

Chasing small amplitudes signals from the deep Earth's interior in GRACE measurements

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Origin of the GRACEFUL project



GRACEFUL is an ERC Synergy project dedicated to detect deep Earth's signals in satellite observations of the magnetic field, gravity field and rotation

Objectives: explore fluid motions in the **outer core** and cristallisation/dissolution processes at the **core mantle boundary**

Motivation: discovery of spatio-temporal correlations between the magnetic field and gravity field at interannual time-scales by Mandea et al., 2012 (PNAS) and 2015 (JGR)



Challenges for the analysis of the gravity field

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Processes causing changes at various temporal and spatial scales in the Earth's gravity field (Mandea et al., 2012)

Challenge:

Gravity signals originating from the deep Earth interior are masked by shallower geophysical processes, including the:

- global water cycle
- glacial isostatic adjustment
- earthquakes

What we need to do :

- correct GRACE and GRACE-FO observations for surface effects
- estimate the uncertainties on satellite gravity data and surface corrections.



Satellite gravity data processing



Collection of time-lapse satellite gravity data from GRACE and GRACE Follow-On

Average of monthly surface mass anomalies from 3 mascon (JPL, CSR, GSFC) and 6 spherical harmonic (JPL, CSR, GFZ, ITSG, COST-G, CNES-GRGS) solutions from April 2002 to December 2021

Application of geophysical corrections

- Glacial isostatic adjustment : ICE6G-D
- Ocean-Atmosphere circulation : AOD1B Release 6 (GAD)
- Hydrology : ISBA-CTRIP and WGHM
- > Additional correction for lakes based on satellite altimetry from hydroweb (Blazquez et al., in prep)

Processing of residual surface mass anomalies

- > Application of geographical masks over ice-sheets and regions impacted by large earthquakes
- > Focus on interannual time scales: removal of linear trend, annual-sinusoid and semi-annual sinusoid
- > Application of a spatial filter (radius: 250 km) based on diffusion (Goux et al., submitted to GJI)

Amplitude of interannual gravity signals







Amplitude of detrended and deseasoned mean residual mass anomalies



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2002 2003 2004 2005 2006 2007 2008 2009 2010 2011 2012 2013 2014 2015 2016 2017

Time

















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🕤 The case of the Zambezi river basin

Amplitude of detrended and deseasoned mean residual mass anomalies



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The case of the Zambezi river basin





The case of the Zambezi river basin





The case of the Zambezi river basin





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Amplitude of detrended and deseasoned mean residual mass anomalies

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The case of the Congo basin





The case of the Congo basin

The case of the Congo basin

Exploration of residual signal in GRACE data

What do residual mass variations look like in "quiet" regions of the world?

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Dominant signal in residual surface mass anomalies

EOF 1 (GRACE - GAD - ISBA): 10.45 percent

1.00 0.75 0.50 0.25 0.00 -0.25-0.50-0.75-1.001.00 0.75 0.50 0.25 0.00 -0.25-0.50-0.75-1.00

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Dominant signal in residual **geoid** anomalies

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Flow at the Core Mantle Boundary

Bouguer anomaly: 1cm EWH \rightarrow 420 nGal

Gravity change at the surface (right) associated with the flow at the CMB in 2000 (left).

The flow model is the ensemble average of the flow models in Barrois et al. (2017) truncated at degree 11

From Dumberry & Mandea (2021) Surveys in Geophysics <u>10.1007/s10712-021-09656-2</u>

Sea level fingerprints associated with the loss of ice from glaciers and ice sheets and from changes in terrestrial water storage.

Hsu & Velicogna, 2017 (GRL)

<u>https://doi.org/10.1002/2</u> <u>017GL074070</u>

- The GRACE and GRACE-FO have been corrected for surface effects. The hydrological models ISBA-CTRIP and WGHM struggle to predict pluri-annual to decadal variations in the water storage.
- In "quiet" regions, residual mass anomalies display characteristics that would be expected from the deep Earth's interior: amplitudes < 2 cm, very large scale features (1 000 to 10 000 km), remarkable continuity between the oceans and the continents, typical time-scales from a few years to 1.5 decade.</p>
- > The origin of such signals remains unknown: large-scale climate or deep Earth's interior?

