

# living planet symposium | BONN 23–27 May 2022

TAKING THE PULSE  
OF OUR PLANET FROM SPACE



## THERMODYNAMIC PROFILING OF THE ATMOSPHERE WITH A SPACEBORNE RAMAN LIDAR FOR ATLAS

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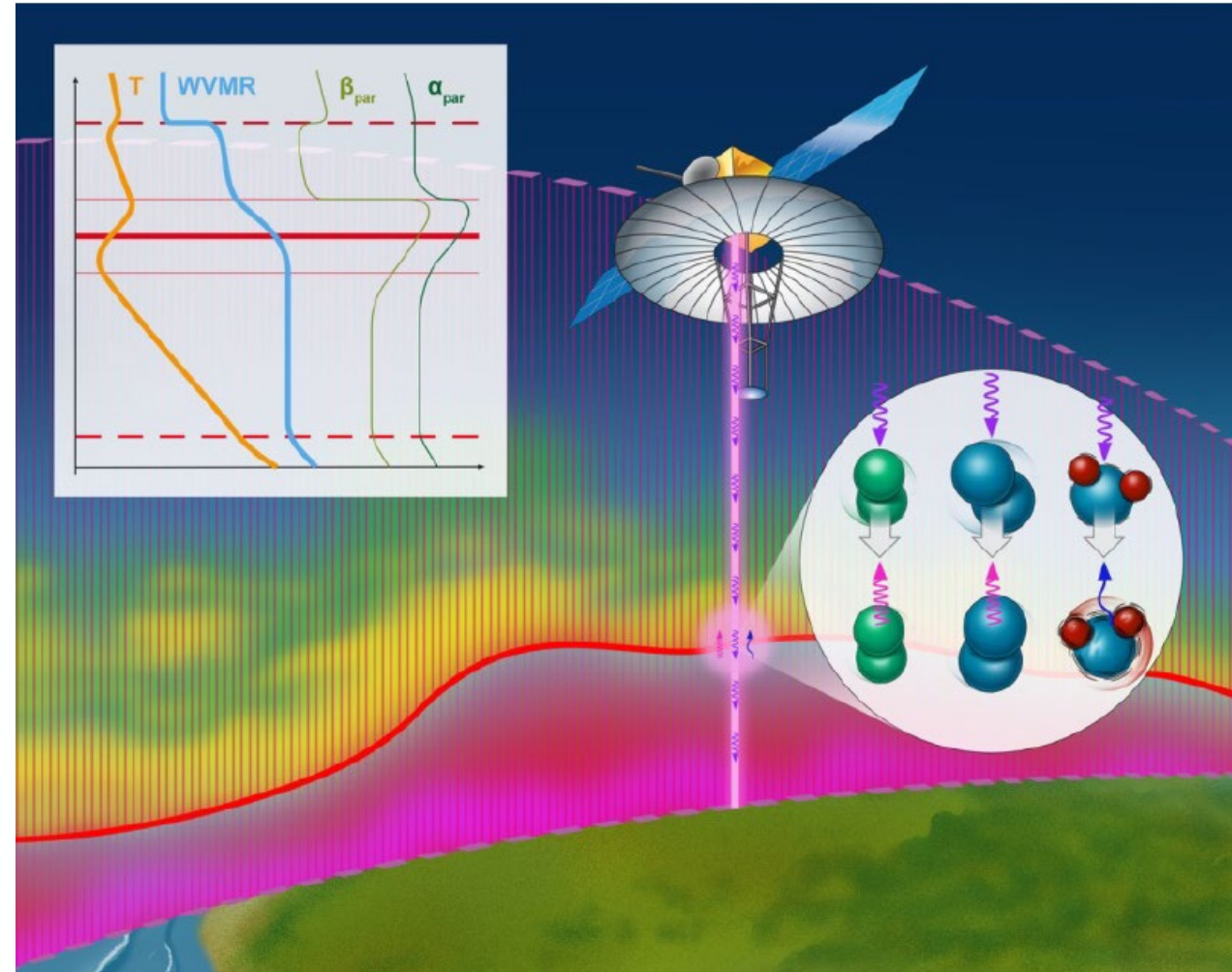
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## ATLAS

### Atmospheric Thermodynamic LidAr in Space

The mission concept aims to develop the first Raman Lidar in space, capable to measure simultaneously atmospheric temperature (T) and water vapour mixing ratio (WVMR) with high temporal and spatial resolutions.

Further data products comprise relative humidity profiles, the particle extinction and backscatter coefficient profiles at 354.7 nm and the PBL depth over land and the oceans.



Accurate, high temporal and spatial resolution observations of thermodynamic profiles in the lower troposphere from the surface to the interfacial layer at the top of the planetary boundary layer (PBL) are critically essential for improving weather forecasting and re-analyses, and for understanding the Earth system.

Global scale measurements of 3D thermodynamic profiles would have a great impact in several research areas [Wulfmeyer et al., 2015]:

- Radiative transfer, as well as regional and global water and energy budgets,
- Land-atmosphere feedback including the surface energy balance and its dependence on soil properties and land cover,
- Mesoscale circulations and convection initiation
- Data assimilation



The space-borne Raman Lidar considered in ATLAS collects six lidar signals:

- The Elastic backscattered signal at the laser wavelength
- The WV roto-vibrational Raman signal
- The high- and low-quantum number O<sub>2</sub>-N<sub>2</sub> rotational Raman signals both in the anti-Stokes and Stokes branches

$$WVMR(z) = k \Delta Trs(z) \frac{P_{WV}(z)}{P_{ref}(z)}$$

$$T(z) = \frac{a}{\ln \left( \frac{P_{HiJ,S}(z) + P_{HiJ,AS}(z)}{P_{LoJ,S}(z) + P_{LoJ,AS}(z)} \right) - b}$$

$$\beta_{\lambda_0}^{par}(z) = \beta_{\lambda_0}^{mol}(z) \left[ \frac{P_{\lambda_0}(z)}{c P_{ref}(z)} - 1 \right]$$

$$\alpha_{\lambda_0}^{par}(z) = \frac{1}{2} \frac{d}{dz} \ln \left( \frac{n(z)}{P_{ref}(z) z^2} \right) - \alpha_{\lambda_0}^{mol}(z)$$

## LIDAR TRANSMITTER

Source	Injection-seeded frequency tripled, diode-laser pumped Nd:YAG
Laser Wavelength	354.7 nm
Single-shot pulse energy	1 J
Repetition rate	200 Hz

