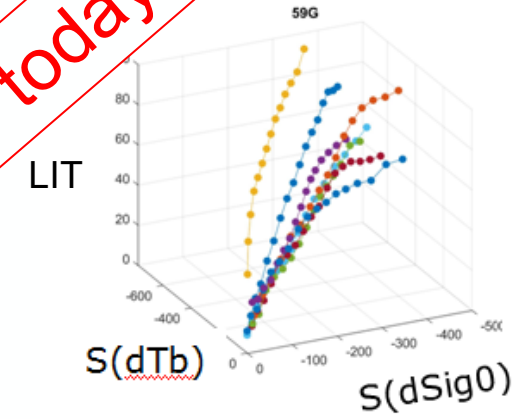


Empirical Approach²



Requires in situ measurements or modelled LIT

Poster today!



$$LIT = f(\text{Sig0}, T_b)$$

$$LIT = f(d\text{Sig0}/dt)$$

$$LIT = f(dT_b/dt)$$

¹ Mangilli, A., P. Thibaut, C.R. Duguay, and J. Murfitt, accepted. A new approach for the estimation of lake ice thickness from conventional radar altimetry. *IEEE Transactions on Geoscience and Remote Sensing*.

² Zakharova, E., S. Agafonova, C. Duguay, N. Frolova, and A. Kouraev, 2021. River ice phenology and thickness from satellite altimetry. Potential for climate studies and ice bridge road operation. *The Cryosphere*, 15: 5387-5407.

ICE AND SNOW PROPERTIES ENCOUNTERED ON NORTHERN LAKES

Ice with low and no snow



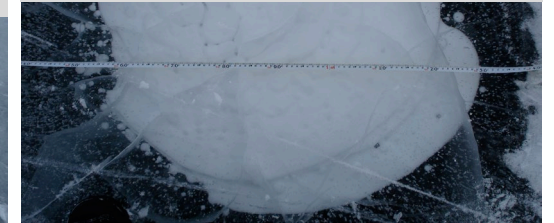
Roughness at ice-water interface



Clear ice / grey ice with small bubbles



Clear ice with large bubble



Snow on ice



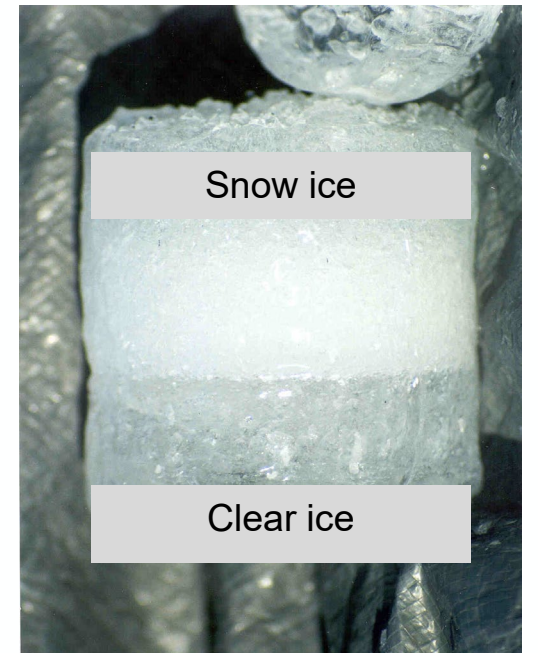
Pressure ridge
(2 m high, several km long)



Slushing / snow ice



Snow ice



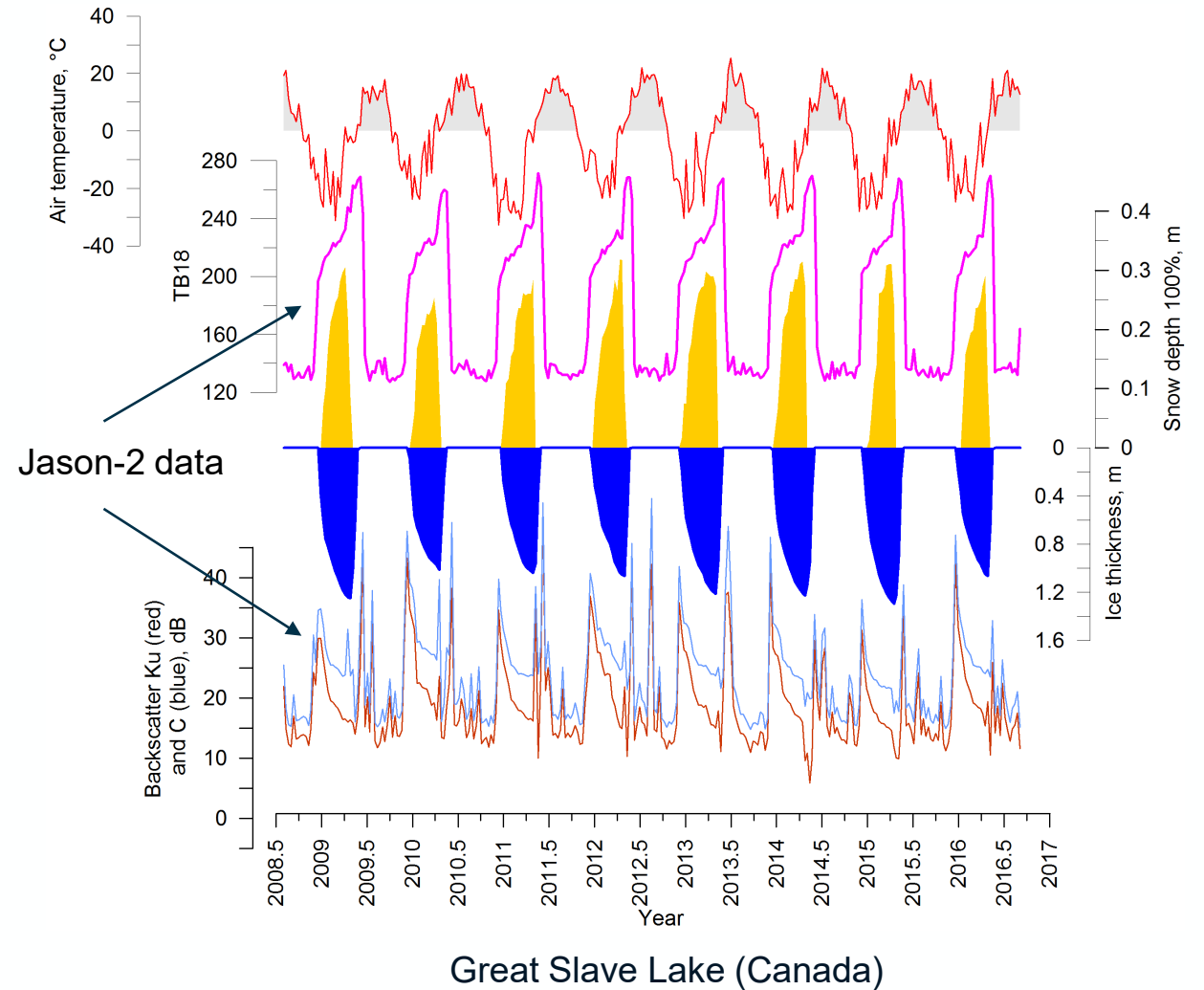
Clear ice

Objective

Examine the sensitivity of brightness temperature (Tb) and backscatter (σ°) measurements from altimetry missions to LIT of varied ice and overlying snow properties.

