

# Upward continuation of satellite altimeter data for GOCE validation

**Sebera J., Bosch W., Bouman J., Bezděk A., Klokočník J.,  
Kostelecký J.**

Czech Technical University in Prague  
Deutsches Geodätisches Forschungsinstitut, Germany  
Astronomical Institute of the Academy of Sciences of the Czech Republic  
Research Institute of Geodesy, Topography and Cartography, Czech Republic

**ESA Living Planet Symposium, Bergen, Norway**  
[www.congrex.nl](http://www.congrex.nl)

28 June - 2 July, 2010

## 1 Introduction

- Satellite altimetry

## 2 Upward experiment

- Global approximation
- Input signal

## 3 Numerical comparisons

- Upwarded 0th derivative:  $T$
- Upwarded 1st derivative:  $T_r, T_{\bar{z}}$
- Upwarded 2nd derivative:  $T_{rr}, T_{\bar{z}\bar{z}}$

## 4 Concluding remarks

## 1 Introduction

- Satellite altimetry

## 2 Upward experiment

- Global approximation
- Input signal

## 3 Numerical comparisons

- Upwarded 0th derivative:  $T$
- Upwarded 1st derivative:  $T_r, T_{\bar{z}}$
- Upwarded 2nd derivative:  $T_{rr}, T_{\bar{z}\bar{z}}$

## 4 Concluding remarks

## 1 Introduction

- Satellite altimetry

## 2 Upward experiment

- Global approximation
- Input signal

## 3 Numerical comparisons

- Upwarded 0th derivative:  $T$
- Upwarded 1st derivative:  $T_r, T_{\bar{z}}$
- Upwarded 2nd derivative:  $T_{rr}, T_{\bar{z}\bar{z}}$

## 4 Concluding remarks

## 1 Introduction

- Satellite altimetry

## 2 Upward experiment

- Global approximation
- Input signal

## 3 Numerical comparisons

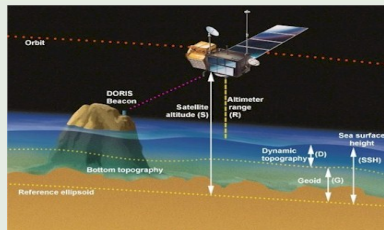
- Upwarded 0th derivative:  $T$
- Upwarded 1st derivative:  $T_r, T_{\bar{z}}$
- Upwarded 2nd derivative:  $T_{rr}, T_{\bar{z}\bar{z}}$

## 4 Concluding remarks

## Roughly on satellite altimetry

- SA gives the geoid undulation since it holds  $MSS = DOT + N$  (MSS-mean sea surface, DOT-dynamic ocean topography).
- Having DOT at hand we can determine disturbing/anomalous potential from the Bruns formula  $T = N\gamma$ .

### SA - principle (from AVISO-CNES)



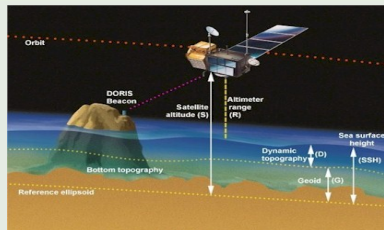
### Some assumptions

- DOT is also unknown  $\Rightarrow$  not solved here
- SA coverage isn't global  $\Rightarrow$  not considered
- $\Rightarrow T$  globally is our input for the base functions experiments

## Roughly on satellite altimetry

- SA gives the geoid undulation since it holds  $MSS = DOT + N$  (MSS-mean sea surface, DOT-dynamic ocean topography).
- Having DOT at hand we can determine disturbing/anomalous potential from the Bruns formula  $T = N\gamma$ .

### SA - principle (from AVISO-CNES)



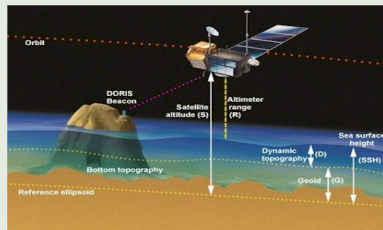
### Some assumptions

- DOT is also unknown  $\Rightarrow$  not solved here
- SA coverage isn't global  $\Rightarrow$  not considered
- $\Rightarrow T$  globally is our input for the base functions experiments

## Roughly on satellite altimetry

- SA gives the geoid undulation since it holds  $MSS = DOT + N$  (MSS-mean sea surface, DOT-dynamic ocean topography).
- Having DOT at hand we can determine disturbing/anomalous potential from the Bruns formula  $T = N\gamma$ .

### SA - principle (from AVISO-CNES)



### Some assumptions

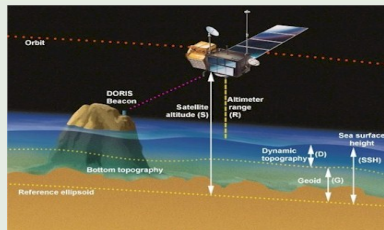
- DOT is also unknown  
 $\Rightarrow$  **not solved here**
- SA coverage isn't global  $\Rightarrow$  **not considered**
- $\Rightarrow T$  globally is our input for the base functions experiments



## Roughly on satellite altimetry

- SA gives the geoid undulation since it holds  $MSS = DOT + N$  (MSS-mean sea surface, DOT-dynamic ocean topography).
- Having DOT at hand we can determine disturbing/anomalous potential from the Bruns formula  $T = N\gamma$ .

### SA - principle (from AVISO-CNES)



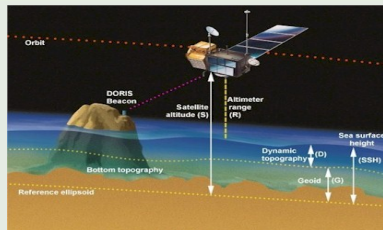
### Some assumptions

- DOT is also unknown  $\Rightarrow$  **not solved here**
- SA coverage isn't global  $\Rightarrow$  **not considered**
- $\Rightarrow T$  globally is our input for the base functions experiments

## Roughly on satellite altimetry

- SA gives the geoid undulation since it holds  $MSS = DOT + N$  (MSS-mean sea surface, DOT-dynamic ocean topography).
- Having DOT at hand we can determine disturbing/anomalous potential from the Bruns formula  $T = N\gamma$ .

