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TAKING THE PULSE
OF OUR PLANET FROM SPACE



Response of GNSS reflected signals to wave spectra across cyclones

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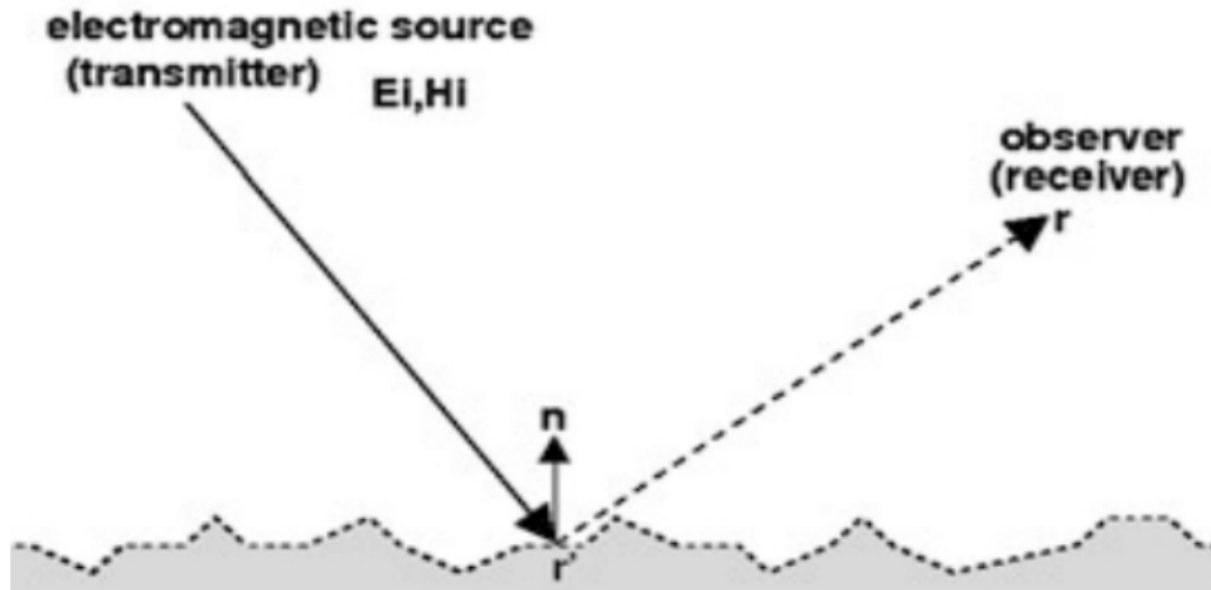
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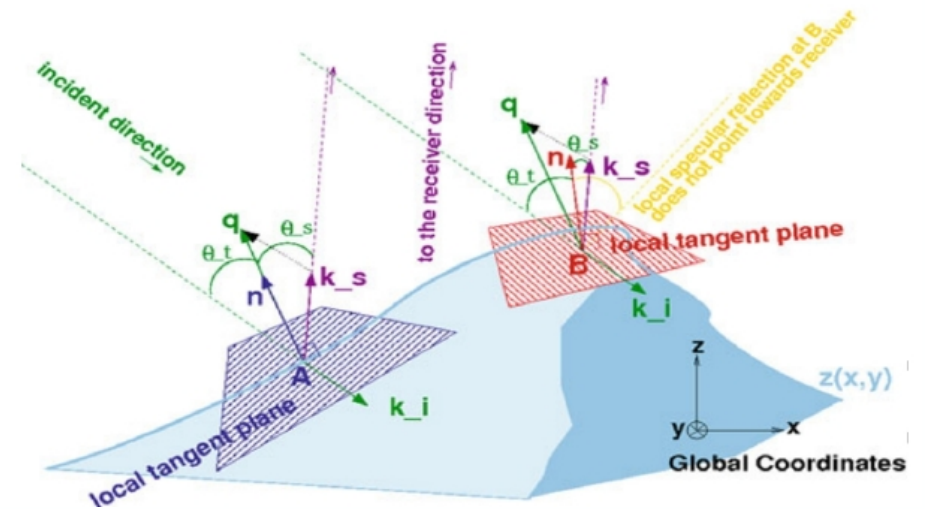
- On GNSS-Reflectometry (GNSS-R)
- Question
- Suggestion
- Methodology
- Case study: GNSS-R Data
- Case study: 2D Wave Model
- Preliminary results

- Reflectometry using satellite navigation signals (GNSS-R): bistatic forward scattering at L-band ($\lambda_{EM} \sim 19 \text{ cm}$)