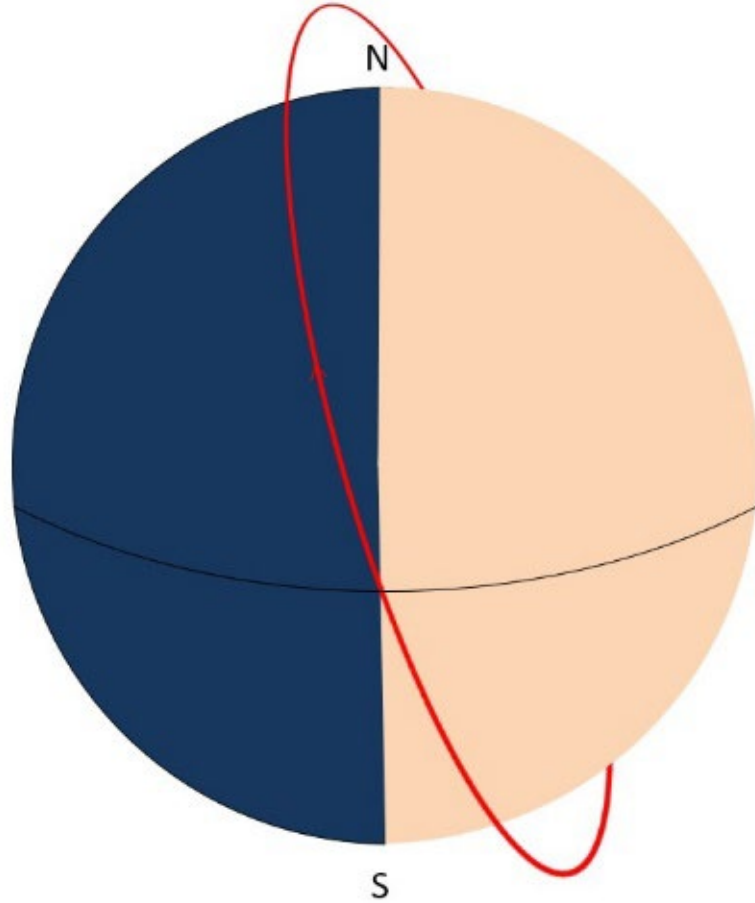
## LIDAR RECEIVER

Telescope	f/5 a-focal Cassegrain
Telescope diameter	2 m
Field-of-View (FWHM)	25 $\mu$ rad

## SPECTRAL SELECTION AND DETECTION

Spectral selection devices	Interference Filters (IFs)
Detection Devices	Photodiodes or Photomultipliers
Quantum efficiencies	85 %



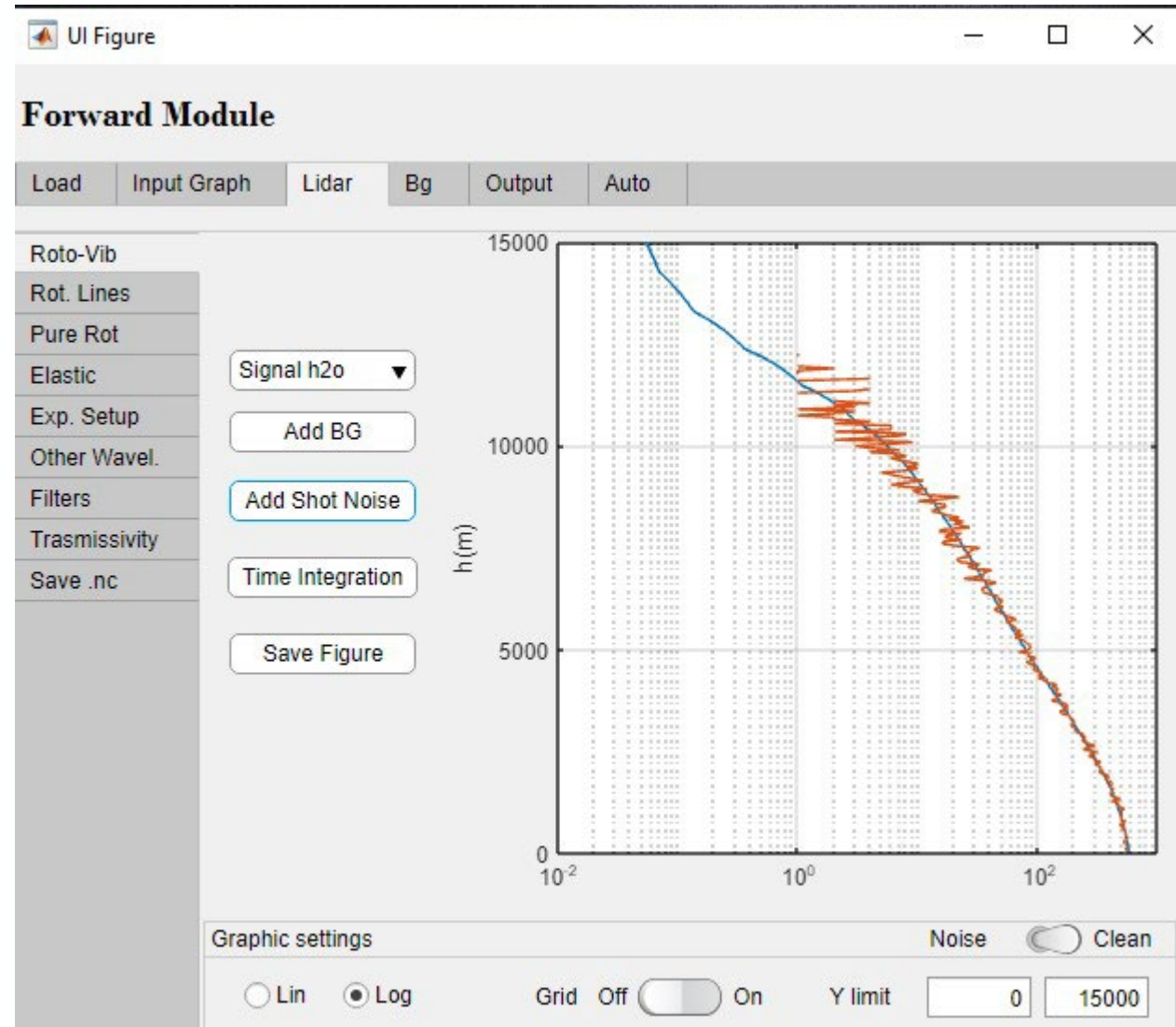


The selected orbit for the satellite is a sun synchronous polar orbit with a local time descending node at 06:00 (dawn-dusk orbit) and an orbital height of 450 km.