### SA - principle (from AVISO-CNES)



### Some assumptions

- DOT is also unknown  $\Rightarrow$  **not solved here**
- SA coverage isn't global  $\Rightarrow$  **not considered**
- $\Rightarrow T$  globally is our input for the base functions experiments

## Numerical test of the approximation

- ① Start with a signal on the geoid -  $T$
- ② Use both kinds of approximation  $\Rightarrow \{C_{nm}^e, S_{nm}^e\}$  and  $\{C_{nm}^s, S_{nm}^s\}$
- ③ Map both sets by upward operators onto the potential functionals at satellite altitude and compare EHS and SHS

### Grid settings

- Regular grid on **sphere**  
(geocentric co-latitude  $\theta$ ) is not regular on the **ellipsoid**  
(reduced co-latitude  $\vartheta$ )
- Trade-off  $\Rightarrow$  mixture of both
- For SHA  $P \in \{r, \theta + \theta(\vartheta), \lambda\}$
- For EHA  $P \in \{u, \vartheta(\theta) + \vartheta, \lambda\}$

### Harmonic analysis

- $\Rightarrow$  "Semi-regular" grid  
 $f = (2N_{max} - 1, 2N_{max})$
- $\Rightarrow$  WLS solution for blocks used
- $\Rightarrow$  Latitudinal weights  
$$W_i(\theta) = 2 \frac{\sin \theta_i}{\sum_{i=1}^{2N-1} \sin \theta_i}$$

## Numerical test of the approximation

- ① Start with a signal on the geoid -  $T$
- ② Use both kinds of approximation  $\Rightarrow \{C_{nm}^e, S_{nm}^e\}$  and  $\{C_{nm}^s, S_{nm}^s\}$
- ③ Map both sets by upward operators onto the potential functionals at satellite altitude and compare EHS and SHS

### Grid settings

- Regular grid on **sphere**  
(geocentric co-latitude  $\theta$ ) is not regular on the **ellipsoid**  
(reduced co-latitude  $\vartheta$ )
- Trade-off  $\Rightarrow$  mixture of both
- For SHA  $P \in \{r, \theta + \theta(\vartheta), \lambda\}$
- For EHA  $P \in \{u, \vartheta(\theta) + \vartheta, \lambda\}$

### Harmonic analysis

- $\Rightarrow$  "Semi-regular" grid  
 $f = (2N_{max} - 1, 2N_{max})$
- $\Rightarrow$  WLS solution for blocks used
- $\Rightarrow$  Latitudinal weights  
$$W_i(\theta) = 2 \frac{\sin \theta_i}{\sum_{i=1}^{2N-1} \sin \theta_i}$$

## Numerical test of the approximation

- ① Start with a signal on the geoid -  $T$
- ② Use both kinds of approximation  $\Rightarrow \{C_{nm}^e, S_{nm}^e\}$  and  $\{C_{nm}^s, S_{nm}^s\}$
- ③ Map both sets by upward operators onto the potential functionals at satellite altitude and compare EHS and SHS

### Grid settings

- Regular grid on **sphere**  
(geocentric co-latitude  $\theta$ ) is not regular on the **ellipsoid**  
(reduced co-latitude  $\vartheta$ )
- Trade-off  $\Rightarrow$  mixture of both
- For SHA  $P \in \{r, \theta + \theta(\vartheta), \lambda\}$
- For EHA  $P \in \{u, \vartheta(\theta) + \vartheta, \lambda\}$

### Harmonic analysis

- $\Rightarrow$  "Semi-regular" grid  
 $f = (2N_{max} - 1, 2N_{max})$
- $\Rightarrow$  WLS solution for blocks used
- $\Rightarrow$  Latitudinal weights  
$$W_i(\theta) = 2 \frac{\sin \theta_i}{\sum_{i=1}^{2N-1} \sin \theta_i}$$

## Numerical test of the approximation

- ① Start with a signal on the geoid -  $T$
- ② Use both kinds of approximation  $\Rightarrow \{C_{nm}^e, S_{nm}^e\}$  and  $\{C_{nm}^s, S_{nm}^s\}$
- ③ Map both sets by upward operators onto the potential functionals at satellite altitude and compare EHS and SHS

### Grid settings

- Regular grid on **sphere**  
(geocentric co-latitude  $\theta$ ) is not regular on the **ellipsoid**  
(reduced co-latitude  $\vartheta$ )
- Trade-off  $\Rightarrow$  mixture of both
- For SHA  $P \in \{r, \theta + \theta(\vartheta), \lambda\}$
- For EHA  $P \in \{u, \vartheta(\theta) + \vartheta, \lambda\}$

### Harmonic analysis

- $\Rightarrow$  "Semi-regular" grid  
 $f = (2N_{max} - 1, 2N_{max})$
- $\Rightarrow$  WLS solution for blocks used
- $\Rightarrow$  Latitudinal weights  
$$W_i(\theta) = 2 \frac{\sin \theta_i}{\sum_{i=1}^{2N-1} \sin \theta_i}$$

## Numerical test of the approximation

- ① Start with a signal on the geoid -  $T$
- ② Use both kinds of approximation  $\Rightarrow \{C_{nm}^e, S_{nm}^e\}$  and  $\{C_{nm}^s, S_{nm}^s\}$
- ③ Map both sets by upward operators onto the potential functionals at satellite altitude and compare EHS and SHS

### Grid settings

- Regular grid on **sphere**  
(geocentric co-latitude  $\theta$ ) is not regular on the **ellipsoid**  
(reduced co-latitude  $\vartheta$ )
- Trade-off  $\Rightarrow$  mixture of both
- For SHA  $P \in \{r, \theta + \theta(\vartheta), \lambda\}$
- For EHA  $P \in \{u, \vartheta(\theta) + \vartheta, \lambda\}$