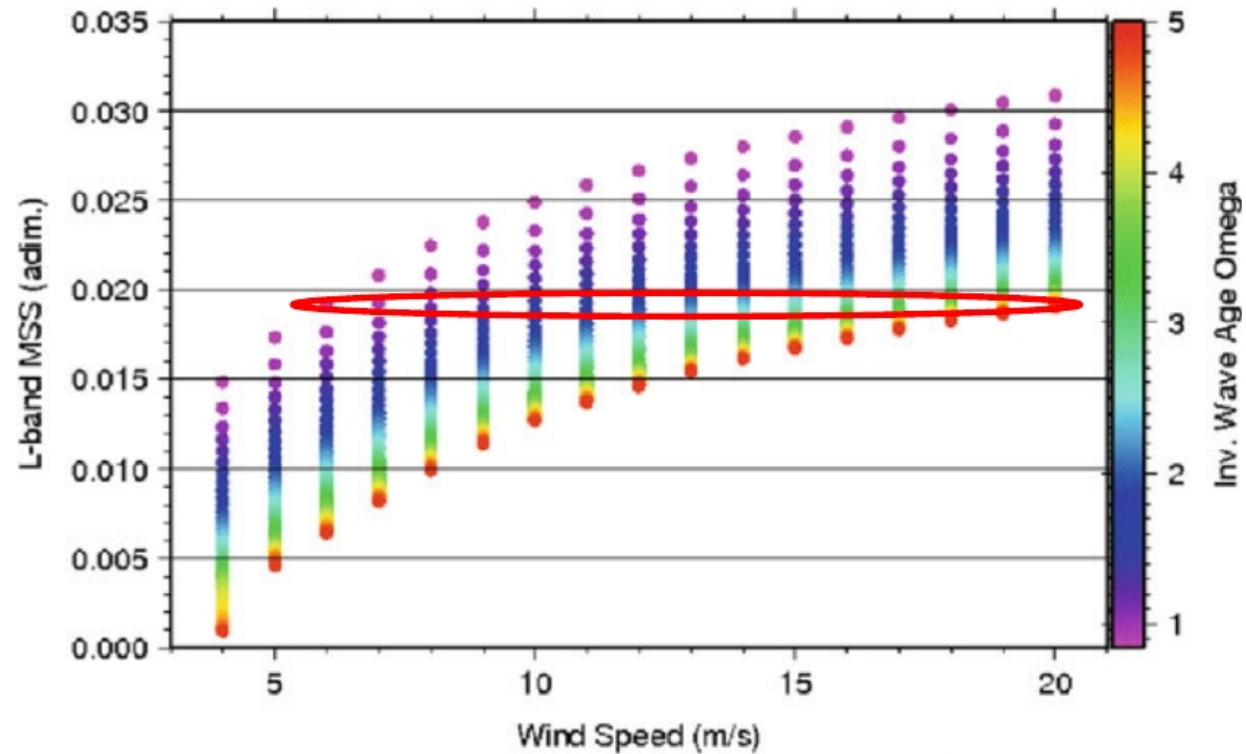
GNSS satellite



GNSS-R signals are **sensitive to L-band filtered surface roughness** ($\lambda_{sea} \geq 3 \lambda_{GNSS} \sim 0.6\text{m}$).



- The statistics of this filtered roughness, through **mean squared slopes (mss)**, depend on **combination** of different variables: wind, fetch, waves age, swell...



Using Elfouhaily et al., 1997 spectrum

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- In tropical cyclones, **complex wave structure**: different wave trains of different fetches/ages juxtapose, non-linear effects...
- Does it make sense to **'invert'** the observables into one of the combined variables (e.g. wind speed)?

QUESTION:

How to use GNSS-R observables around tropical cyclones?

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- Here we suggest to **avoid 'inversion' schemes** (GNSS-R depends on several variables).

Instead, perhaps the **assimilation of the 'observable'** into complex wave models could get the most of them



Note that the 'observable' is not an oceanographic variable

- To assimilate an observable, such as σ_0 , into numerical models, the observable must be expressed as a **forward operator of the state variables \mathbf{x}** of the model. $\sigma_{mod}^0 = H[\vec{\mathbf{x}}]$
- A variational approach to data assimilation then consists of minimization of a cost function J :

$$J(\vec{\mathbf{x}}) = \frac{1}{2}(\vec{\mathbf{x}} - \vec{\mathbf{x}}_B)^T \mathbf{B}^{-1}(\vec{\mathbf{x}} - \vec{\mathbf{x}}_B) + \frac{1}{2}(\sigma_{obs}^0 - \sigma_{mod}^0(\vec{\mathbf{x}}))^T (\mathbf{E} + \mathbf{F})^{-1}(\sigma_{obs}^0 - \sigma_{mod}^0(\vec{\mathbf{x}}))$$

