



esa

SPPA

Sensor Performance, Products and Algorithms

Atmospheric Correction and Vicarious Calibration of MERIS/OLCI

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Content

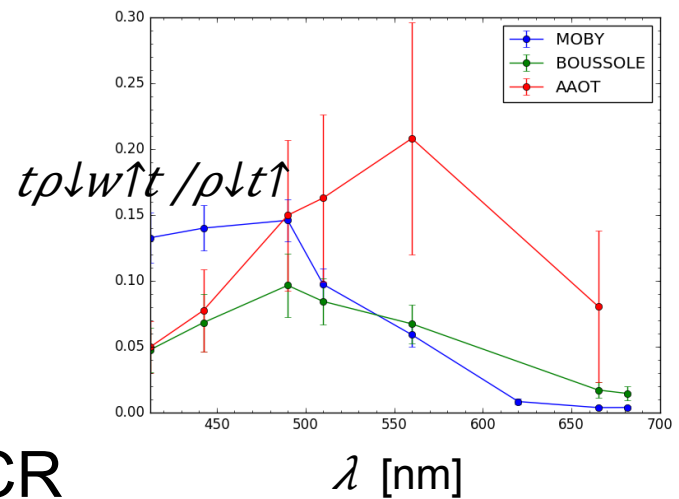
- Atmospheric correction: principle, Black Pixel vs Bright Pixel
- Bright Pixel Atmospheric Correction (BPAC): model, issues, algorithm, validation
- System Vicarious calibration (SVC): MERIS, uncertainties, OLCI

Ocean Colour Radiometry (OCR)

- Colour of the ocean with the atmosphere above

$$\rho_{\downarrow t \uparrow}(\lambda) = t_{\downarrow g \uparrow}(\lambda)(\rho_{\downarrow R \uparrow}(\lambda) + \rho_{\downarrow a \uparrow}(\lambda) + \rho_{\downarrow a R \uparrow}(\lambda) + T\rho_{\downarrow G \uparrow}(\lambda) + t\rho_{\downarrow w \uparrow}(\lambda))$$

Marine reflectance $\rho_{\downarrow w \uparrow}$ is only a small part of the TOA signal (~ 10%)



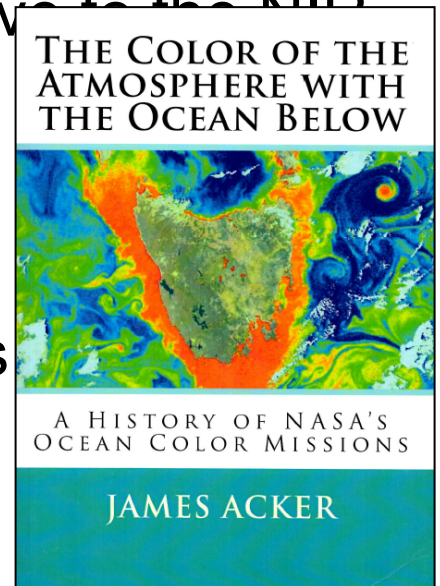
- The sources of uncertainty for OCR

- Atmospheric scattering functions
- Calibration, including SVC

$$\rho_{\downarrow w \uparrow}(\lambda) = g(\lambda) * \rho_{\downarrow gc \uparrow}(\lambda) - \rho_{\downarrow path \uparrow}(\lambda) / t(\lambda)$$

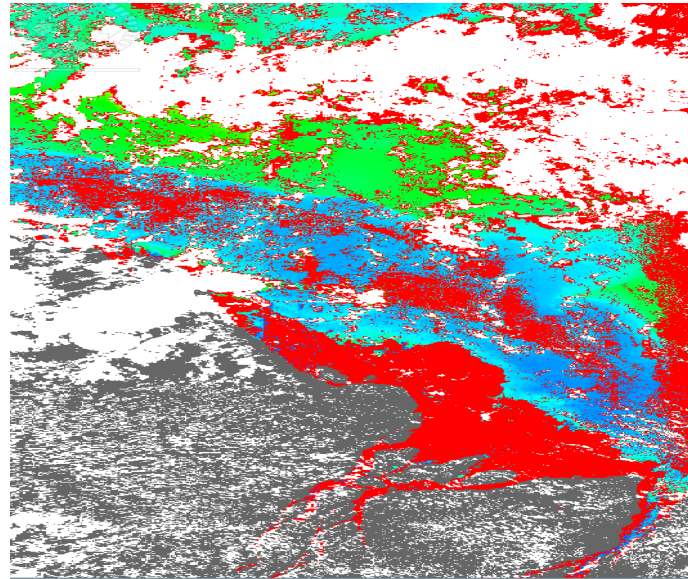
Atmospheric correction

- *Standard AC* used in MERIS/OLCI (Antoine & Morel 1999)
 - Black pixel assumption: aerosol detection in the NIR
 - Pros: no assumption on marine modelling in the VIS
 - Cons: aerosol detection is highly sensitive to the NIR
- The case of Bright pixel
 - Marine signal in the NIR
 - Need to decouple atm. & marine signals