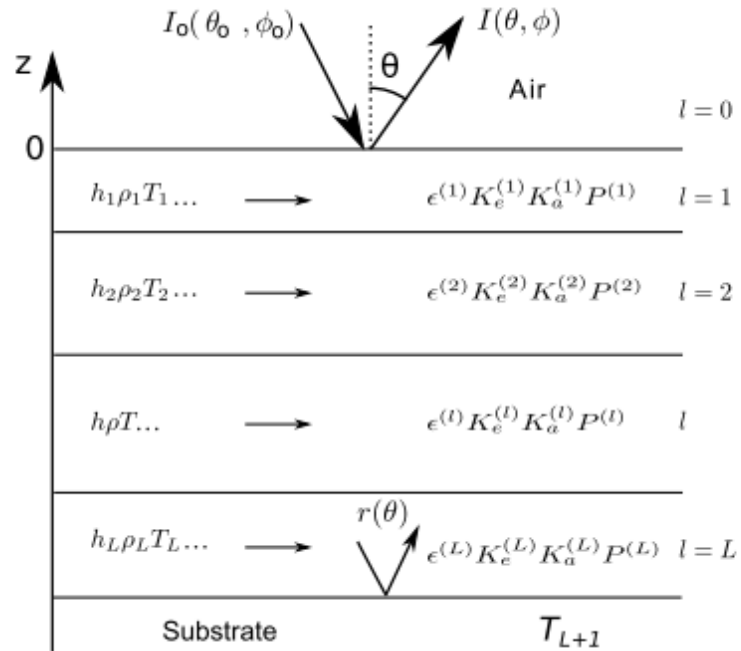
Approach

- Forward simulations using a radiative transfer model with input from a thermodynamic lake ice model.
- Analysis/interpretation of satellite measurements (Tb, σ° and waveforms).
- Comparison of forward simulations with satellite measurements.



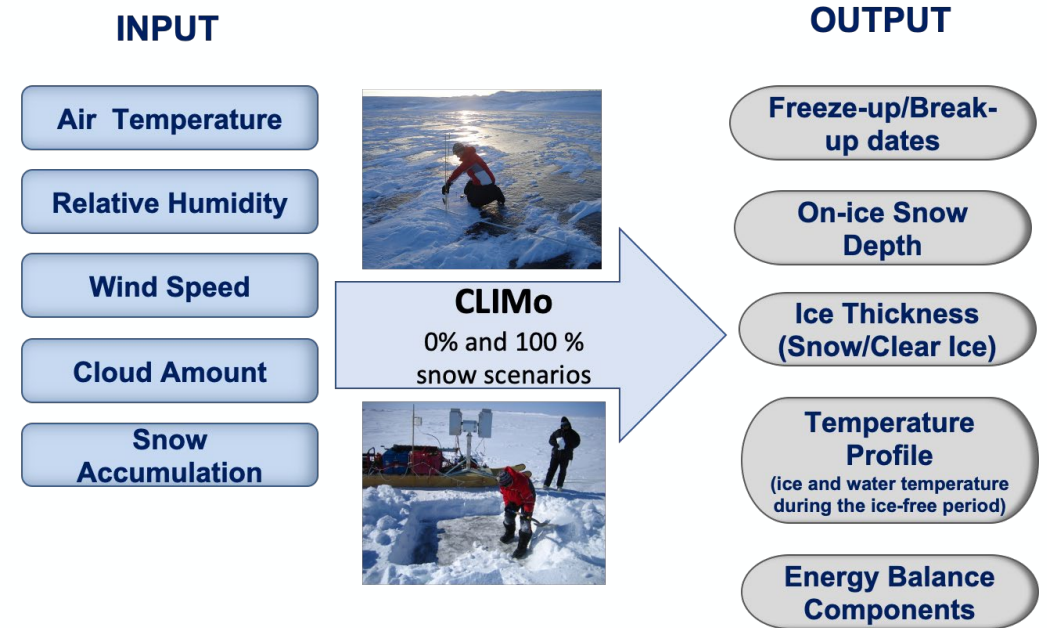
SMRT

Radiative transfer model that is usable for multilayer snowpacks and ice columns in both passive and active microwave



CLIMo

Thermodynamic lake ice model



SMRT: Picard, G., Sandells, M., Löwe, H., 2018. SMRT: An active-passive microwave radiative transfer model for snow with multiple microstructure and scattering formulations (v1.0). *Geoscientific Model Development*, 11: 2763–2788.

CLIMo: Duguay, C.R., G.M. Flato, M.O. Jeffries, P. Ménard, K. Morris, and W.R. Rouse, 2003. Ice cover variability on shallow lakes at high latitudes: Model simulations and observations. *Hydrological Processes*, 17(17): 3465-3483.

CLIMo and SMRT Integration

CLIMo Output

Temp _{surf}	Snow Depth (S _h)
Temp _{L1}	L1 _h
Temp _{L2}	L2 _h
Temp _{L3}	L3 _h
Temp _{L4}	L4 _h
Temp _{L5}	
Water	

SMRT Parameterization

Clear Ice Scenario

$Snow_T = (Temp_{surf} + Temp_{L1}) / 2$	S _h
Clear Ice	L1 _h = Ice _h / 4
$L1_T = (Temp_{L1} + Temp_{L2}) / 2$	
Clear Ice	L2 _h = Ice _h / 4
$L2_T = (Temp_{L2} + Temp_{L3}) / 2$	
Clear Ice	L3 _h = Ice _h / 4
$L3_T = (Temp_{L3} + Temp_{L4}) / 2$	
Clear Ice	L4 _h = Ice _h / 4
$L4_T = (Temp_{L4} + Temp_{L5}) / 2$	
Water	

Snow Ice Scenario

$Snow_T = (Temp_{surf} + Temp_{L1}) / 2$	S _h
Snow Ice	Snow Ice _h = Ice _h * 0.1
$L1_T = (Temp_{L1} + Temp_{L2}) / 2$	
Clear Ice	L1 _h = (Ice _h - Snow Ice _h) / 3
$L2_T = (Temp_{L2} + Temp_{L3}) / 2$	
Clear Ice	L2 _h = (Ice _h - Snow Ice _h) / 3
$L3_T = (Temp_{L3} + Temp_{L4}) / 2$	
Clear Ice	L3 _h = (Ice _h - Snow Ice _h) / 3
$L4_T = (Temp_{L4} + Temp_{L5}) / 2$	
Water	