Dominant spatial feature in residual geoid anomalies from GRACE

Surface gravity change associated with flow at the CMB

Sea level fingerprint associated with ice-melting and TWS changes

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Improve the estimation of residual mass anomalies from satellite gravity data

- Inclusion of additional models (hydrology) and corrections (geocenter, C20)
- Search for optimal corrections, models and parameters

Improve the estimation of the uncertainties

- Estimation of the dispersion of the ensemble of residual mass anomalies
- Search for advanced signal characterization method (aka: go beyond eofs!)

> Identify the mechanisms responsible for the residual gravity field variations

- Comparison with satellite observations of the magnetic field and length of day
- Investigation of deep Earth's dynamic processes by comparison with numerical models
- Investigation of "slow" climate forcings: sea level fingerprints, wind patterns, climate indices...

Thank you for your attention.

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https://graceful.oma.be/

Second component of residual surface mass anomalies

1.00

0.75

0.50

0.25

0.00

-0.25

-0.50

-0.75

-1.00

1.00

0.75

0.50

-0.25 -0.00 --0.25

-0.50

-0.75

-1.00

Third component of residual surface mass anomalies

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ISBA-CTRIP:

- ISBA solves the water and energy balance in the soil, canopy, snow and surface water bodies.
- **CTRIP simulates runoff** through the global river channel network.

ISBA and C-TRIP are coupled through the land surface interface SURFEX, allowing interactions with floodplains and groundwater, modelled with a 2D diffusive aquifer.

Schematic view of ISBA-CTRIP (Decharme et al., 2019)

Hydrological models

Watergap Global Hydrological Model:

precipi-

Eact

soil

tation P 、

canopy

throughfall

T > 0°C

canopy evaporation

snow

runoff R

R_s

surface water

groundwater

use

Ra

aroundwater

fon ingaio

sublimation

vertica

water

local lakes

balance

- WGHM solves a **simple vertical water balance**, except for runoff which allows lateral flows from one cell to another
- WGHM takes into account human influence on the freshwater storage variability through dams building, water abstraction and irrigation
- Water storage in glaciers is included in version 2.2d of WGHM

withdrawals

return flows

P↓ ↑E_{pot}

local wetlands

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inflow from

upstream cells

global lakes/reservoirs

global wetlands

river segment

P. TEpot

P. TEpot

Concept of the GRACE and GRACE-FO missions

- GRACE : April 2002 October 2017
- GRACE Follow On : May 2018 present
- Two identical satellites flying 220 km apart
- Polar orbit ~ 500 km above the Earth
- Provides time-lapse estimate of the Earth's gravity field
- Accurate range (inter-satellite distance) rate measurement of by microwave ranging system (KBR: 1 μm per second) and laser range interferometry (LRI: 1 nm per second)

Credit: NASA/JPL-Caltech/GFZ & Tapley et al., 2019, Nature Climate Change

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Residual geoid anomalies

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Residual geoid anomalies

2 .

-2

Power

bC2 (mm)

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😚 Residual geoid anomalies

Residual geoid anomalies

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- -0.6

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Flow at CMB in 2000

Example of a flow map at the CMB for the year 2000.

The flow model is the ensemble average of the flow models in Barrois et al. (2017) truncated at degree 11

From Dumberry & Mandea (2021)

Surveys in Geophysics 10.1007/s10712-021-09656-2

Geostrophic pressure, vertical displacement at the CMB and at the surface, and gravity change at the surface associated with the flow at the CMB in 2000.

Contribution of subannual, pluri-annual and decadal signals in detrended and deseasoned GRACE residuals corrected for ISBA

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Residual hydrology signals

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Contribution of subannual, pluri-annual and decadal signals in detrended and deseasoned GRACE residuals corrected for WGHM