### Harmonic analysis

- $\Rightarrow$  "Semi-regular" grid  
 $f = (2N_{max} - 1, 2N_{max})$
- $\Rightarrow$  WLS solution for blocks used
- $\Rightarrow$  Latitudinal weights  
$$W_i(\theta) = 2 \frac{\sin \theta_i}{\sum_{i=1}^{2N-1} \sin \theta_i}$$

## Spherical harmonics

very well represent a functional  $f = f(r, \theta, \lambda)$  on Earth  $\sim$  spherical approximation

$$T^s = \frac{GM}{a} \sum_{n,m} \left(\frac{a}{r}\right)^{n+1} (C_{nm}^s \cos m\lambda + S_{nm}^s \sin m\lambda) P_{nm}(\cos \theta) \quad (1)$$

## Ellipsoidal harmonics

are much closer to Earth's geometry, functional  $f = f(u, \vartheta, \lambda)$

$$T^e = \frac{GM}{a} \sum_{n,m} \frac{Q_{nm}\left(\frac{u}{E}\right)}{Q_{nm}\left(\frac{b}{E}\right)} (C_{nm}^e \cos m\lambda + S_{nm}^e \sin m\lambda) P_{nm}(\cos \vartheta) \quad (2)$$

or with Jekeli's renormalization  $S_{nm}\left(\frac{u}{E}\right)/S_{nm}\left(\frac{b}{E}\right) = Q_{nm}\left(\frac{u}{E}\right)/Q_{nm}\left(\frac{b}{E}\right)$

## "Normal" derivatives ( $z$ axis in LNOF)

$$\begin{aligned} T_r^s &= \frac{\partial V^s}{\partial r} \approx \frac{\partial V^e}{\partial \bar{z}} = T_{\bar{z}}^e \\ T_{rr}^s &= \frac{\partial^2 V^s}{\partial r^2} \approx \frac{\partial^2 V^e}{\partial \bar{z}^2} = T_{\bar{z}\bar{z}}^e \end{aligned}$$



## Spherical harmonics

very well represent a functional  $f = f(r, \theta, \lambda)$  on Earth  $\sim$  spherical approximation

$$T^s = \frac{GM}{a} \sum_{n,m} \left(\frac{a}{r}\right)^{n+1} (C_{nm}^s \cos m\lambda + S_{nm}^s \sin m\lambda) P_{nm}(\cos \theta) \quad (1)$$

## Ellipsoidal harmonics

are much closer to Earth's geometry, functional  $f = f(u, \vartheta, \lambda)$

$$T^e = \frac{GM}{a} \sum_{n,m} \frac{Q_{nm}\left(\frac{u}{E}\right)}{Q_{nm}\left(\frac{b}{E}\right)} (C_{nm}^e \cos m\lambda + S_{nm}^e \sin m\lambda) P_{nm}(\cos \vartheta) \quad (2)$$

or with Jekeli's renormalization  $S_{nm}\left(\frac{u}{E}\right)/S_{nm}\left(\frac{b}{E}\right) = Q_{nm}\left(\frac{u}{E}\right)/Q_{nm}\left(\frac{b}{E}\right)$

## "Normal" derivatives ( $z$ axis in LNOF)

$$\begin{aligned} T_r^s &= \frac{\partial V^s}{\partial r} \approx \frac{\partial V^e}{\partial \bar{z}} = T_{\bar{z}}^e \\ T_{rr}^s &= \frac{\partial^2 V^s}{\partial r^2} \approx \frac{\partial^2 V^e}{\partial \bar{z}^2} = T_{\bar{z}\bar{z}}^e \end{aligned}$$

## Spherical harmonics

very well represent a functional  $f = f(r, \theta, \lambda)$  on Earth  $\sim$  spherical approximation

$$T^s = \frac{GM}{a} \sum_{n,m} \left(\frac{a}{r}\right)^{n+1} (C_{nm}^s \cos m\lambda + S_{nm}^s \sin m\lambda) P_{nm}(\cos \theta) \quad (1)$$

## Ellipsoidal harmonics

are much closer to Earth's geometry, functional  $f = f(u, \vartheta, \lambda)$

$$T^e = \frac{GM}{a} \sum_{n,m} \frac{Q_{nm}\left(\frac{u}{E}\right)}{Q_{nm}\left(\frac{b}{E}\right)} (C_{nm}^e \cos m\lambda + S_{nm}^e \sin m\lambda) P_{nm}(\cos \vartheta) \quad (2)$$

or with Jekeli's renormalization  $S_{nm}\left(\frac{u}{E}\right)/S_{nm}\left(\frac{b}{E}\right) = Q_{nm}\left(\frac{u}{E}\right)/Q_{nm}\left(\frac{b}{E}\right)$

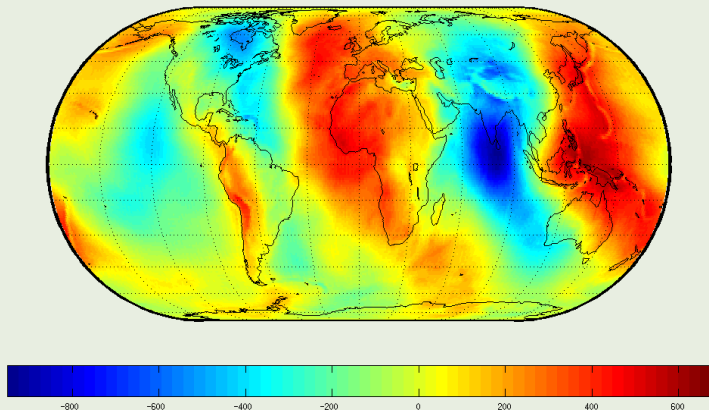
## "Normal" derivatives ( $z$ axis in LNOF)

$$\begin{aligned} T_r^s &= \frac{\partial V^s}{\partial r} \approx \frac{\partial V^e}{\partial \bar{z}} = T_{\bar{z}}^e \\ T_{rr}^s &= \frac{\partial^2 V^s}{\partial r^2} \approx \frac{\partial^2 V^e}{\partial \bar{z}^2} = T_{\bar{z}\bar{z}}^e \end{aligned}$$