Sensitivity Experiments

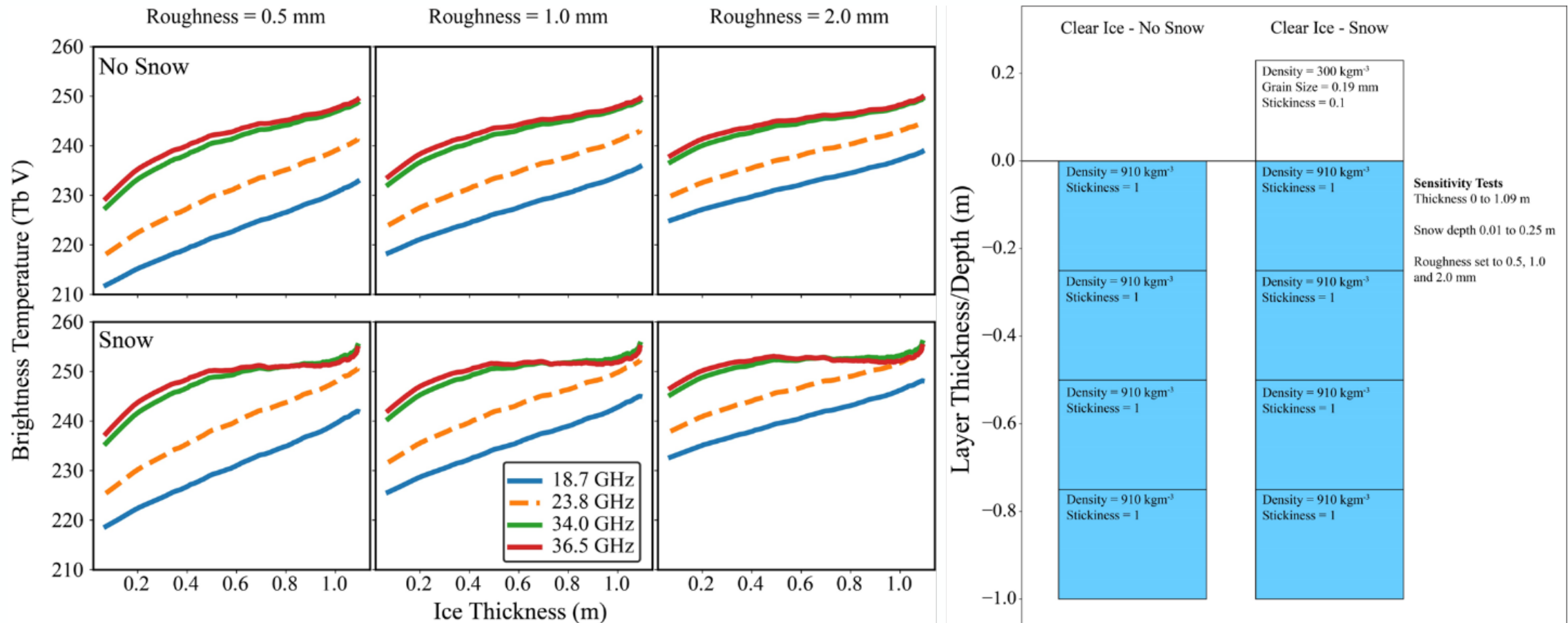
Simulations (3.2 to 36.5 GHz) of Tb and σ^o for various scenarios of ice (LIT, roughness at ice-water interface, clear ice, snow ice, bubble radius) and overlying snow (depth, density, wetness) properties.

Note: Being fundamentally a 1-D model, SMRT cannot account for 3-D features such as pressure ridges and cracks in the ice.

CLIMo inputs to SMRT are ice thickness, snow depth, ice temperature, and snow temperature.

Great Slave Lake (40-m mixed layer depth)

Tb simulations (2000-2020 average)

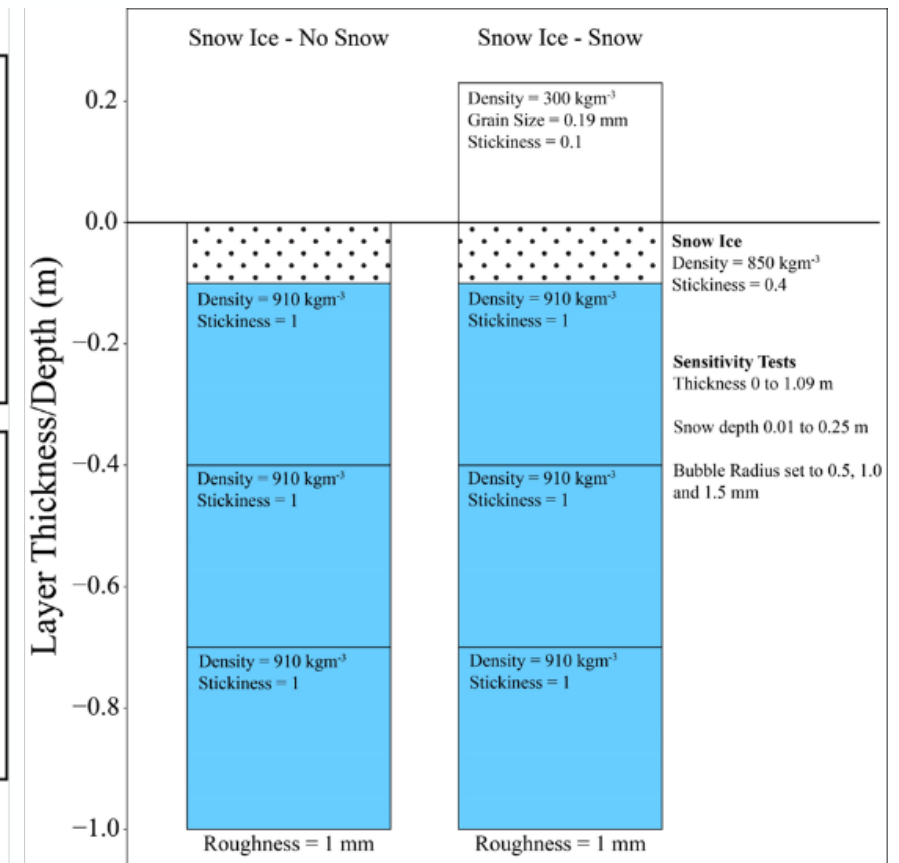
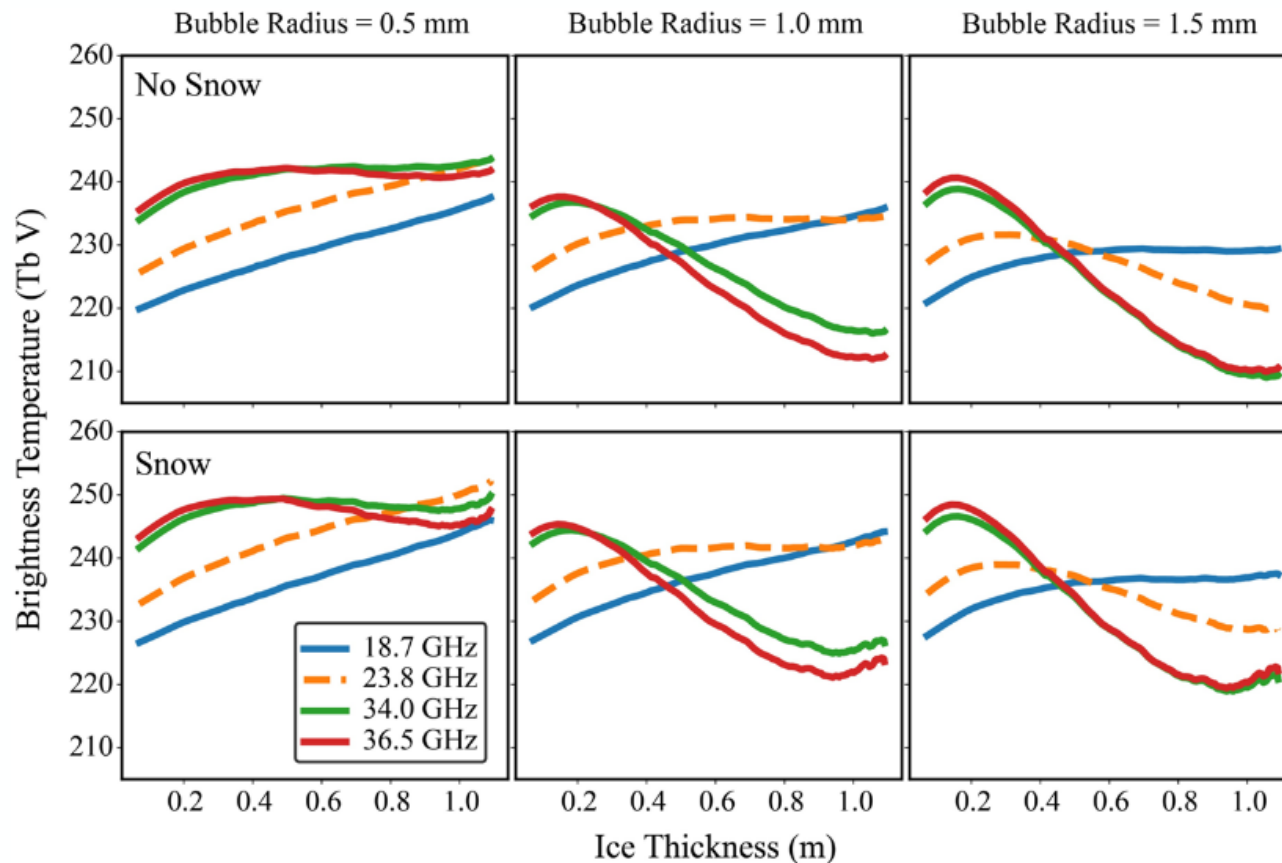