Covariance of the numerical ocean model

Covariance of the observations

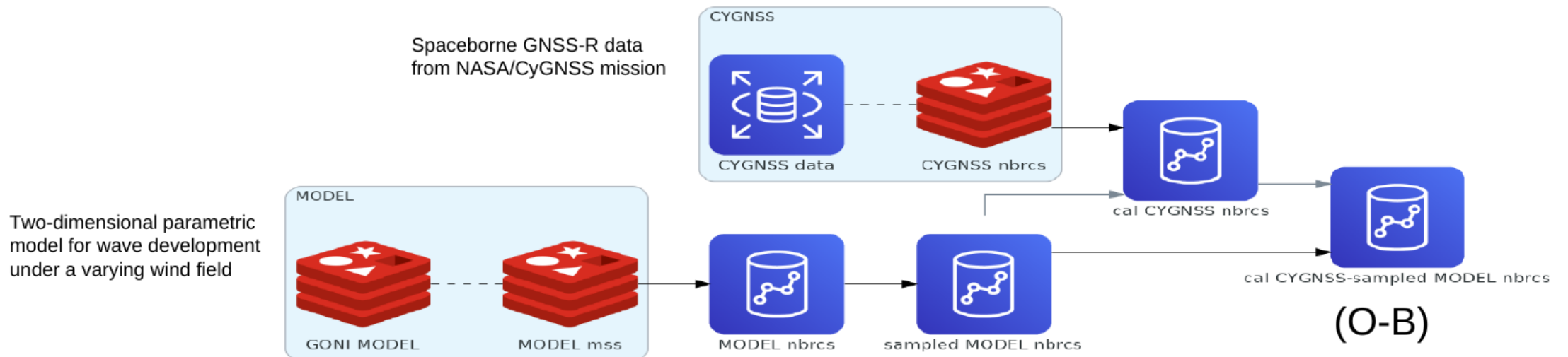
Covariance of the forward operator $H[\mathbf{x}]$

The approach corrects anomalies not originally captured in the numerical model: pre-fit residuals or (O-B).

Certain degree of agreement required for the approach to converge.

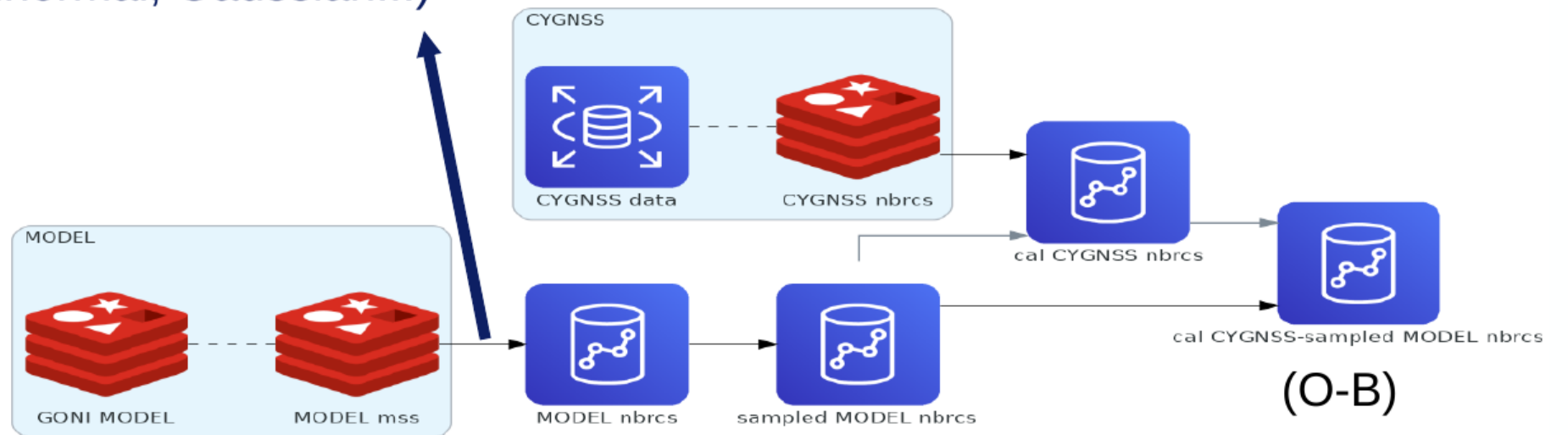
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- The goal of this initial exercise is to check whether spaceborne GNSS-R observables, σ_0 , across TC are sufficiently close to the equivalent σ_0 , estimated by a complex wave model.
- This is the first and required condition for future assimilation of σ_0 in such type of models.
- A case example is presented: Goni typhoon.