3RP BPAC issues

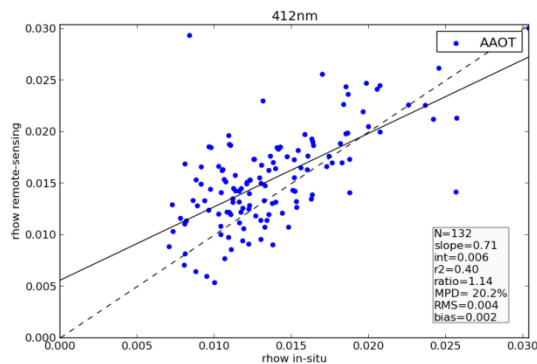
- Negative reflectance for (very) turbid waters



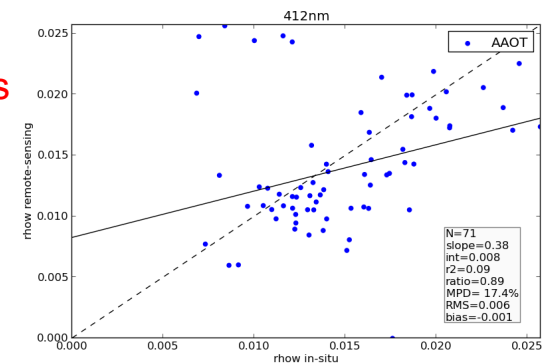
ρ_w (412)

PCD_1_13

- Too strong sensitivity to NIR calibration



Applying NIR vicarious gains



The BPAC model (Moore 2011)

1) Marine IOPs
+ Petzold
phase function

$$b_{bp}(\lambda) = b_{bp}(\lambda_0) \left(\frac{\lambda}{\lambda_0} \right)^{-S_b}$$

$$a_p(\lambda) = a_p(\lambda_0) e^{-S_a(\lambda - \lambda_0)}$$

2) Marine AOP
through radiative
transfer

$$\omega(\lambda) = \frac{b_{bw}(\lambda) + b_{bp}(\lambda)}{a_w(\lambda) + a_p(\lambda) + b_{bw}(\lambda) + b_{bp}(\lambda)}$$

$$\eta(\lambda) = \frac{b_{bw}(\lambda)}{b_{bw}(\lambda) + b_{bp}(\lambda)}$$

$$F'(\lambda, \eta, \omega) = A_0(\lambda) + A_1(\lambda)\eta + a_0(\lambda) + a_1(\lambda)\omega + a_2(\lambda)\omega^2 + a_3(\lambda)\omega^3 + a_4(\lambda)\omega^4$$

$$\rho_w(\lambda) = F'(\lambda, \eta, \omega) * \omega(\lambda)$$

3) Atmospheric
& TOA model

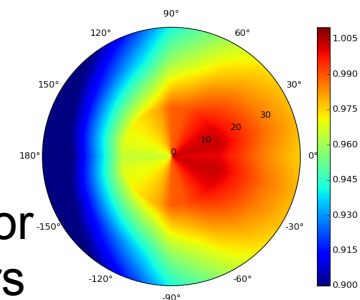
$$t(\lambda) = e^{-\left(0.5\tau_R(\lambda) + (1-f\omega_a)\tau_a(\lambda_0)\left(\frac{\lambda}{\lambda_0}\right)^\alpha\right)M}$$

$$\rho_a(\lambda) = \rho_a(\lambda_0) \left(\frac{\lambda}{\lambda_0} \right)^\alpha$$

$$\rho_{RC}(\lambda) = t(\lambda)\rho_w(\lambda) + \rho_a(\lambda)$$

4) Free variables, justified
by a sensitivity analysis:

$$b_{bp}(\lambda_0), \rho_a(\lambda_0), \alpha$$

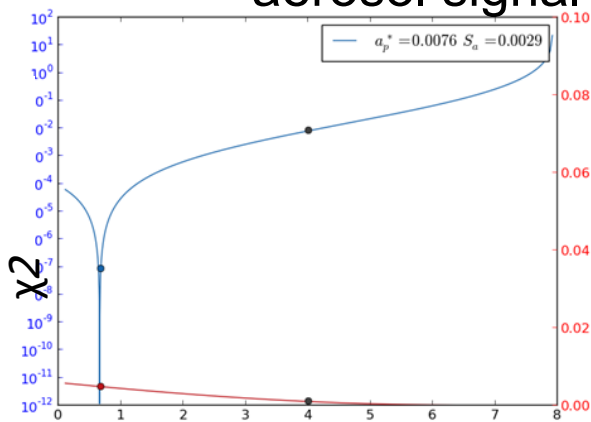


Remark: this model can be used for
BRDF correction over turbid waters

Understanding the BPAC issues

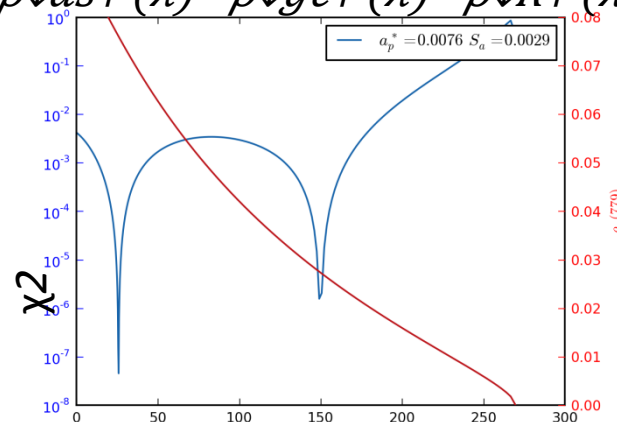
1. Ambiguity in the BPAC model

- Several solutions may be possible depending on TSM and bands
- For any TSM, compute the χ^2 of the log-log fit of the remaining aerosol signal $\rho_{\downarrow a s \uparrow}(\lambda) = \rho_{\downarrow g c \uparrow}(\lambda) - \rho_{\downarrow R \uparrow}(\lambda) - t(\lambda)\rho_{\downarrow R \uparrow}(\lambda)$



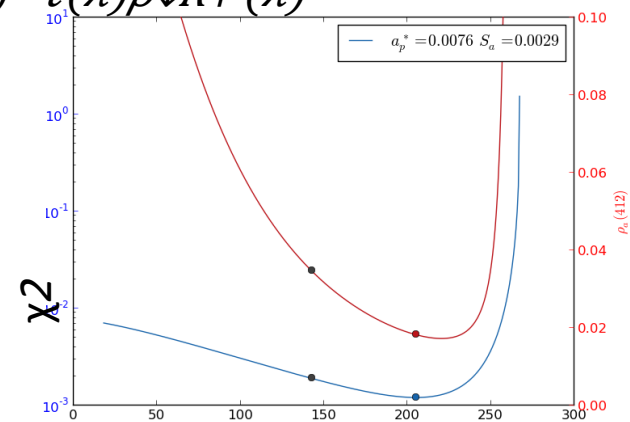
TSM [mg/L]

Clear water: unique solution when considering 709+779+865



TSM [mg/L]

Very turbid water: two solutions when considering 709+779+865



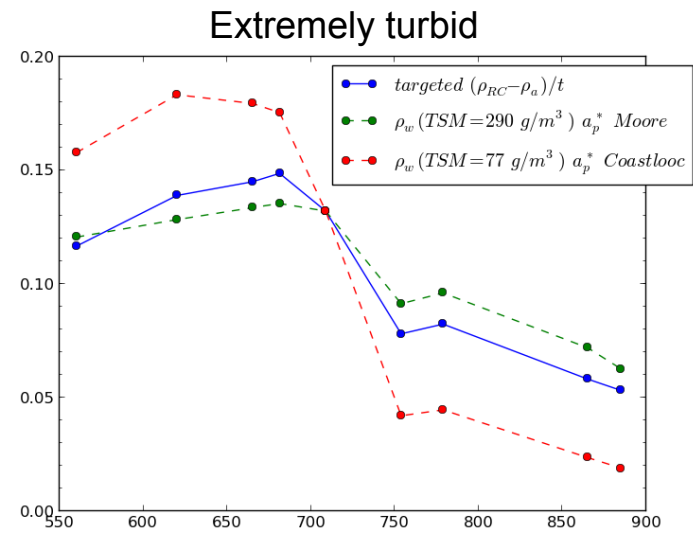
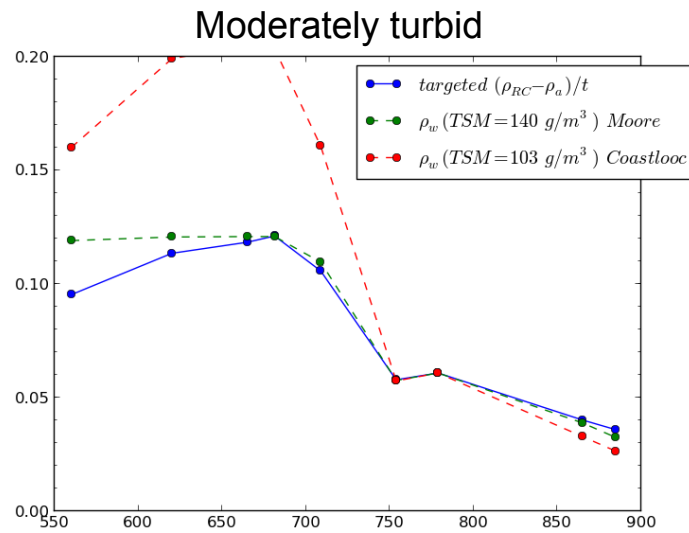
TSM [mg/L]