FORWARD SIMULATIONS: INFLUENCE OF SNOW ICE ON RELATION BETWEEN T_b AND ICE THICKNESS

Great Slave Lake (40-m mixed layer depth)

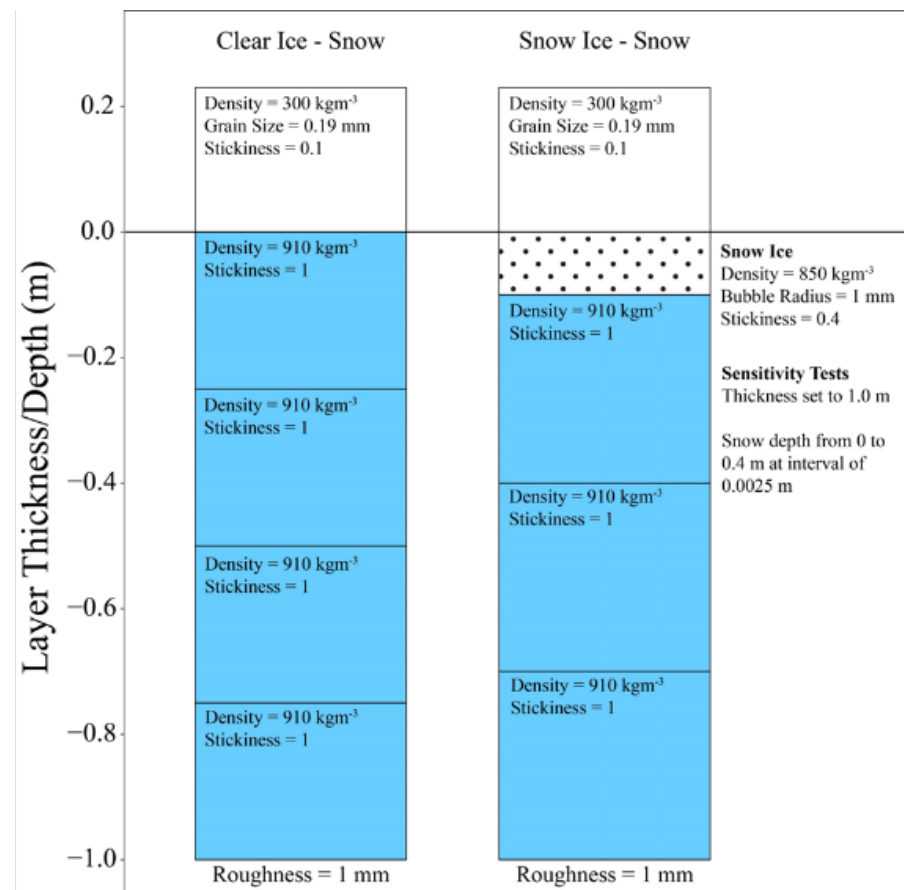
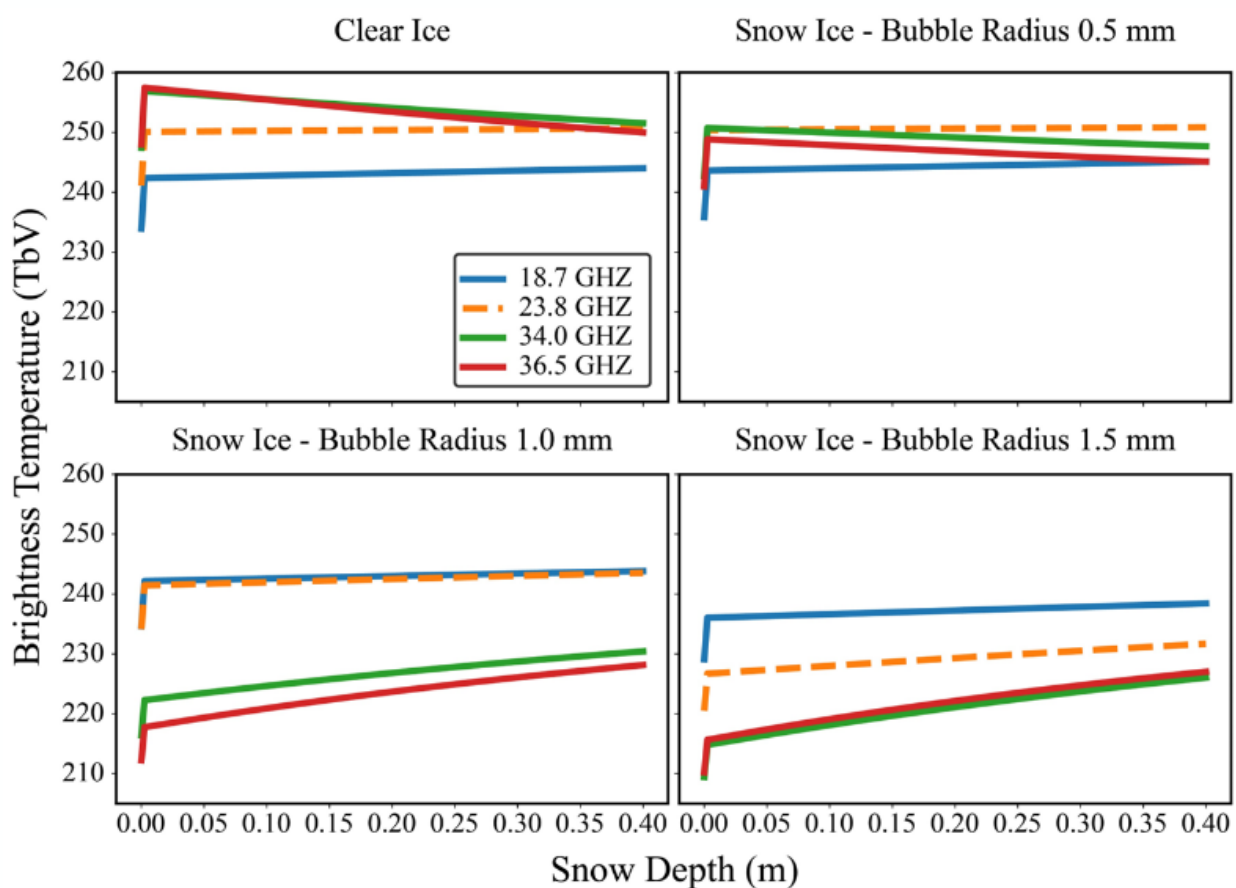
T_b simulations (2000-2020 average)