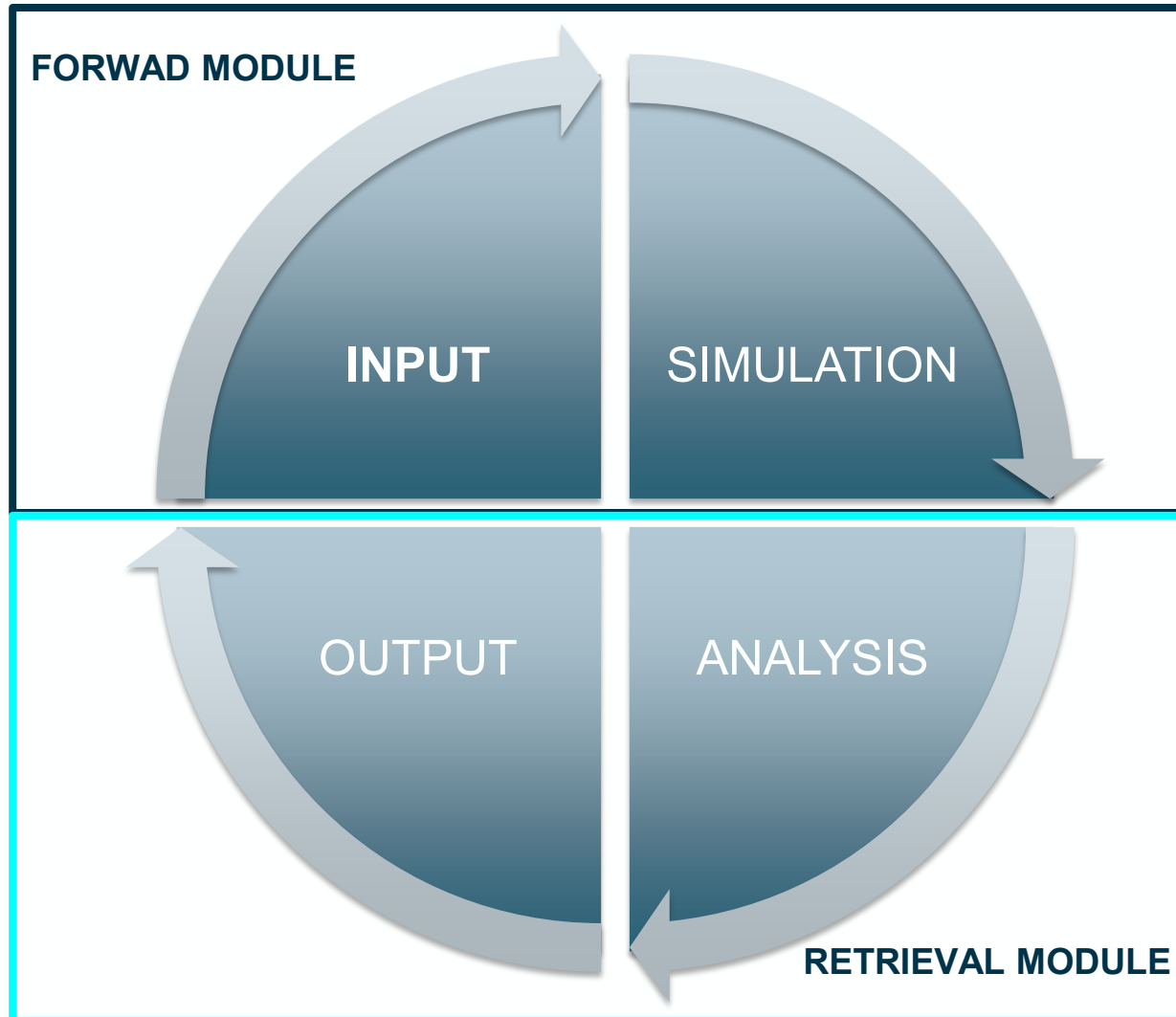
This kind of orbit guarantees a sun zenith angle on the sub-satellite point close to 90°

# Specification assessment

An assessment of the specifications of the different lidar sub-systems has been performed with an analytical simulation model for space-borne Raman lidar systems [Di Girolamo et al., 2018] and verified through an end-to-end numerical simulation model.