## Spherical and ellipsoidal harmonic analysis of $T$ on the geoid

- Disturbing potential on the geoid, ITG03 model,  $N_{max} = 180$
- Ellipsoidal and spherical analysis  $N_{max} = 180$

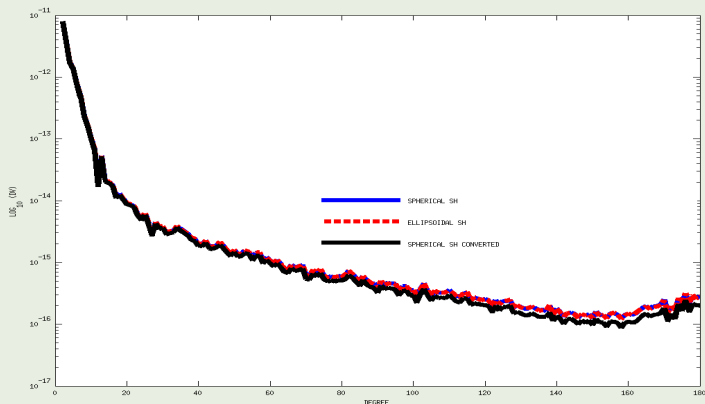
$T^s$  from SHS, [ $m^2 \cdot s^{-2}$ ]



## Spherical and ellipsoidal harmonic analysis of $T$ on the geoid

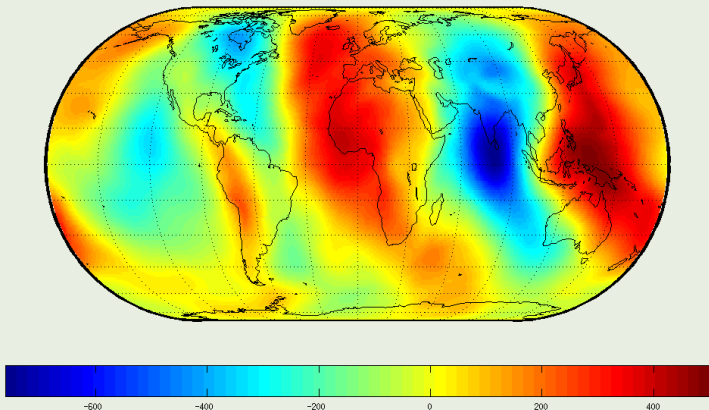
- Disturbing potential on the geoid, ITG03 model,  $N_{max} = 180$
- Ellipsoidal and spherical analysis  $N_{max} = 180$

### Degree variances of the derived coefficients



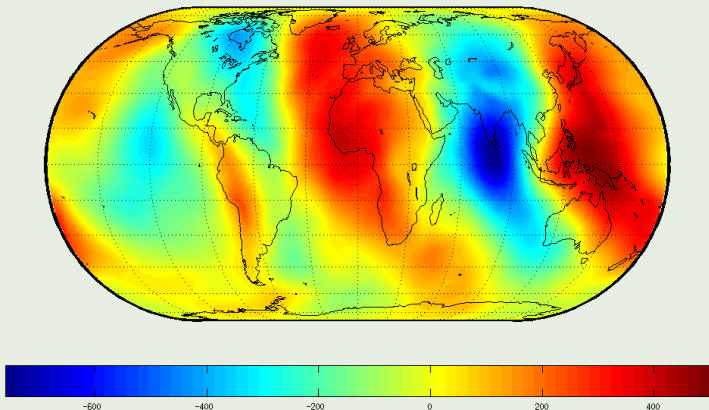
Synthesis of  $T^s$ ,  $T^e$ ,  $T_{conv}^s$ ,  $u = b + 255$  km,  $N_{max} = 180$

$T^s$  from SHS, [ $m^2 \cdot s^{-2}$ ]



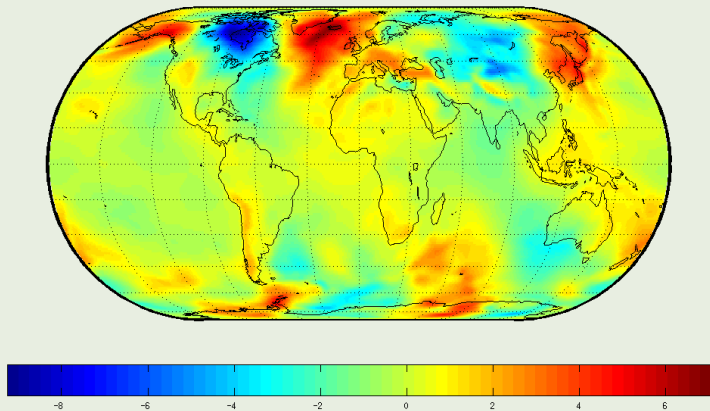
Synthesis of  $T^s$ ,  $T^e$ ,  $T_{conv}^s$ ,  $u = b + 255$  km,  $N_{max} = 180$

$T^e$  from EHS, [ $m^2 \cdot s^{-2}$ ]



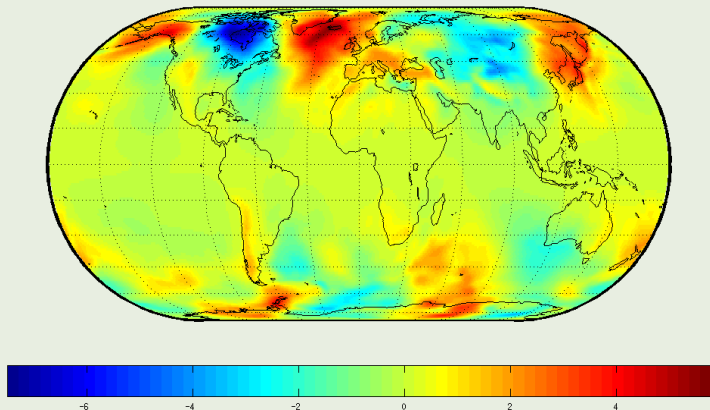
Synthesis of  $T^s$ ,  $T^e$ ,  $T_{conv}^s$ ,  $u = b + 255$  km,  $N_{max} = 180$

$$T^s - T^e, \text{RMS} = 1.74 \text{ m}^2 \cdot \text{s}^{-2}$$



Synthesis of  $T^s$ ,  $T^e$ ,  $T_{conv}^s$ ,  $u = b + 255$  km,  $N_{max} = 180$