Snow ice set at 10% of the total ice column



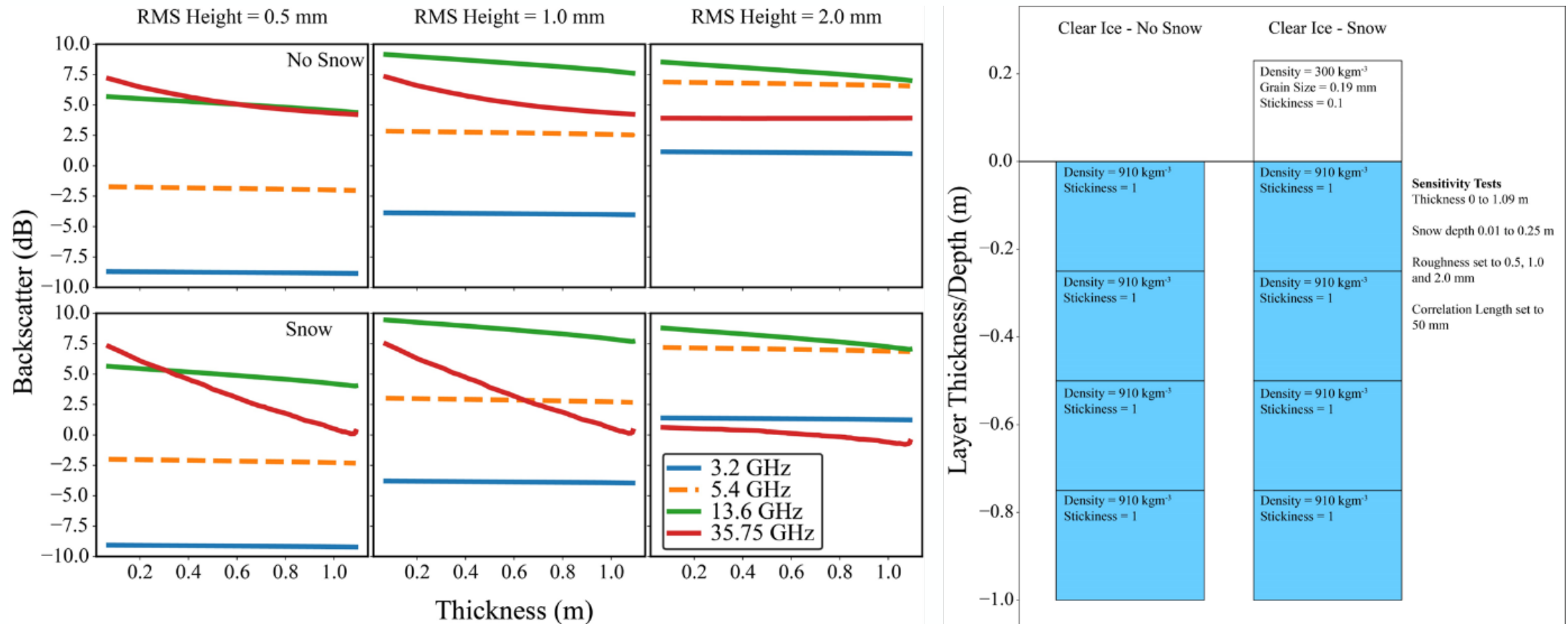
Great Slave Lake (40-m mixed layer depth)

T_b simulations (2000-2020 average)



Great Slave Lake (40-m mixed layer depth)

σ° simulations (2000-2020 average)