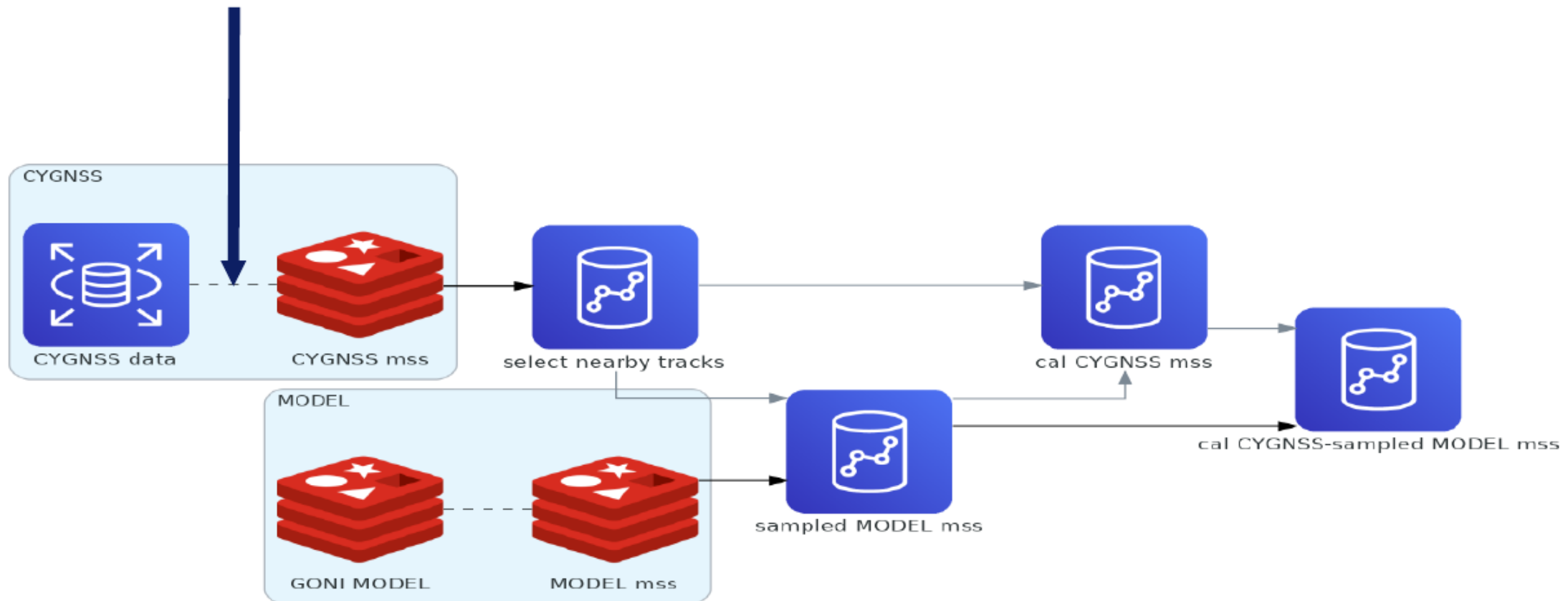


- Why not assimilating GNSS-R measured mss?

PDF(slopes): the model has all required information (e.g., Gram-Charlier, binormal, Gaussian...)



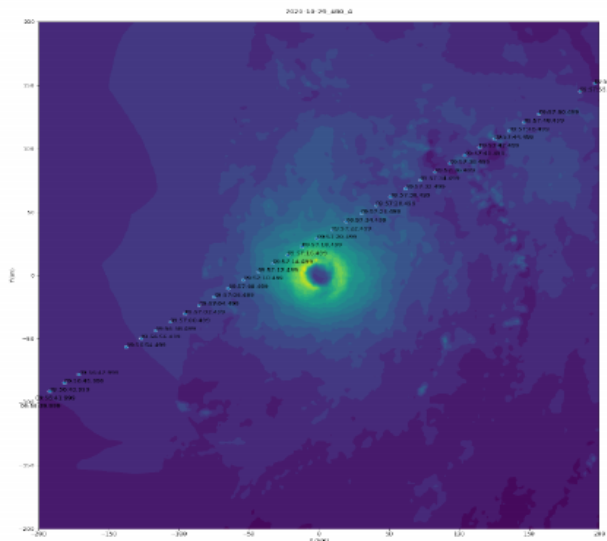
PDF(slopes): assumed by the 'data provider' in the inversion procedure → Gaussian



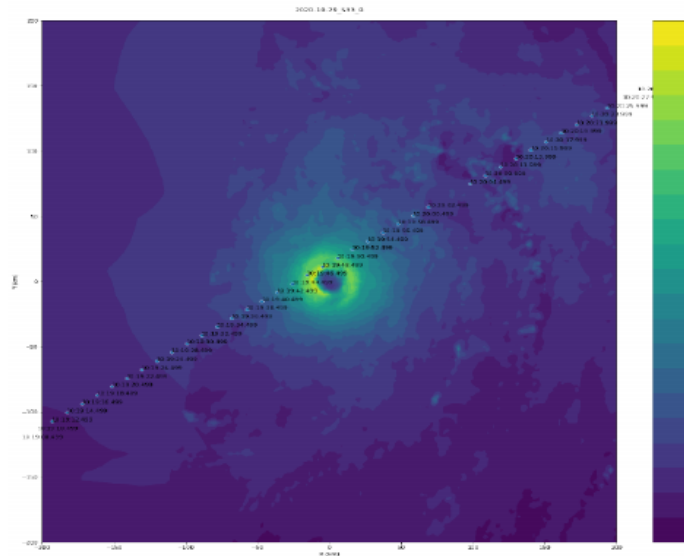
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- Case study for Goni Typhoon 2020-10-29
- **GNSS-R data:**
 - Level 1b NBRCS (σ_0) NASA/CyGNSS
 - 3 tracks crossing Goni on 2020-10-29/30