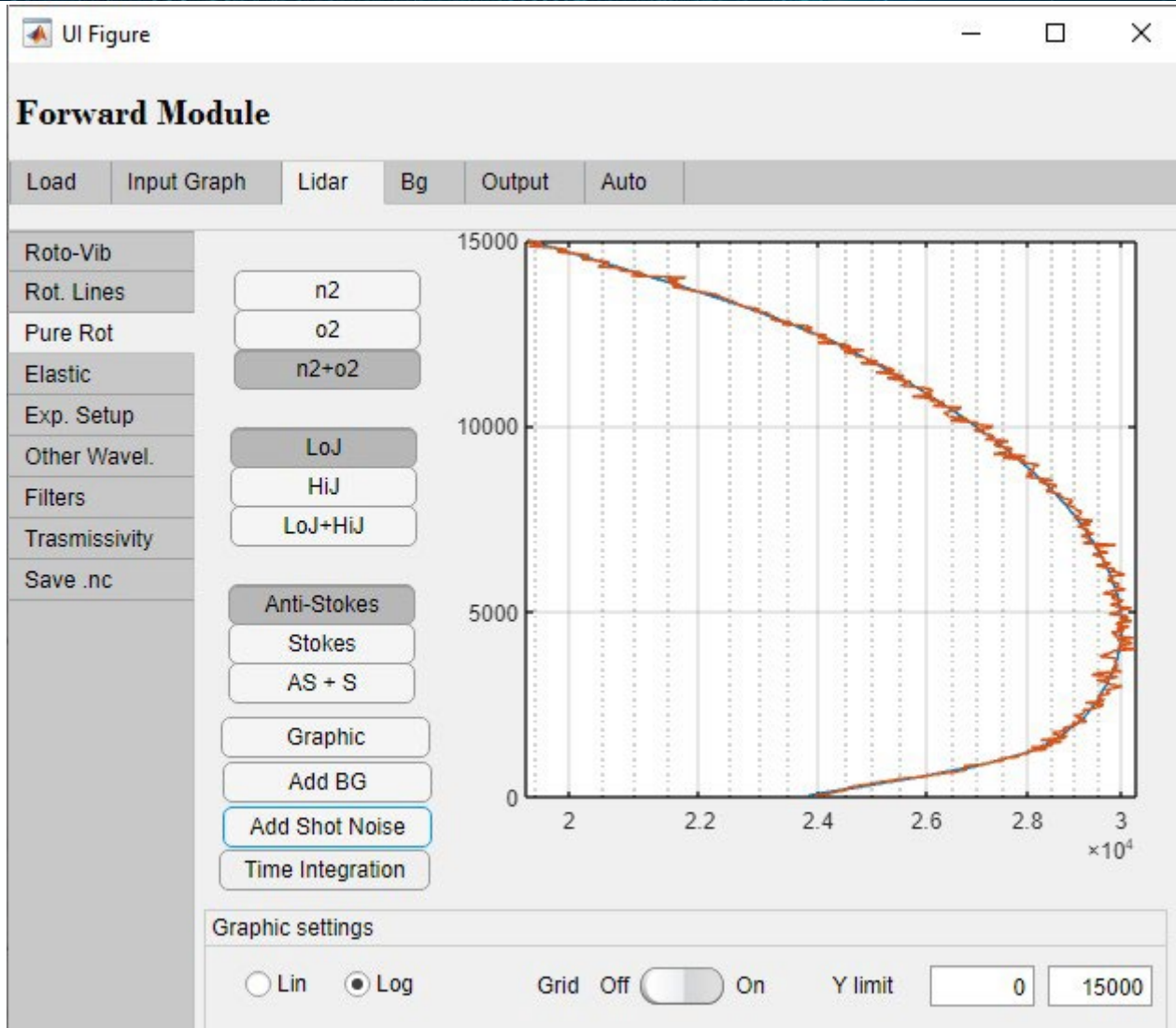






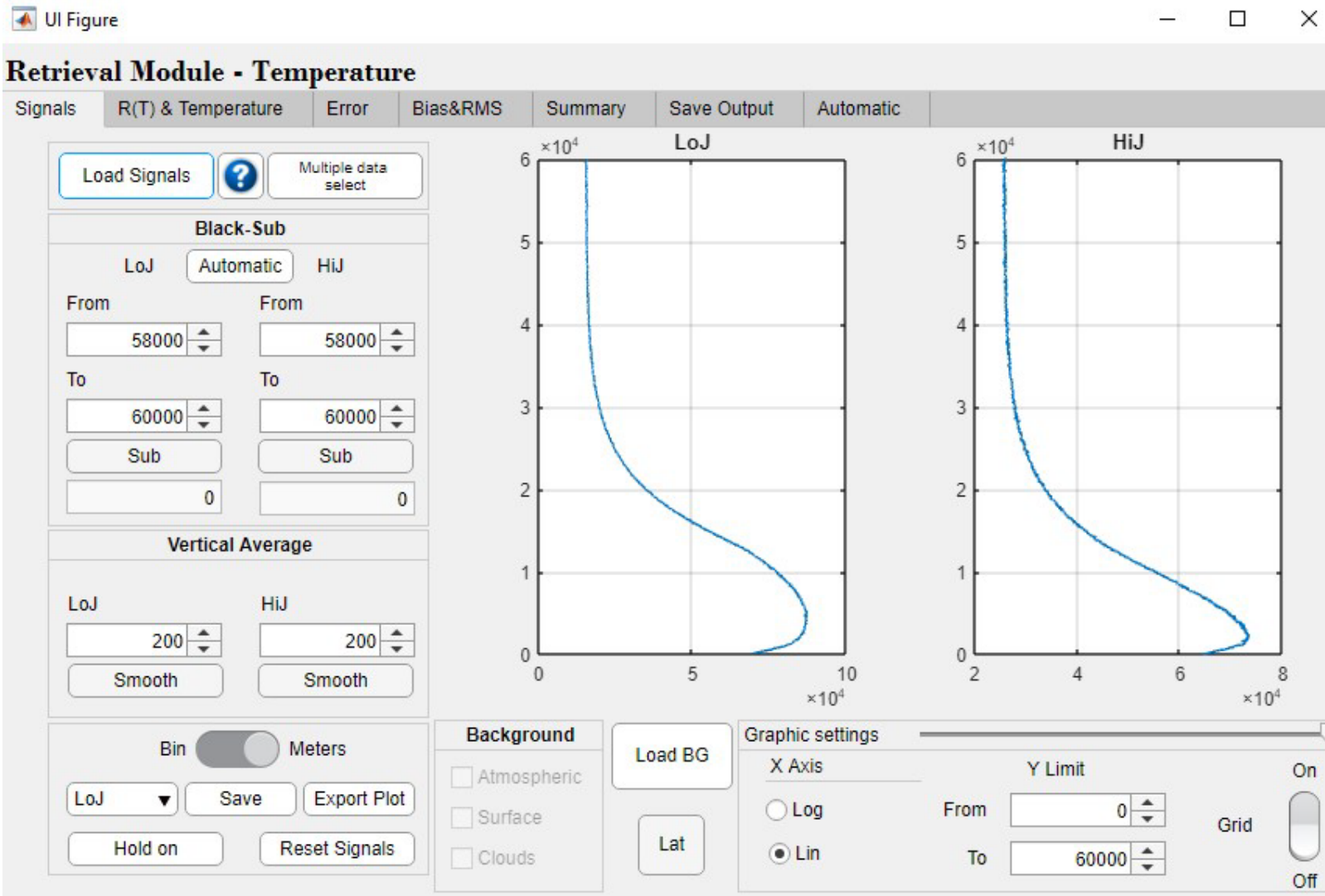
## INPUT DATA:

- Thermodynamic profiles (temperature, pressure, air density)
- Optical properties (backscatter and extinction coefficient)
- Surface albedo, clouds ...



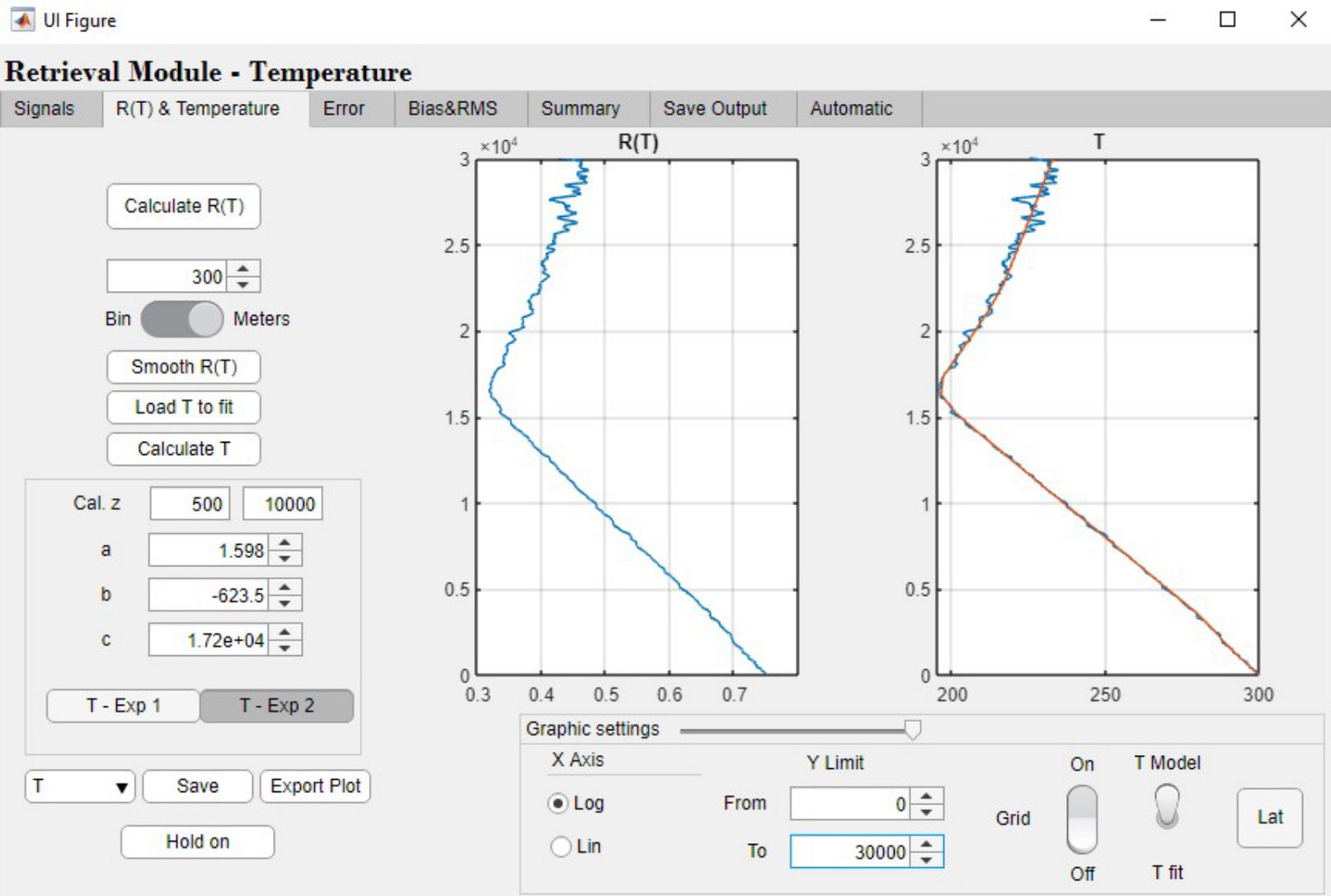
## SIMULATION:

- Behavior of all the devices in the experimental system
- Simulation of the signals
- Background contribution
- Shot-noise contribution



## SIGNAL ANALYSIS

- Background subtraction
- Vertical average
- Application of Raman Lidar techniques



## OUTPUT:

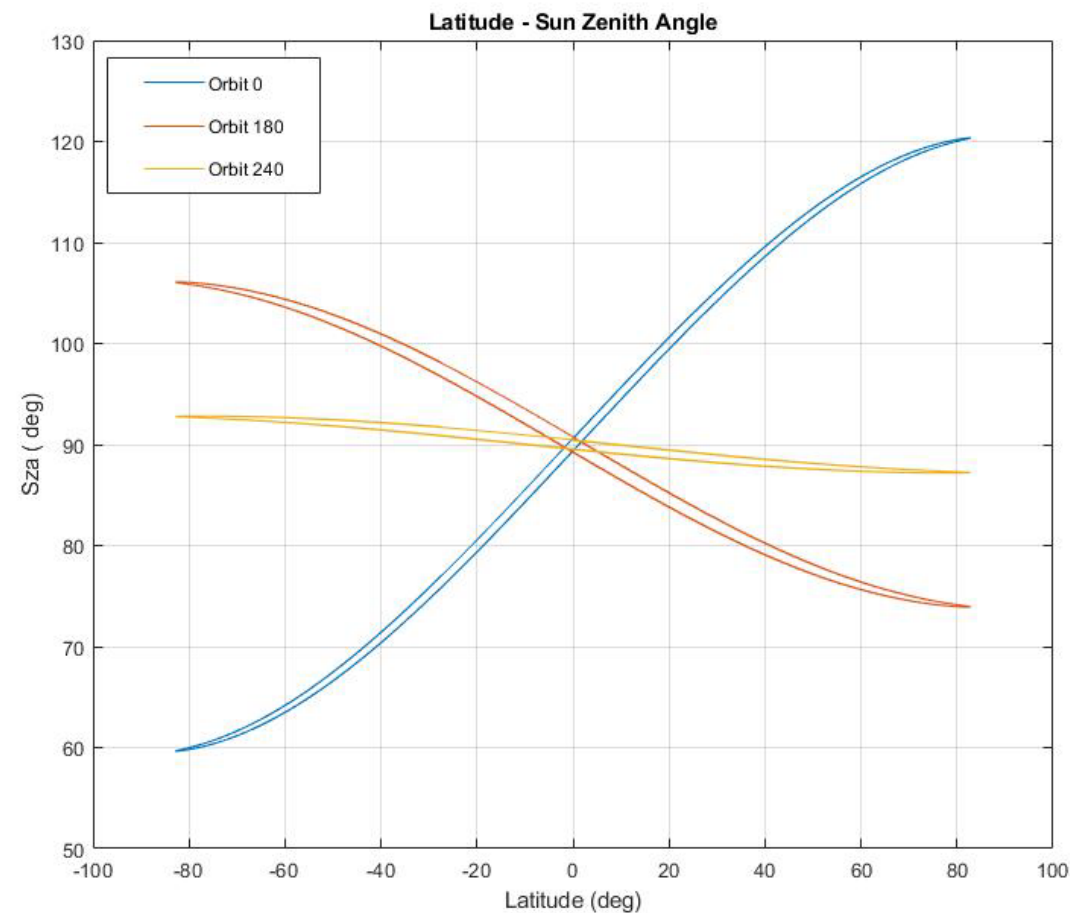
- Retrieval of water vapour mixing ratio and atmospheric temperature profiles
- Statistical uncertainties for temperature (K) and WVMR (g/kg; %)
- Bias for temperature (K) and WVMR (g/kg; %)

To verify the capabilities of ATLAS, simulations along several orbits around the Earth were performed.