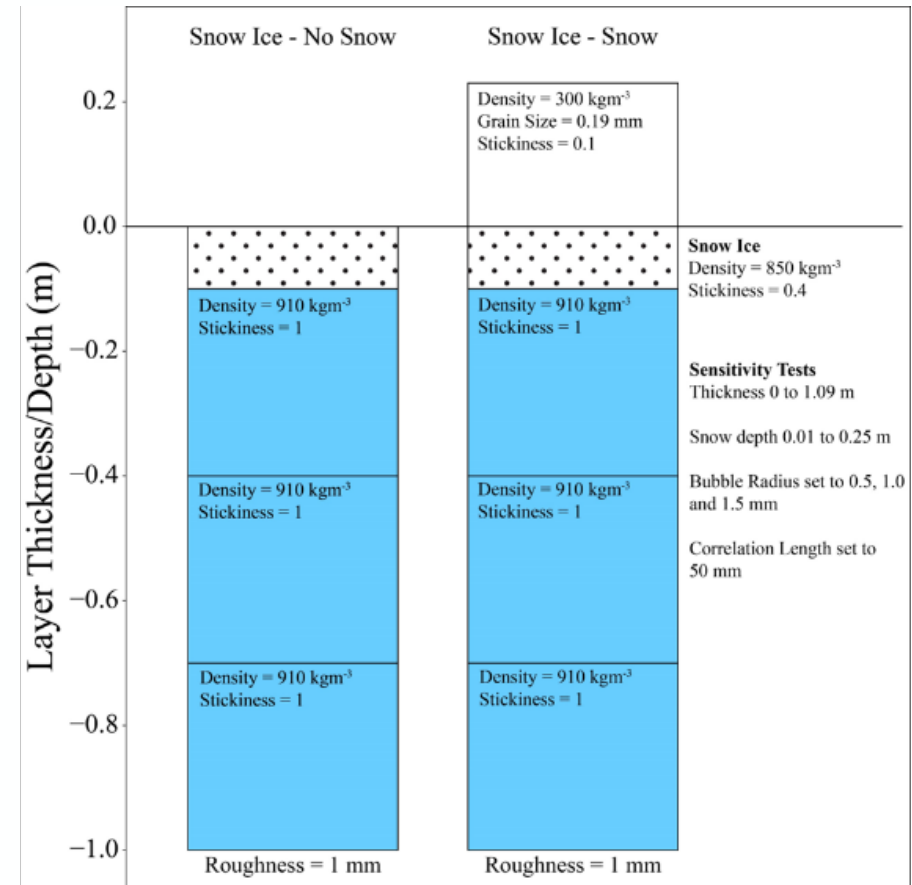
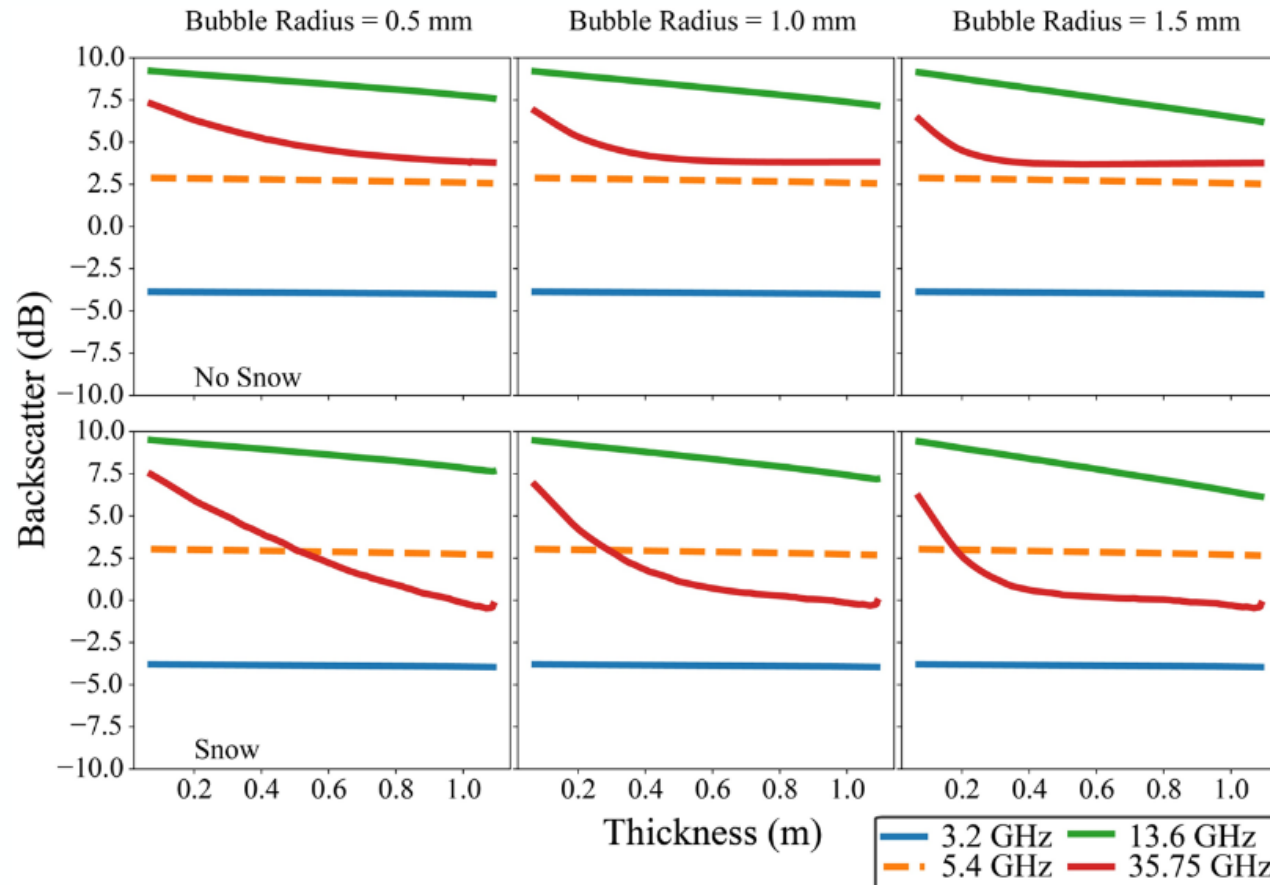


FORWARD SIMULATIONS: INFLUENCE OF SNOW ICE ON RELATION BETWEEN σ° AND ICE THICKNESS

Great Slave Lake (40-m mixed layer depth)

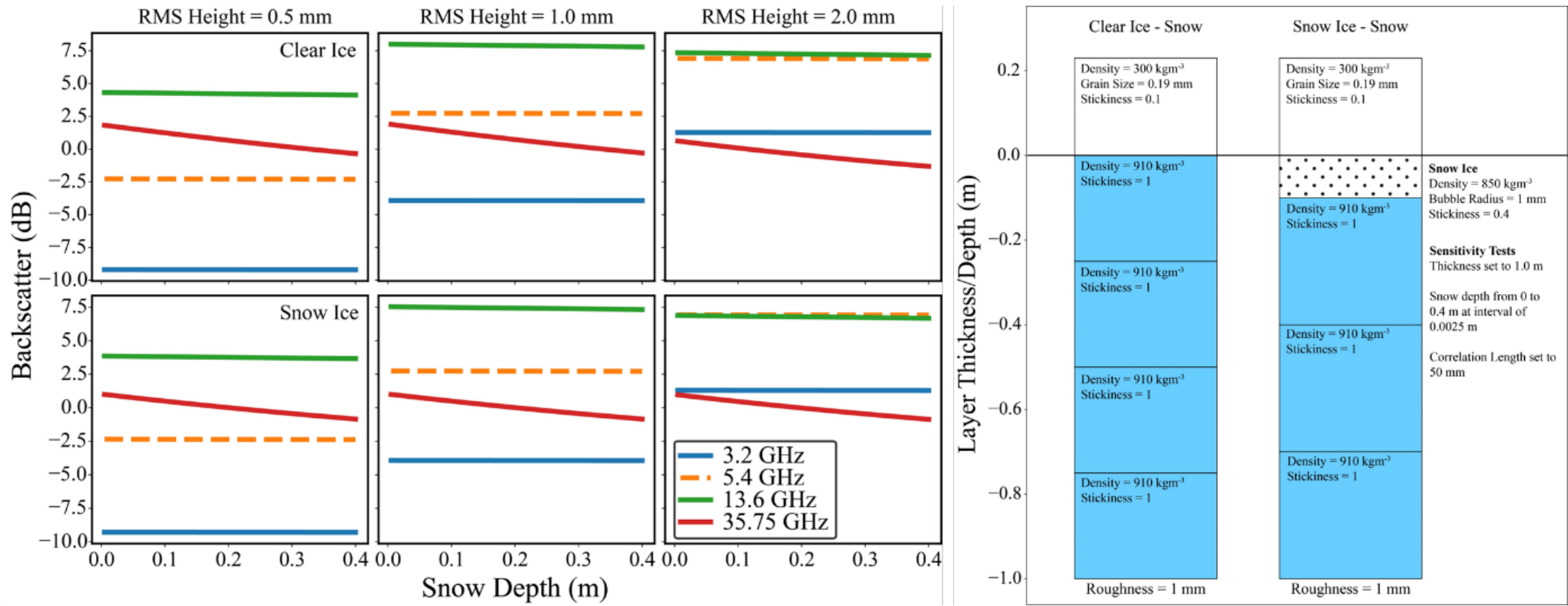
σ° simulations (2000-2020 average)