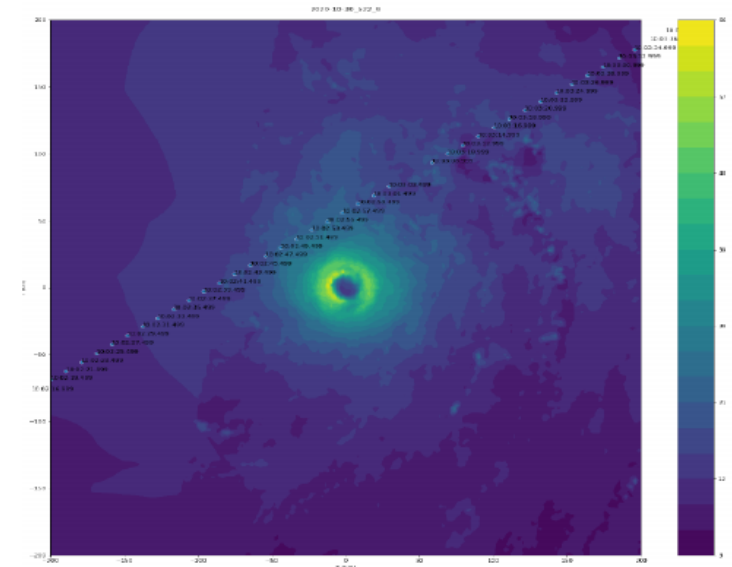
2020-10-29_480_4



2020-10-29_533_8



2020-10-30_522_8



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Two-dimensional parametric model for wave development under varying wind field

(Kudryavtsev, Yurovskaya & Chapron, 2021)

- The system describes the development of surface waves under a varying wind field in both space and time, and the evolution of swell propagation in the absence of wind forcing.
- Input: 2D wind field at 1 km resolution from Sentinel-1 SAR images (provided by IFREMER).
- Governing Equations:

Energy and momentum conservation
(Hasselmann et al., 1976; Phillips, 1977):

$$\frac{\partial E}{\partial t} + c_{gj} \frac{\partial E}{\partial x_j} = S^E$$

$$\frac{\partial M_i}{\partial t} + c_{gj} \frac{\partial M_i}{\partial x_j} = S_i^M$$

$$\int \int d\varphi d\omega$$

and some algebra...

$E(\omega, \varphi) = A(\varphi - \varphi_p) F(\omega)$ - energy spectral density