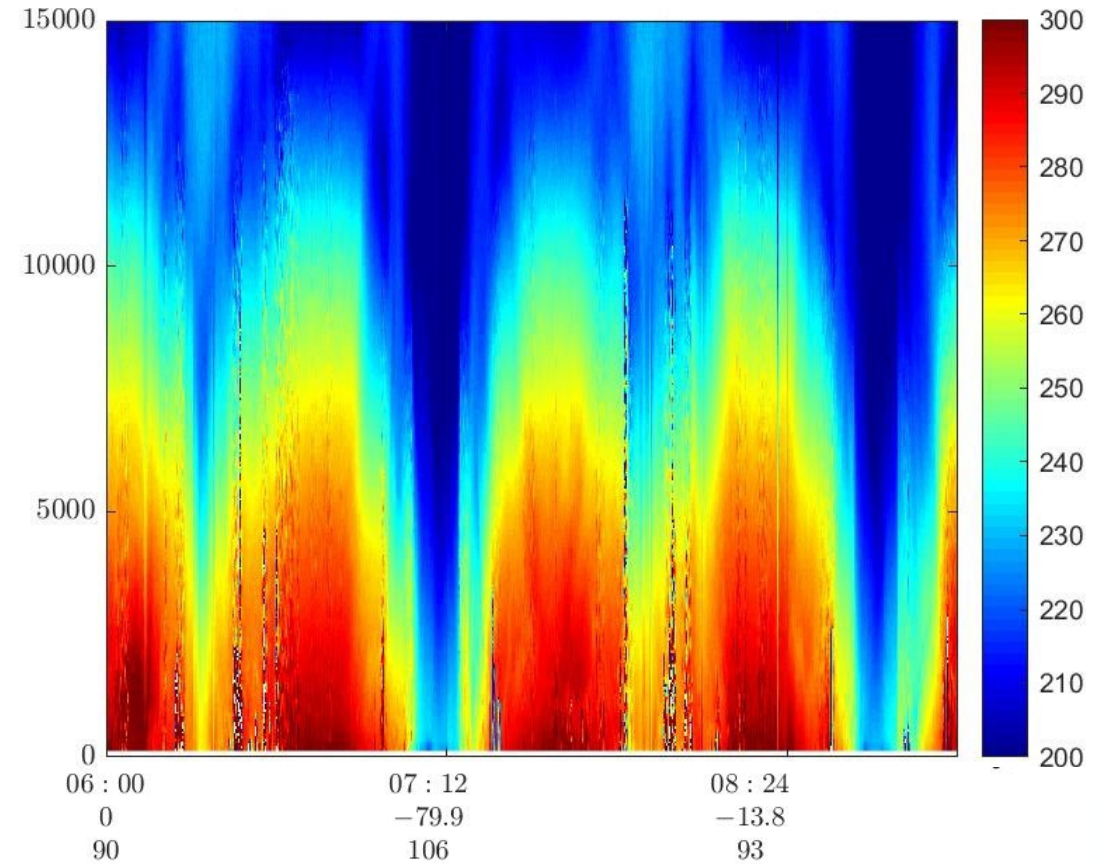
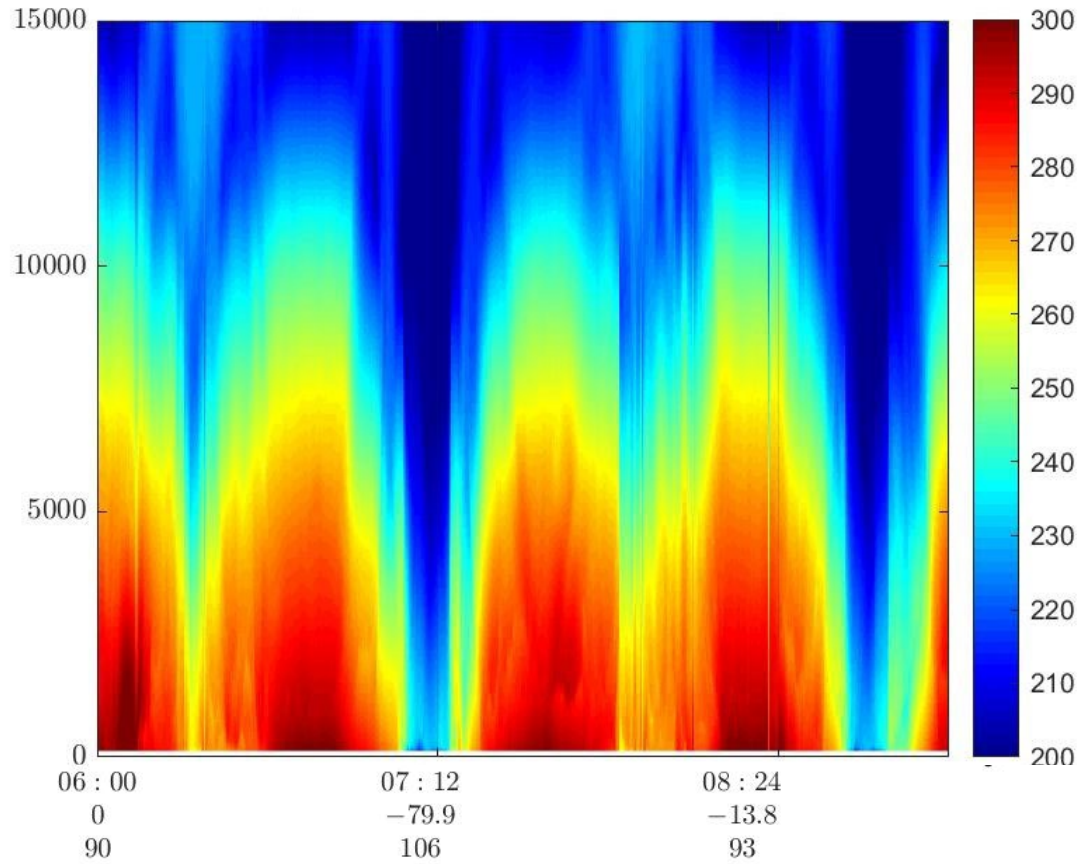
The orbit were selected to consider the different variability of the sza during the year:

- Day 0:  $90^{\circ} \pm 30^{\circ}$  (High background)
- Day 180:  $90^{\circ} \pm 15^{\circ}$  (Mean background)
- Day 240:  $90^{\circ} \pm 3^{\circ}$  (Low background)

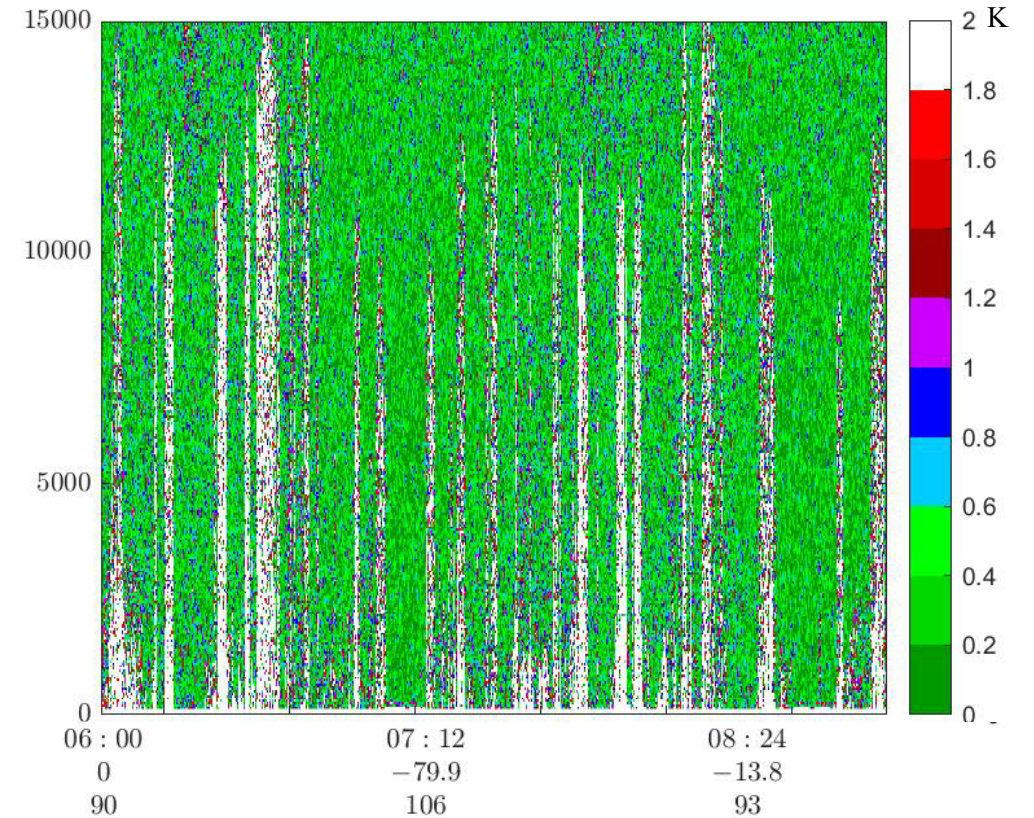
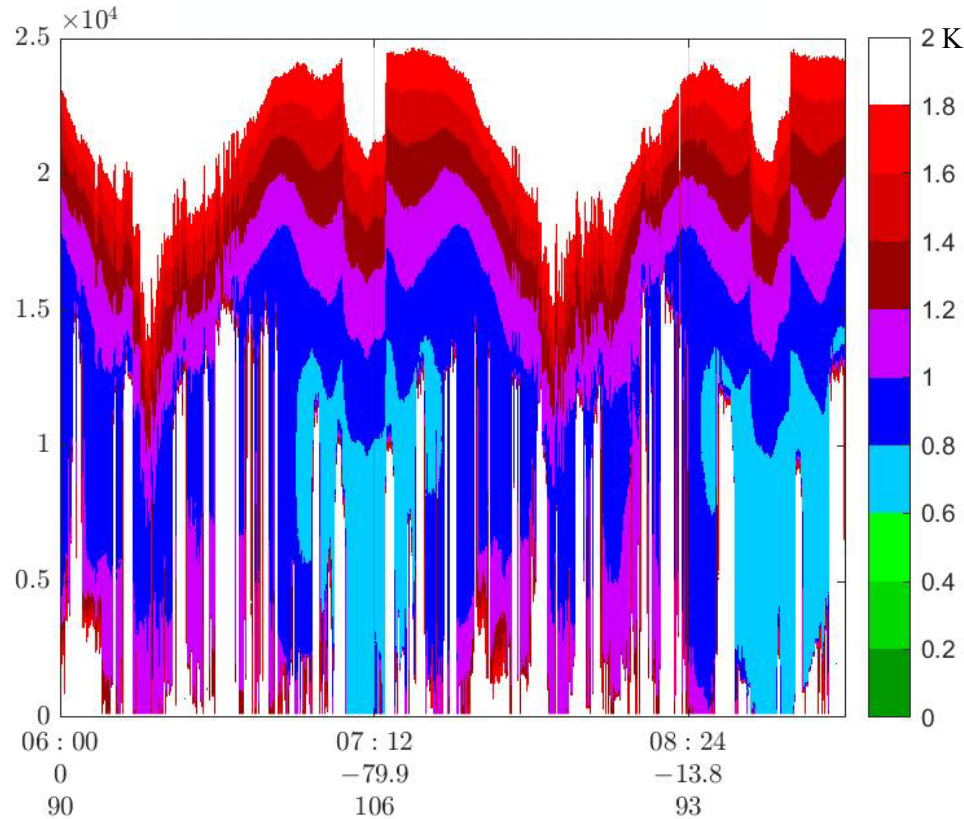
The data are extracted from NASA's Goddard Earth Observing System Model (GEOS-5) analysis