Snow ice set at 10% of the total ice column



Great Slave Lake (40-m mixed layer depth)

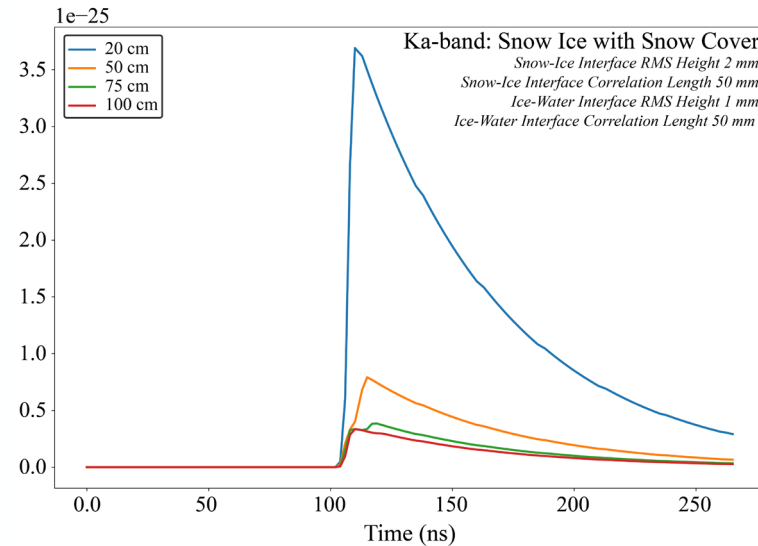
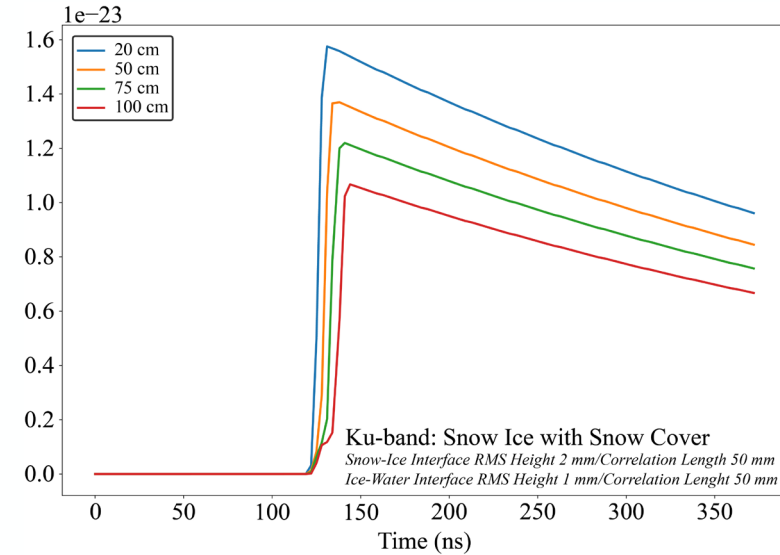
σ° simulations (2000-2020 average)



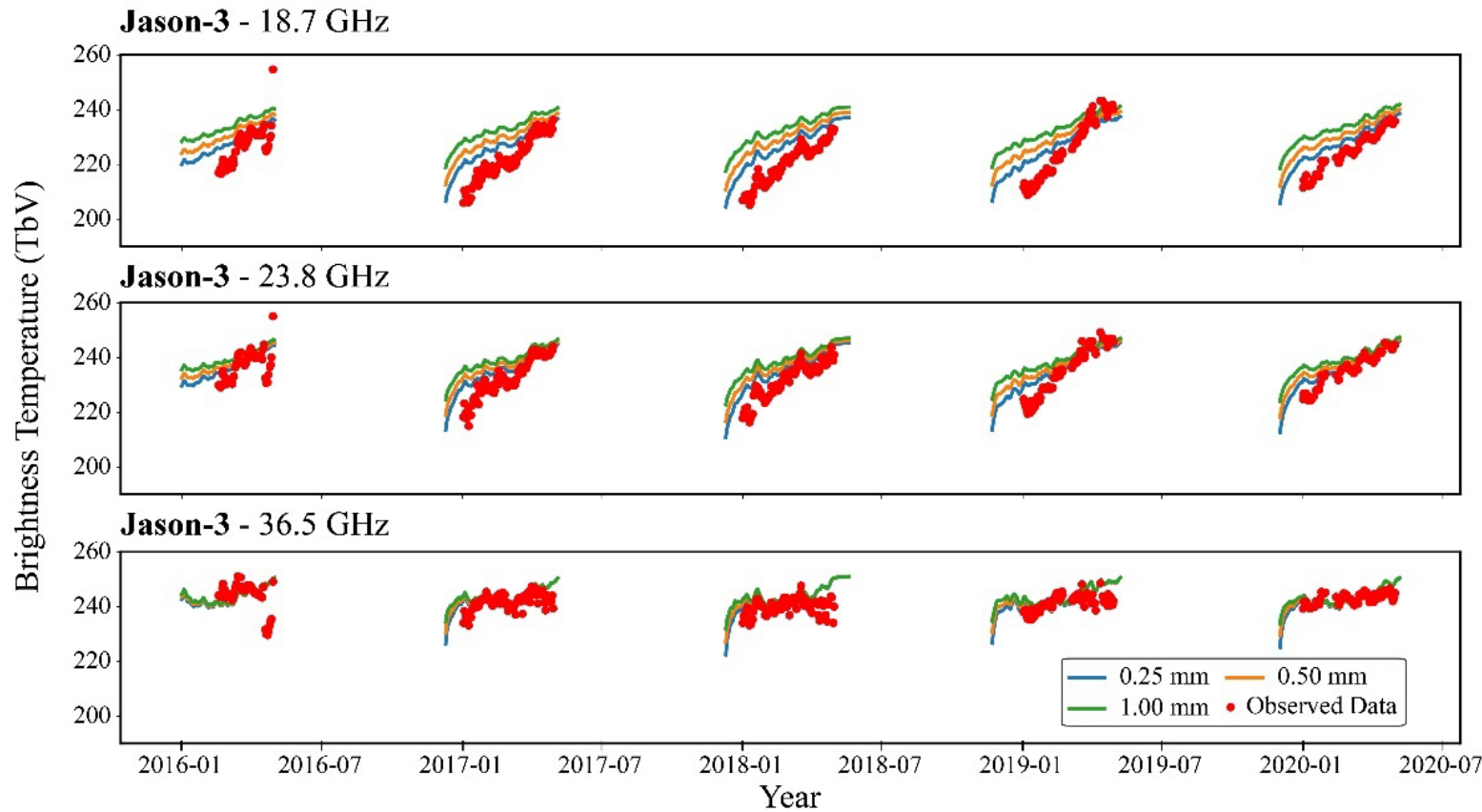
New “nadir LRM altimetry” module implemented in SMRT now being tested for lake ice