$M_i = \kappa_i E / \omega = \kappa_i \omega E / g$ - momentum spectral density

$\kappa_i = [\cos \varphi, \sin \varphi]$

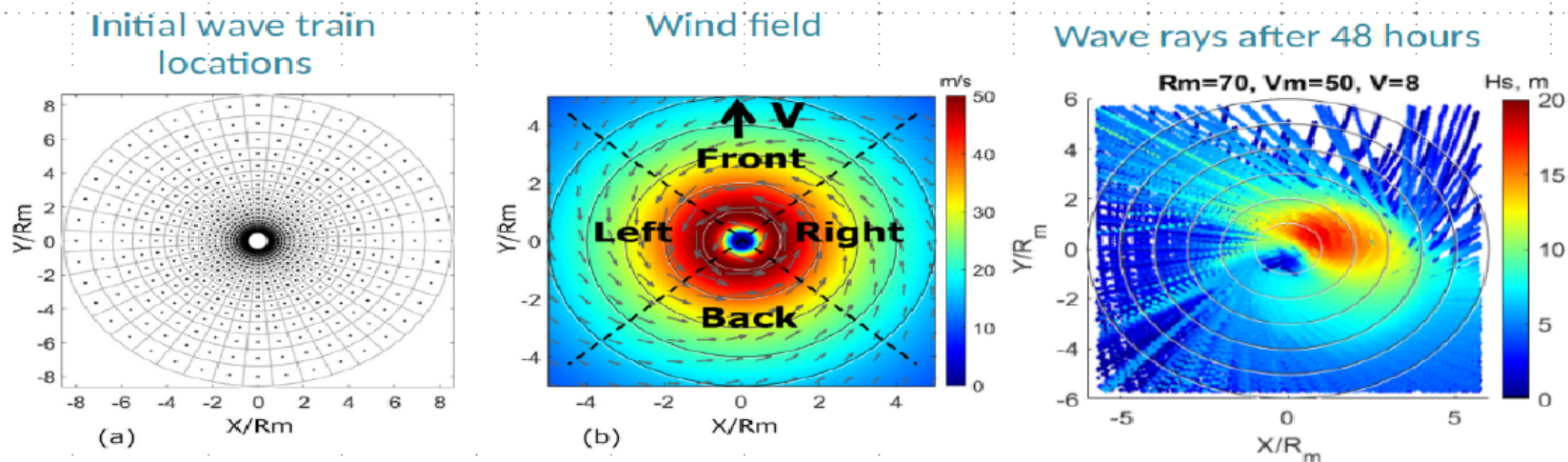
$S^E = S_w - S_D + S_N$ - energy source

$S_i^M = \kappa_i \omega S^E / g$ - momentum source

c_{gj} - group velocity

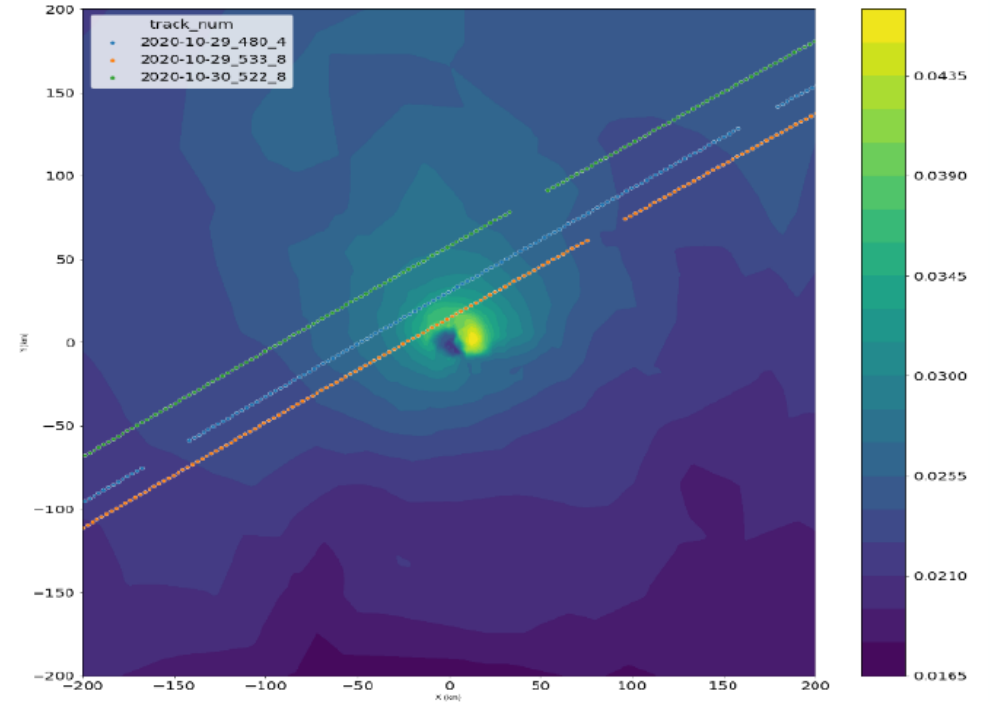
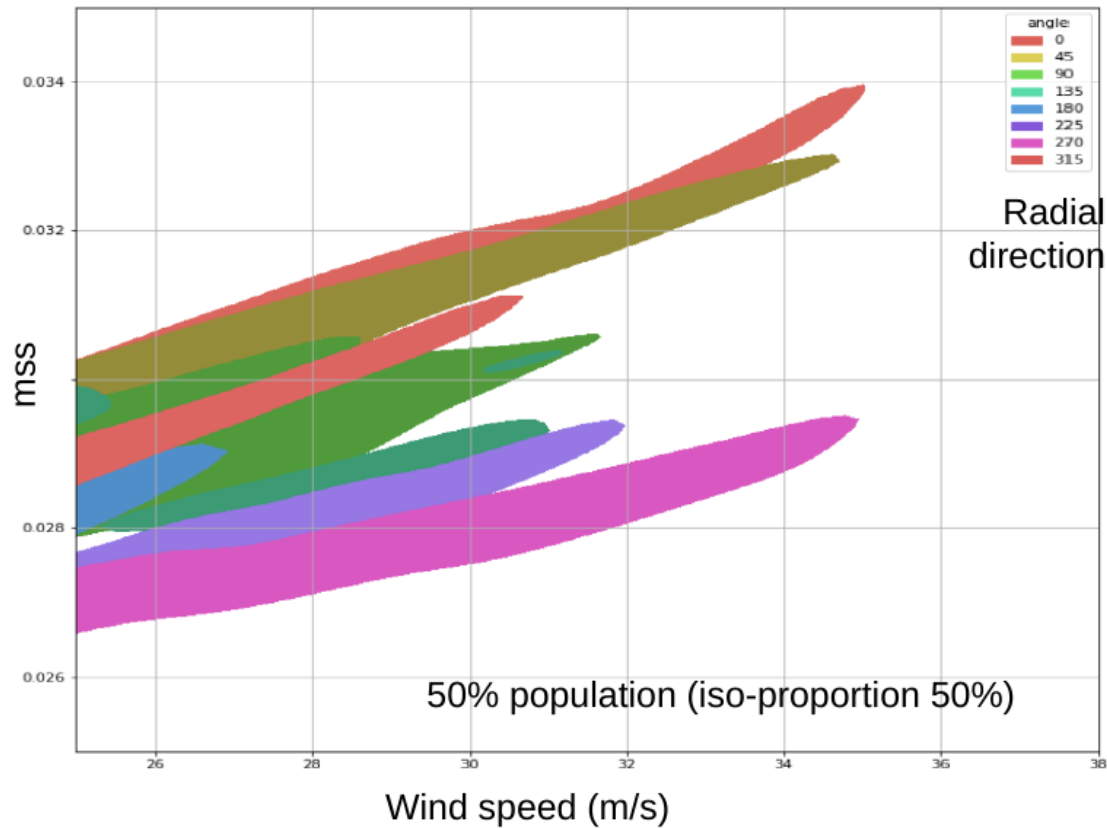
<p>Wind input:</p> <p>$S_w = \beta \omega A(\varphi - \varphi_p) F(\omega)$ <small>(Miles, 1957)</small></p> <p>$\beta = c_p (u_* / c)^2 \cos^2(\varphi - \varphi_w)$ - growth rate <small>(Plant, 1982; Meirink et al. 2003)</small></p>	<p>Dissipation:</p> <p>Wave breaking</p> <p>$D = \omega_p e (k_p^2 e / \varepsilon_T^2)^n$</p> <p>$D = \int S_D d\varphi d\omega$</p> <p>$e = \int E d\varphi d\omega$</p>	<p>Non-linear interactions:</p> <p>Four-wave interactions <small>(Hasselmann, 1962)</small></p> <p>Energy transfer towards low frequencies:</p> <p>$\langle S_N \rangle \sim E^3$</p> <p><small>(Zakharov, 2010; Badulin et al., 2007)</small></p>
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- The model is solved in the storm frame of reference.
- Each wave train location and wave parameters (peak frequency, energy and direction) at each moment of time are obtained with the use of 4th order Runge-Kutta scheme.
- Wave-rays visualize how wave trains develop and travel through the TC varying wind field, and how they leave the storm area as swell systems.
- The wave train with maximal wave length/energy in a given grid cell can be treated as the primary wave system. Each grid cell can be considered to analyze multiple wave systems and their time evolution.
- For this exercise: we used the L-band filtered mss, at each grid cell.



