

# Uncertainty in evapotranspiration mapping from thermal infrared remote sensing data with EVASPA

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Evapotranspiration (ET) can be derived using various models based on thermal infrared data

```
    -> uncertainty in Ts: {
        - instrument
        - atmosphere
    -> large number of models
        -> diversity of algorithms => uncertainties
    -> lots of other data required :
        - albedo
        - vegetation density
        - weteorological data
        -> various sources of data => uncertainties
```

Many unknowns remain concerning the uncertainties in the derivation of ET, in particular for discriminating uncertainties from input data and models

## **Models of Latent Heat Flux (LE)**

[see Lagouarde and Boulet 2016]

-> Contextual model : LE ~ EF x (Rn – G)

```
      EF = evaporative fraction <- Ts vs. albedo or NDVI

      Rn = net radiation <-</td>
      albedo, emissivity, Ts, solar irradiance atmospheric irradiance

      G = ground heat flux <- Rn, NDVI, fCOVER
```



ex: S-SEBI (Roerink et al. 2000)





Uncertainty analysis : we use the EVASPA concept (Gallego et al. 2013, Olioso et al. 2018, Allies et al. 2020)

-> Uncertainty is defined as the standard deviation of the ensemble simulations (or the range of variations as in Mira et al., RSE 2016)

-> Uncertainty related to the impact of one variable or model is computed by keeping only the variations related to this model or variable

-> We compute two levels of uncertainty: (Allen et al., 2011, Blatchford et al., RSE 2019).

## **Example -> Meteorological variables : solar radiation**



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## **Example -> Meteorological variables : solar radiation**



## Uncertainties

# variability in input data source and algorithms

		Novice case			Expert case				
	common inputs								
Surface temperature	Ts	4 2 atmospheric corrections (				Un. Valencia & CzechGlobe)			
	NDVI	2 calculations (different spectral windows)							
	albedo	13 models			11 models				
Incident solar radiation	emissivity	20 models				5 models			
	<u>R</u> g	2 reanalysis 1 satellite pr		oduct 2 in situ stations		1 reanalysis	1 satellite product		1 in situ station
Atmospheric radiation	<i>L</i> ↓	2 reanalysis		2 in situ stations		2 reanalysis		2 in situ stations	
G to Rn ratio	ξ (G/Rn)	7 equations				6 equations			
	Evaporative fraction model								
Evaporative fraction	EF	5 algorithms				5 algorithms with slope limitations			

#### Surface temperature data from TASI

Two sets of data were provided based on the TES algorithm

```
We used data on the 18<sup>th</sup> and the 20<sup>th</sup> at : 

- around 10h UTC (= 12h CEST) -> morning

- around 12h UTC (= 14h CEST) -> midday

notation: D1H1, D1H2, D2H1, D2H2
```

However, results were very similar for each days (and even each acquisition time) -> D2H2

## **Uncertainty in Ts**

## Standard deviation of the 4 Ts maps



Standard deviation Ts (K)

## Ts standard deviation map D2H2 (K)



#### **Evapotranspiration map**

## Average of all the calculations for D2H2

## Novice case – average ET =3.4 mm $d^{-1}$



#### Expert case - average ET = 2.8 mm d<sup>-1</sup>







## **Uncertainty in ET** (mm d<sup>-1</sup>) D2H2 case





#### ET uncertainty break-down, D2H2

#### **Summary**

- uncertainty in ET is large, up to 1.5 mm.d<sup>-1</sup> (for D2H2 and when expressed as standard deviation)
- uncertainty is significantly lower in the expert case than in the novice case
- ranking of uncertainty sources highlights

- impact of Ts is low
 - largest impacts: evaporative fraction, G/Rn ratio (ξ)
 - model formulations have a larger impact than input data

#### Perspectives

- work on other situations
- extending to aerodynamic one-source and two source-models (TSEB, SPARSE...)
- transfer to time serie processing (ex. using MODIS data)
- derive an uncertainty algorithm to be used in the TRISHNA data processing for associating uncertainty to ET

## **Ranking of uncertainty sources**



Uncertainty ranking for D2H1 (July 20<sup>th</sup> morning)