Paper describing the new extension applied to Antarctica:

Larue et al., 2021. Radar altimeter waveform simulations in Antarctica with the Snow Microwave Radiative Transfer Model (SMRT). *Remote Sensing of Environment*, 263, 112534.



Great Bear Lake (Centre)



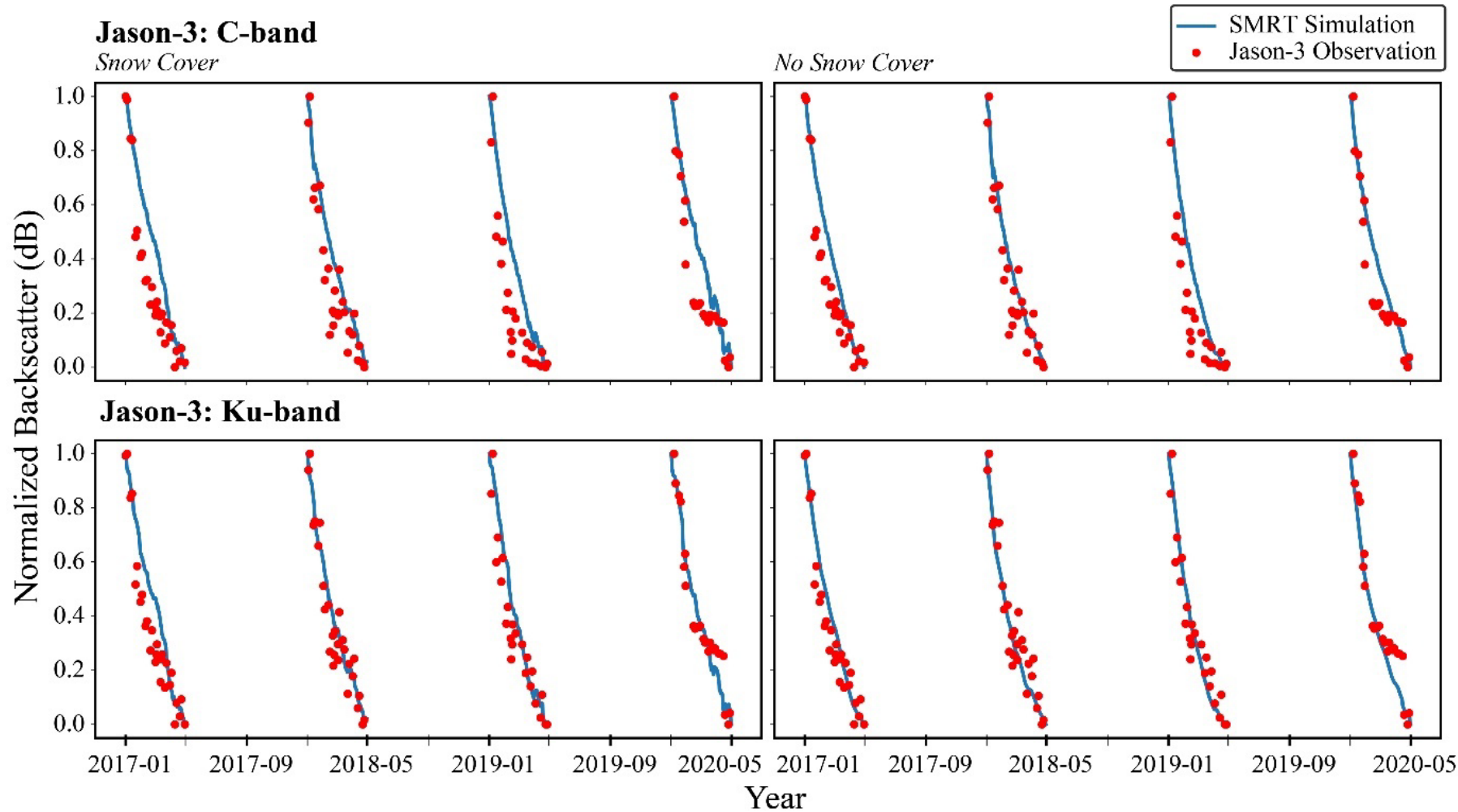
RMSE = 6.85
 $\rho = 0.98^*$

RMSE = 4.70
 $\rho = 0.95^*$

RMSE = 4.30
 $\rho = 0.39^*$

*Denotes significance $p < 0.05$

Great Bear Lake (Keith Arm)



$0.13 \leq \text{RMSE} \leq 0.17$
 $\rho = 0.77^*$

$0.08 \leq \text{RMSE} \leq 0.11$
 $0.78^* \leq \rho \leq 0.79^*$

*Denotes significance $p < 0.05$

- A thermodynamic lake ice model (CLIMo) has now been “coupled” with SMRT, which allows for more realistic (beyond synthetic) simulations of the impact of varied ice and snow properties with ice growth on brightness temperature, backscatter and waveforms.
 - The new “nadir LRM altimetry” module needs to be further tested and compared to altimeter waveforms.
- The quality of LIT retrievals from physical/analytical and empirical approaches can be affected by ice and overlying snow properties/conditions in addition to ice thickness.
- Sensitivity experiments further support the use of 13.6 GHz radar (Ku-band) and 18-19 GHz radiometer data for the retrieval of LIT and the potential of Ka-band (ca. 34-37 GHz range) used alone or in combination with Ku-band for the retrieval of on-ice snow depth.
 - These results have implications for future work investigating the retrieval of LIT and snow depth from the upcoming SWOT and CRISTAL missions.
- Extension of SMRT to high-resolution SAR altimetry mode is needed for the simulation of waveforms from current (CryoSat-2, Sentinel-3, Sentinel-6) and future missions (SWOT, CRISTAL).

living planet symposium | BONN 23–27 May 2022

TAKING THE PULSE
OF OUR PLANET FROM SPACE



THANK YOU FOR YOUR ATTENTION!

LIAM Project website:
<https://www.h2ogeomatics.com/lake-ice-from-altimetry-missions-li>