Case study: 2D wave model

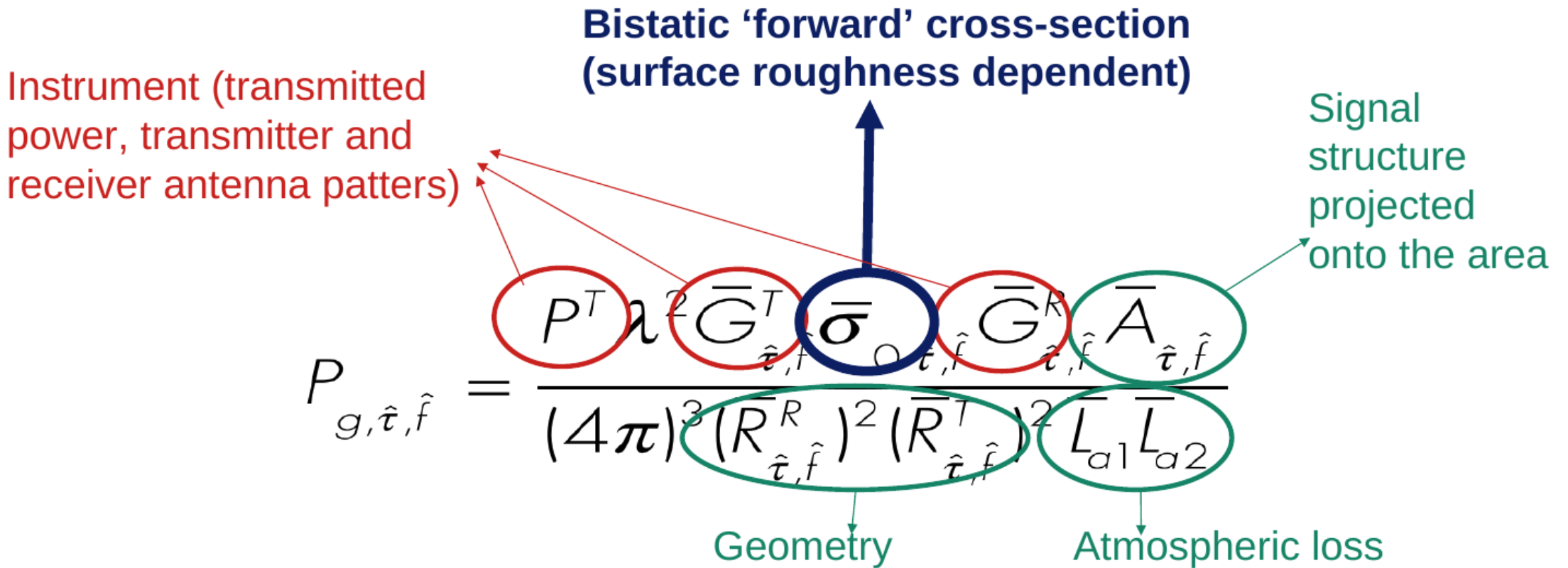
Outcome of the L-band mss model around
TC Goni on 2020-10-29
(three CyGNSS tracks overplotted)