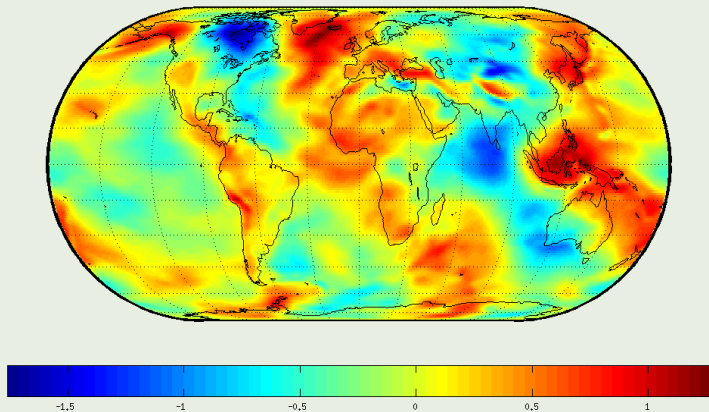
$$T^s - T_{conv}^s, \text{RMS} = 1.37 \text{ m}^2 \cdot \text{s}^{-2}$$

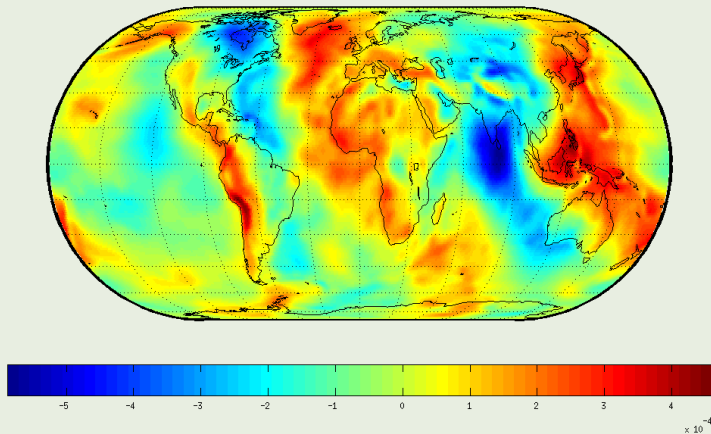




Synthesis of  $T^s$ ,  $T^e$ ,  $T_{conv}^s$ ,  $u = b + 255$  km,  $N_{max} = 180$

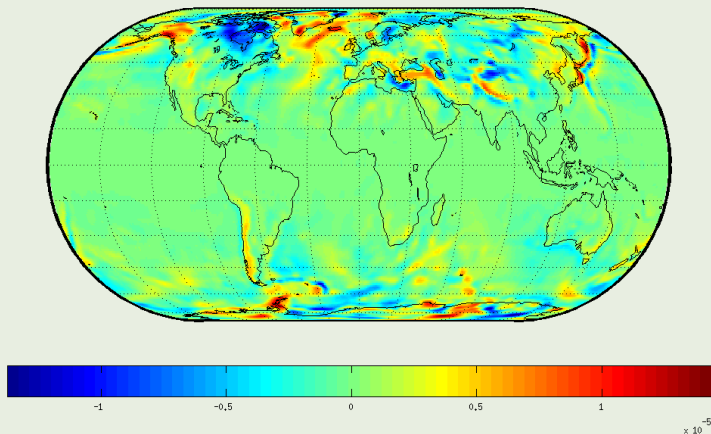
$$T_{conv}^s - T^e, \text{RMS} = 0.41 \text{ m}^2 \cdot \text{s}^{-2}$$



Synthesis of  $T_r, T_z$ ,  $u = b + 255 \text{ km}$ ,  $N_{max} = 180$  $T_r^s$  from SHS, [ $m \cdot s^{-2}$ ]

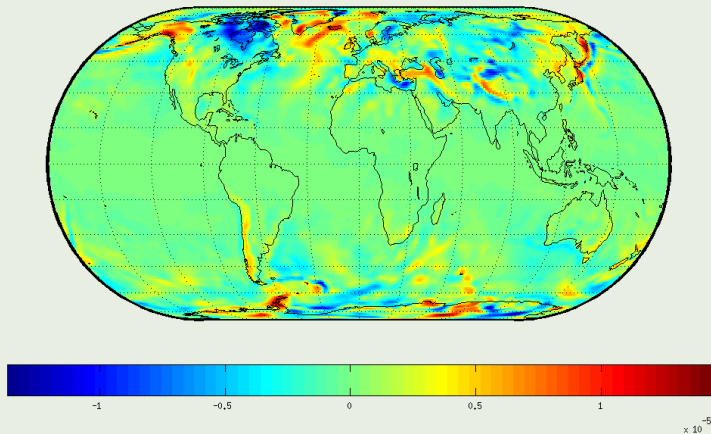
Synthesis of  $T_r, T_z$ ,  $u = b + 255 \text{ km}$ ,  $N_{max} = 180$

$$T_r^s - T_z^e, \text{RMS} = 0.253 \text{ mGal}$$



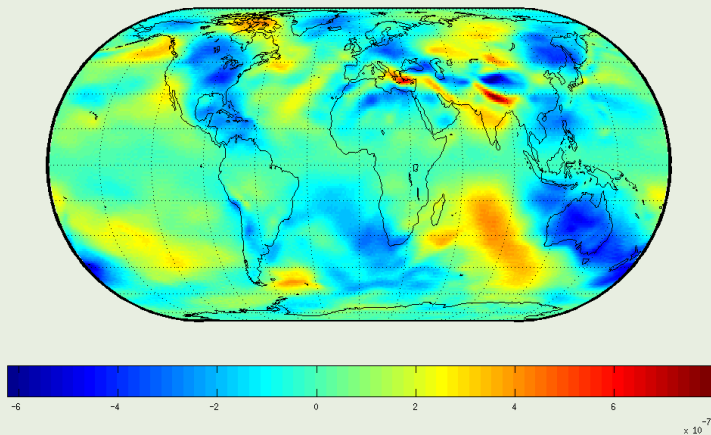
Synthesis of  $T_r, T_z$ ,  $u = b + 255 \text{ km}, N_{max} = 180$