# Input and Retrieved Temperature (K)



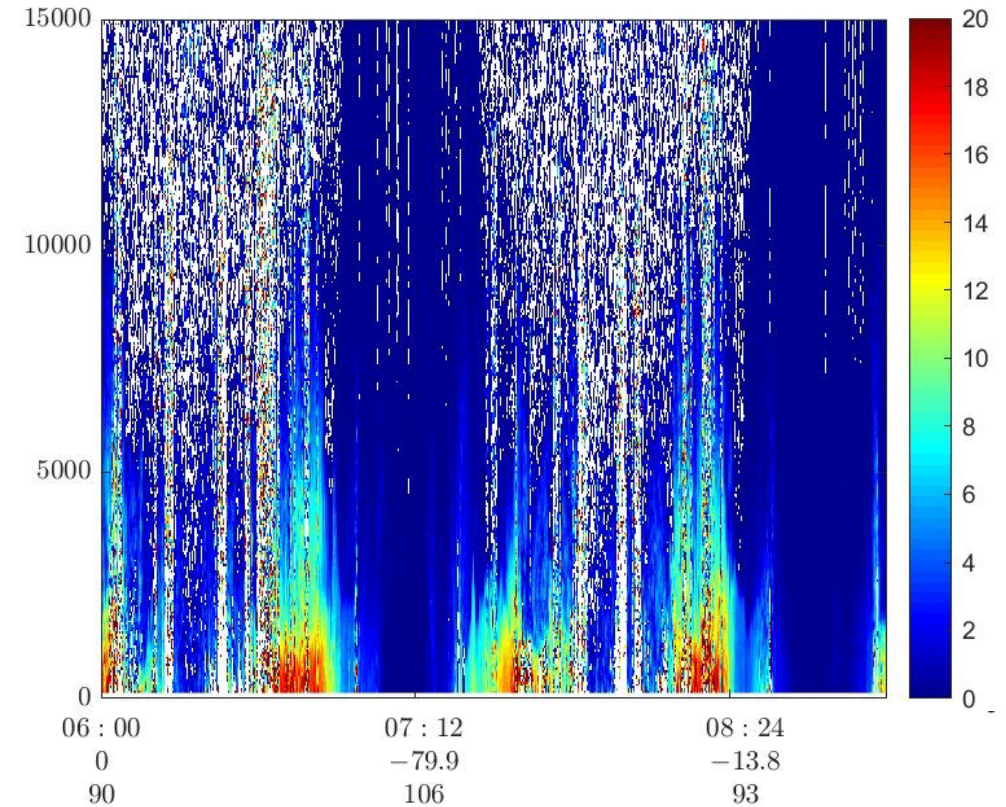
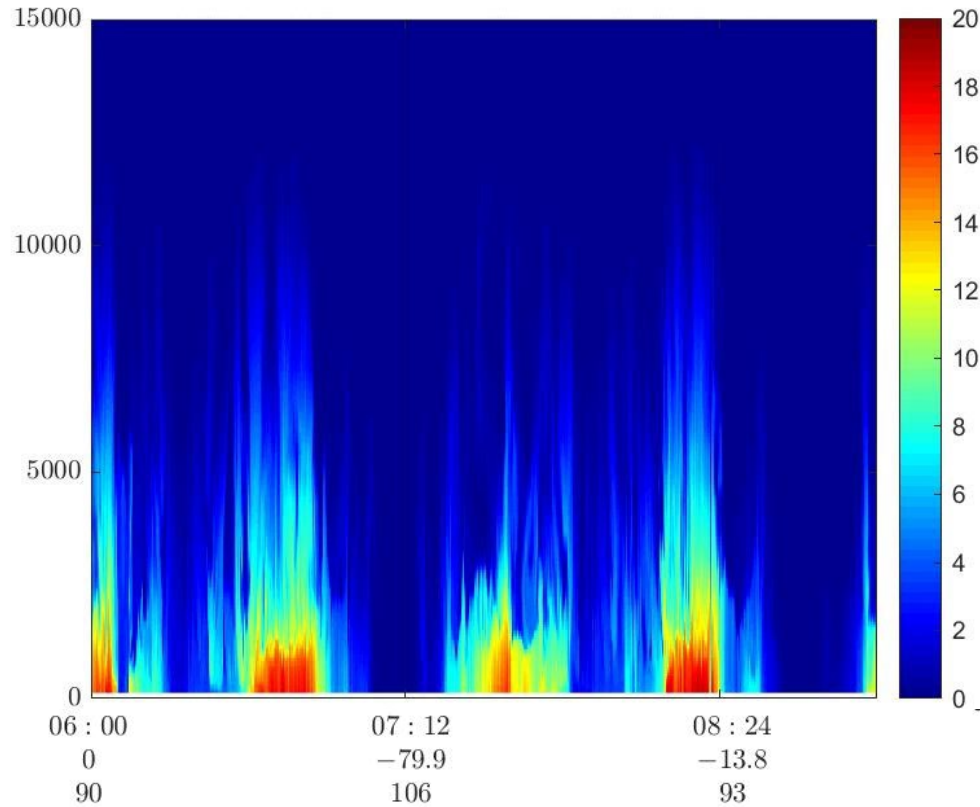
# Statistical uncertainty (K) and bias (K)