Relationship between mss and wind speed
given by the model, for high winds around Goni

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The received GNSS-R power, at a given 'pixel', can be modeled as:



From measurements:

$$\bar{\sigma}_{0, DDMA} = \frac{P_{g, DDMA} (4\pi)^3 L_{a1} L_{a2} I}{P^T \lambda^2 G_{SP}^T G_{SP}^R R_{SP}^{Total} A_{DDMA}}$$

Calibration required

Forward operator around the specular point (model):

$$\sigma_0 = \pi |\mathcal{R}|^2 PDF(\text{slope} = 0; \{\kappa_s\})$$

From measurements:

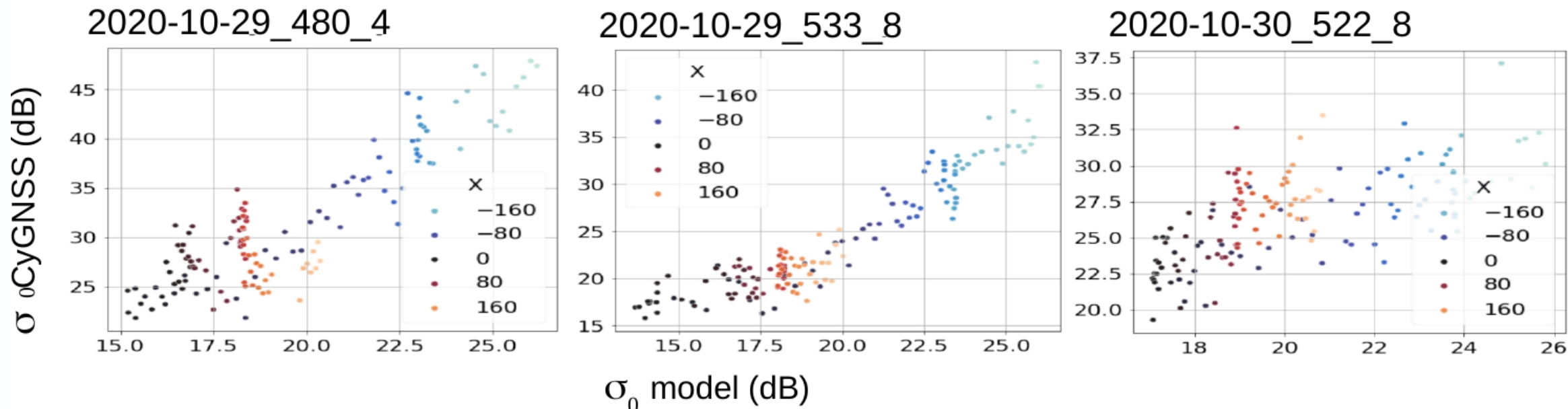
$$\bar{\sigma}_{0, DDMA} = \frac{P_{g, DDMA} (4\pi)^3 L_{a1} L_{a2}}{P^T \lambda^2 G_{SP}^T G_{SP}^R R_{SP}^{Total} A_{DDMA}}$$

Calibration required

Forward operator around the specular point (model):

$$\sigma_0 = \pi |\mathcal{R}|^2 PDF(\text{slope} = 0; \{\kappa_s\})$$

↓
 For simplicity, here we used Gaussian with $\sigma^2 = \text{mss}$
 Should be better done, actual slopes' statistics of the model



Trackwise (for each track): linear fit and calibration of the CyGNSS σ_0

$$\sigma_0^{cal} = \frac{\sigma_0^{CyGNSS} - a}{b}$$

- Preliminary studies to assess the potential of spaceborne GNSS-R ‘observables’ (σ_0) for assimilation into complex wave models across TC.
- Assimilation of σ_0 rather than geophysical variables inverted from GNSS-R (wind, mss) is preferred as
 - GNSS-R is sensitive to a combination of variables;
 - assimilation of σ_0 moves all the assumptions at the model-side (thus consistent with and taking advantage of the richness of the model).
- There is the need of calibration of the data, to ‘align’ it with the model. We suggest trackwise calibration based on linear fit.
- This approach tested on a case study, TC Goni, 2020, 2D wave model [Kudryavtsev, Yurovskaya & Chapron, 2021], initialized with S1 SAR-wind [IFREMER].
- Good agreement between modelled σ_0 and calibrated GNSS-R σ_0 .
- Some ‘anomalies’ detected, where the GNSS-R data could ‘correct’ the model.