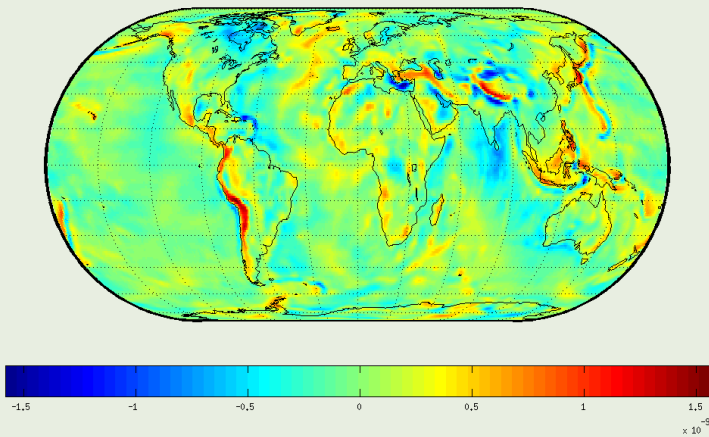
$$T_r^s - T_{r,conv}^s, \text{ RMS} = 0.255 \text{ mGal}$$



Synthesis of  $T_r, T_z$ ,  $u = b + 255 \text{ km}, N_{max} = 180$

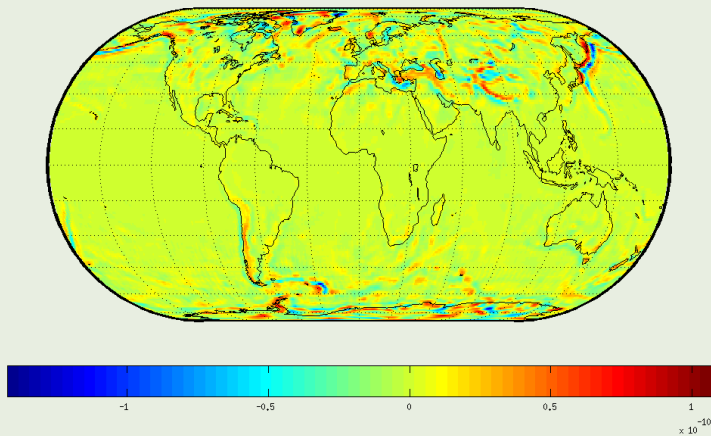
$$T_{r,conv}^s - T_z^e, \text{ RMS} = 0.014 \text{ mGal}$$



Synthesis of  $T_{rr}, T_{zz}$  at  $u = b + 255$  km $T_{rr}^s$  from SHS, [ $s^{-2}$ ]

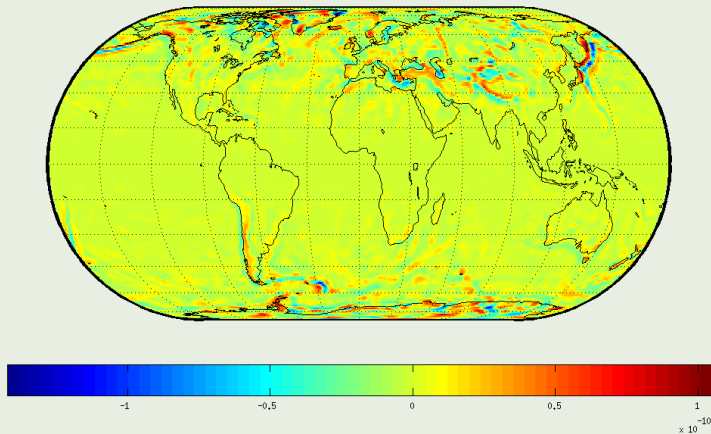
Synthesis of  $T_{rr}, T_{zz}$  at  $u = b + 255$  km

$$T_{rr}^s - T_{zz}^e, \text{RMS} = 15.47 \text{ mE}$$



Synthesis of  $T_{rr}, T_{zz}$  at  $u = b + 255$  km

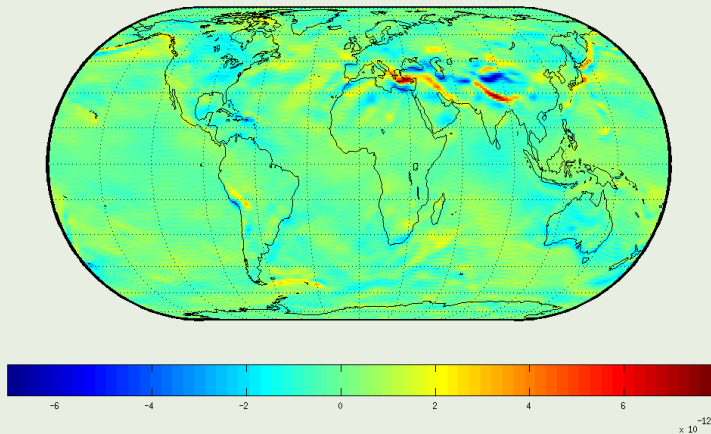
$$T_{rr}^s - T_{rr,conv}^s, \text{ RMS} = 15.44 \text{ mE}$$





Synthesis of  $T_{rr}, T_{zz}$  at  $u = b + 255$  km

$$T_{rr,conv}^s - T_{zz}^e, \text{ RMS} = 0.69 \text{ mE}$$



## Concluding remarks

RMS of diffs.	$T$ [ $m^2 \cdot s^{-2}$ ]	$T_r, T_{\bar{z}}$ [ $mGal$ ]	$T_{rr}, T_{\bar{z}\bar{z}}$ [ $mE$ ]
Sph - Ell	1.74	0.253	15.5
Sph - Sph(conv.)	1.37	0.255	15.4
Ell - Sph(conv.)	<b>0.41</b>	<b>0.014</b>	<b>0.7</b>