$$\Delta T(z) = \frac{\partial T(z)}{\partial R} R(z) \sqrt{\frac{N_{LoJ}(z) + bg_{LoJ}}{N_{LoJ}^2(z)} + \frac{N_{HiJ}(z) + bg_{HiJ}}{N_{HiJ}^2(z)}}$$

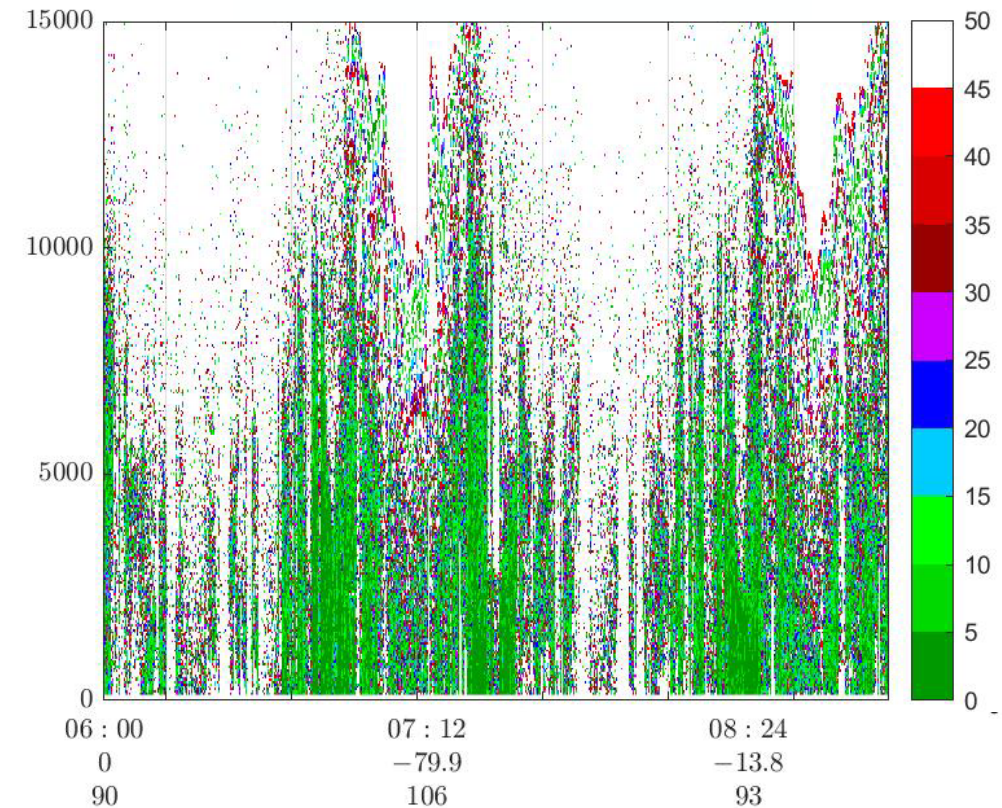
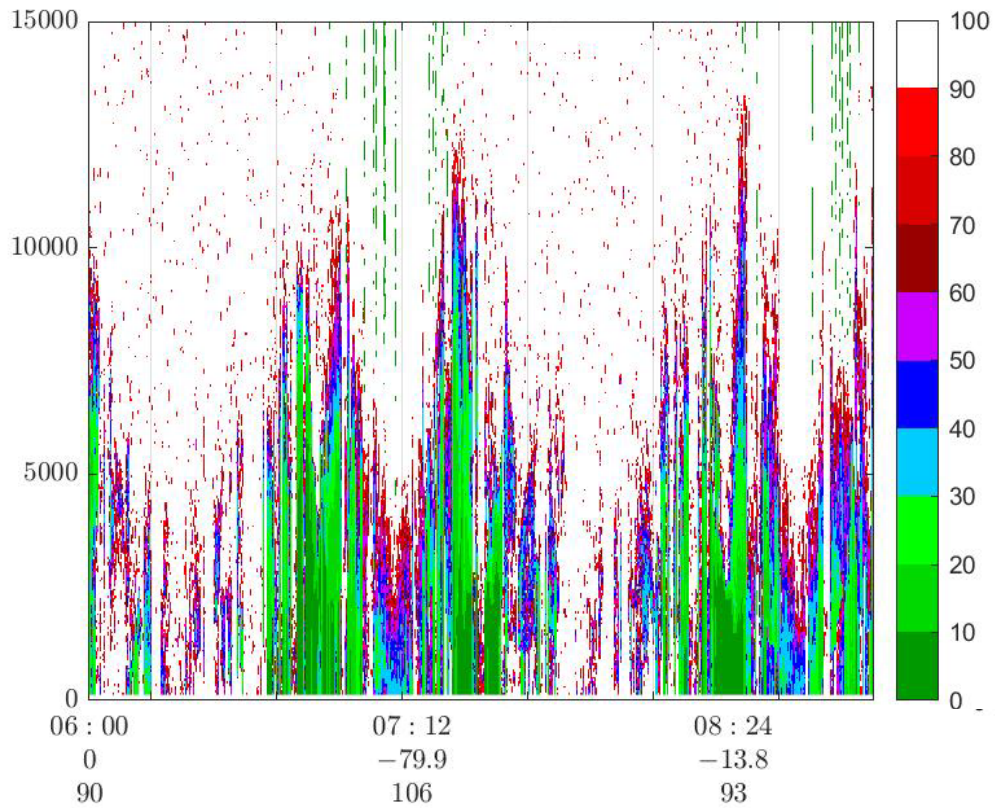
$$\Delta T = |T_{inp} - T_{ret}|$$

# Input and Retrieved Water Vapour Mixing Ratio (g/kg)





# Relative Statistical Uncertainty (%) and relative Bias (%)



$$\frac{\Delta\chi_{H_2O}(z)}{\chi_{H_2O}(z)} = 100 \times \sqrt{\frac{N_{H_2O}(z) + bg_{H_2O}}{N_{H_2O}^2(z)} + \frac{N_{ref}(z) + bg_{ref}}{N_{ref}^2(z)}}$$

$$\frac{\Delta\chi_{H_2O}}{\chi_{H_2O}} = 100 \times \frac{|\chi_{inp} - \chi_{ret}|}{\chi_{inp}}$$