- We have compared three sets of coefficients coming from one input (2x spherical and 1x ellipsoidal) via harmonic synthesis on the  $u = b + 255$  km for three orders of derivative of  $T$ .
- Good agreement achieved when SHS with converted coefficients and EHS were used
- $\implies$  when validation uses the global approximation of the ground data, EH and SH(converted) "suit" more to this task
- $\implies$  global gravity field models based on the ellipsoidal analysis might have principal advantages (e.g. EGM08)

## Concluding remarks

RMS of diffs.	$T$ [ $m^2 \cdot s^{-2}$ ]	$T_r, T_{\bar{z}}$ [ $mGal$ ]	$T_{rr}, T_{\bar{z}\bar{z}}$ [ $mE$ ]
Sph - Ell	1.74	0.253	15.5
Sph - Sph(conv.)	1.37	0.255	15.4
Ell - Sph(conv.)	<b>0.41</b>	<b>0.014</b>	<b>0.7</b>

- We have compared three sets of coefficients coming from one input (2x spherical and 1x ellipsoidal) via harmonic synthesis on the  $u = b + 255$  km for three orders of derivative of  $T$ .
- Good agreement achieved when SHS with converted coefficients and EHS were used
- $\implies$  when validation uses the global approximation of the ground data, EH and SH(converted) "suit" more to this task
- $\implies$  global gravity field models based on the ellipsoidal analysis might have principal advantages (e.g. EGM08)

## Concluding remarks

RMS of diffs.	$T$ [ $m^2 \cdot s^{-2}$ ]	$T_r, T_{\bar{z}}$ [ $mGal$ ]	$T_{rr}, T_{\bar{z}\bar{z}}$ [ $mE$ ]
Sph - Ell	1.74	0.253	15.5
Sph - Sph(conv.)	1.37	0.255	15.4
Ell - Sph(conv.)	<b>0.41</b>	<b>0.014</b>	<b>0.7</b>

- We have compared three sets of coefficients coming from one input (2x spherical and 1x ellipsoidal) via harmonic synthesis on the  $u = b + 255$  km for three orders of derivative of  $T$ .
- Good agreement achieved when SHS with converted coefficients and EHS were used
- $\implies$  when validation uses the global approximation of the ground data, EH and SH(converted) "suit" more to this task
- $\implies$  global gravity field models based on the ellipsoidal analysis might have principal advantages (e.g. EGM08)

## Concluding remarks

RMS of diffs.	$T$ [ $m^2 \cdot s^{-2}$ ]	$T_r, T_{\bar{z}}$ [ $mGal$ ]	$T_{rr}, T_{\bar{z}\bar{z}}$ [ $mE$ ]
Sph - Ell	1.74	0.253	15.5
Sph - Sph(conv.)	1.37	0.255	15.4
Ell - Sph(conv.)	<b>0.41</b>	<b>0.014</b>	<b>0.7</b>

- We have compared three sets of coefficients coming from one input (2x spherical and 1x ellipsoidal) via harmonic synthesis on the  $u = b + 255$  km for three orders of derivative of  $T$ .
- Good agreement achieved when SHS with converted coefficients and EHS were used
- $\Rightarrow$  when validation uses the global approximation of the ground data, EH and SH(converted) "suit" more to this task
- $\Rightarrow$  global gravity field models based on the ellipsoidal analysis might have principal advantages (e.g. EGM08)

## Concluding remarks

RMS of diffs.	$T$ [ $m^2 \cdot s^{-2}$ ]	$T_r, T_{\bar{z}}$ [ $mGal$ ]	$T_{rr}, T_{\bar{z}\bar{z}}$ [ $mE$ ]
Sph - Ell	1.74	0.253	15.5
Sph - Sph(conv.)	1.37	0.255	15.4
Ell - Sph(conv.)	<b>0.41</b>	<b>0.014</b>	<b>0.7</b>

- We have compared three sets of coefficients coming from one input (2x spherical and 1x ellipsoidal) via harmonic synthesis on the  $u = b + 255$  km for three orders of derivative of  $T$ .
- Good agreement achieved when SHS with converted coefficients and EHS were used
- $\Rightarrow$  when validation uses the global approximation of the ground data, EH and SH(converted) "suit" more to this task
- $\Rightarrow$  global gravity field models based on the ellipsoidal analysis might have principal advantages (e.g. **EGM08**)

**Thank you!**

# Synthesis of $|\nabla T^s|$ , $|\nabla T^e|$ , $u = b + 255 \text{ km}$ , $N_{max} = 180$

$$|\nabla T^s|^2 = \left(\frac{\partial T}{\partial r}\right)^2 + \left(\frac{1}{r} \frac{\partial T}{\partial \theta}\right)^2 + \left(\frac{1}{r \sin \theta} \frac{\partial T}{\partial \lambda}\right)^2$$

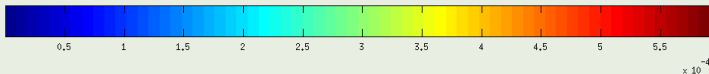
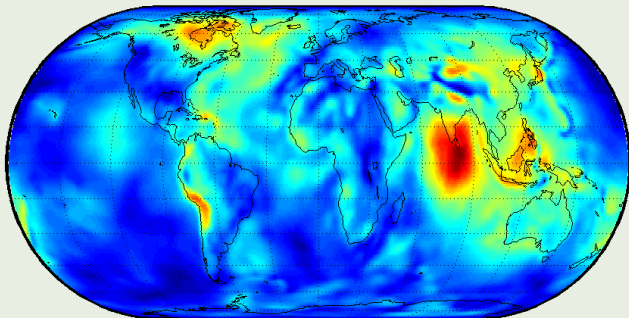
$$|\nabla T^e|^2 = \left(\frac{1}{w} \frac{\partial T}{\partial u}\right)^2 + \left(\frac{1}{w \sqrt{u^2 + E^2}} \frac{\partial T}{\partial \vartheta}\right)^2 + \left(\frac{1}{\sqrt{u^2 + E^2} \sin \vartheta} \frac{\partial T}{\partial \lambda}\right)^2$$

$$w = \sqrt{\frac{u^2 + E^2 \cos^2 \vartheta}{u^2 + E^2}}$$



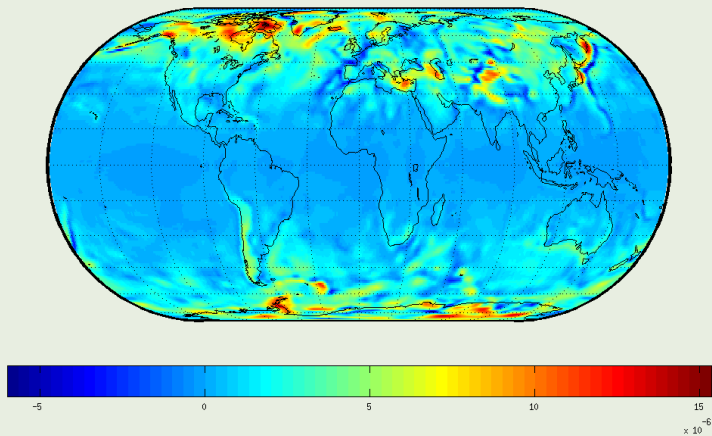
Synthesis of  $|\nabla T^s|$ ,  $|\nabla T^e|$ ,  $u = b + 255 \text{ km}$ ,  $N_{max} = 180$

$|\nabla T^s|$  from SHS,  $[m \cdot s^{-2}]$



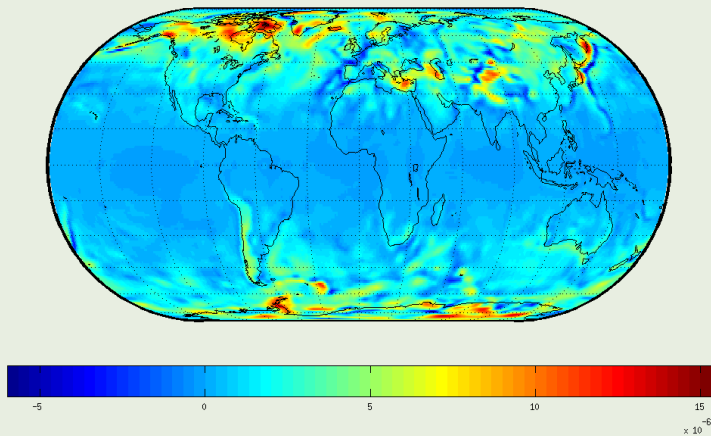
Synthesis of  $|\nabla T^s|$ ,  $|\nabla T^e|$ ,  $u = b + 255 \text{ km}$ ,  $N_{max} = 180$

$|\nabla T^s| - |\nabla T^e|$ , RMS = 0.307 mGal



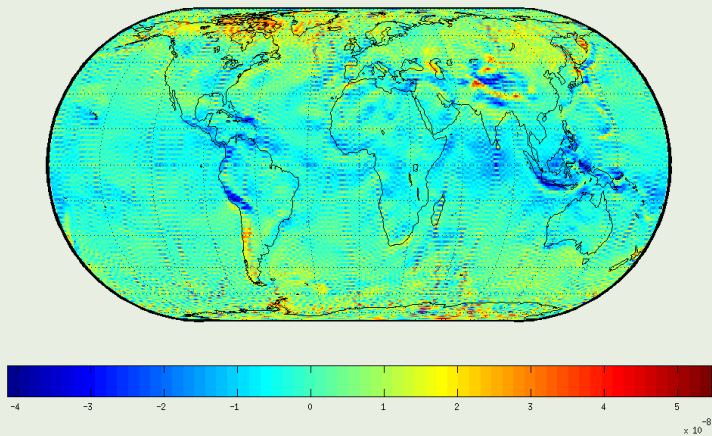
Synthesis of  $|\nabla T^s|$ ,  $|\nabla T^e|$ ,  $u = b + 255 \text{ km}$ ,  $N_{max} = 180$

$|\nabla T^s| - |\nabla T^s_{conv}|$ , RMS = 0.307 mGal



Synthesis of  $|\nabla T^s|$ ,  $|\nabla T^e|$ ,  $u = b + 255 \text{ km}$ ,  $N_{max} = 180$

$|\nabla T^s_{conv}| - |\nabla T^e|$ , RMS = 0.001 *mGal*



$$T_{\bar{z}} = \frac{GM}{a} \frac{v}{L} \sum_{n,m} \frac{\frac{\partial S_{nm}(\frac{u}{E})}{\partial u}}{S_{nm}(\frac{b}{E})} (A_{nm} \cos m\lambda + B_{nm} \sin m\lambda) P_{nm}(\cos \vartheta)$$

$$T_{\bar{z}\bar{z}} = \frac{GM}{a} \frac{v^2}{L^2} \sum_{n,m} \frac{\frac{\partial^2 S_{nm}(\frac{u}{E})}{\partial u^2}}{S_{nm}(\frac{b}{E})} (A_{nm} \cos m\lambda + B_{nm} \sin m\lambda) P_{nm}(\cos \vartheta)$$

$$- \frac{GM}{a} \frac{uE^2 \sin^2 \vartheta}{L^4} \sum_{n,m} \frac{\frac{\partial S_{nm}(\frac{u}{E})}{\partial u}}{S_{nm}(\frac{b}{E})} (A_{nm} \cos m\lambda + B_{nm} \sin m\lambda) P_{nm}(\cos \vartheta)$$

$$+ \frac{GM}{a} \frac{E^2 \sin \vartheta \cos \vartheta}{L^4} \sum_{n,m} \frac{S_{nm}(\frac{u}{E})}{S_{nm}(\frac{b}{E})} (A_{nm} \cos m\lambda + B_{nm} \sin m\lambda) \frac{\partial P_{nm}(\cos \vartheta)}{\partial \vartheta}$$

$$v^2 = u^2 + E^2$$

$$L^2 = u^2 + E^2 \cos^2 \vartheta$$