Very turbid water: unique solution when considering the five NIR bands

Understanding the BPAC issues

2. Suitability of the marine model with respect to a_p

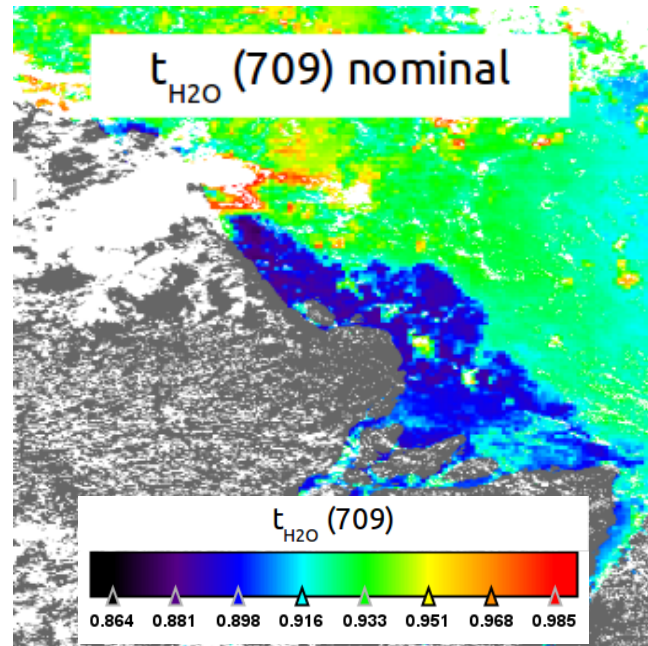
- Particulate absorption is the most important fixed parameter over very turbid water
- Various a_p exist in the literature



Understanding the BPAC issues

3. Water vapour issue

- MERIS band 709 nm is theoretically very useful for remote-sensing of coastal waters, but:
 - It is impacted by H₂O absorption (885 nm as well)
 - H₂O correction is based on the 900/885 ratio using the “black pixel” assumption



TOA errors over the Amazon:

510 nm: ~ 0.6%

681 nm: ~ 0.1%

753 nm: ~ 0.1%

709 nm: ~ 3.5%

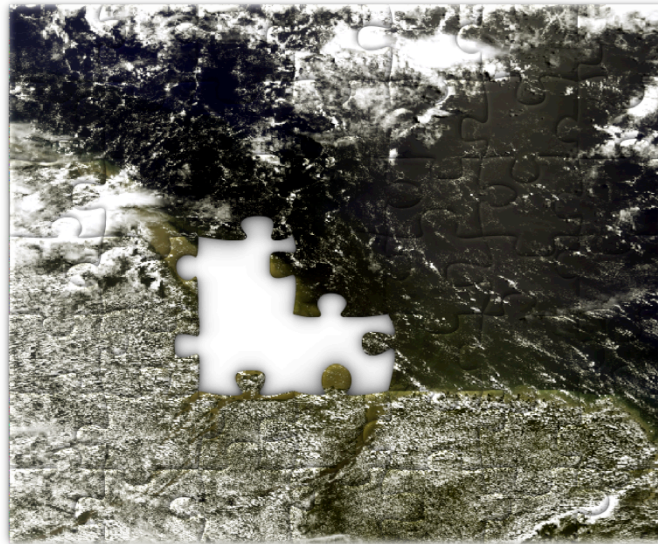
779 nm: ~ 0.1%

865 nm: ~ 0.2%

885 nm: ~ 1.5%

Understanding the BPAC issues

1. Ambiguity in the BPAC model



2. Suitability of the marine model with respect to a_p

3. Water vapour issue

Algorithm for MERIS 4th reproc. and OLCI (ATBD 2.6)

1. Ambiguity in the BPAC model

→ Removed when χ^2 mainly constructed by 753, 779, 865 nm

$$\chi^2 = \sum_{\lambda} \left(\frac{r(\lambda)}{\sigma(\lambda)} \right)^2$$

$$r(\lambda) = t(\lambda)\rho_w(\lambda) + \rho_a(\lambda_0) \left(\frac{\lambda}{\lambda_0} \right)^\alpha - \rho_{RC}(\lambda)$$

$\sigma^2(709)$ and $\sigma^2(885)$ growing with TSM

No more iterative first guess with TSM=0

3. Water vapour issue

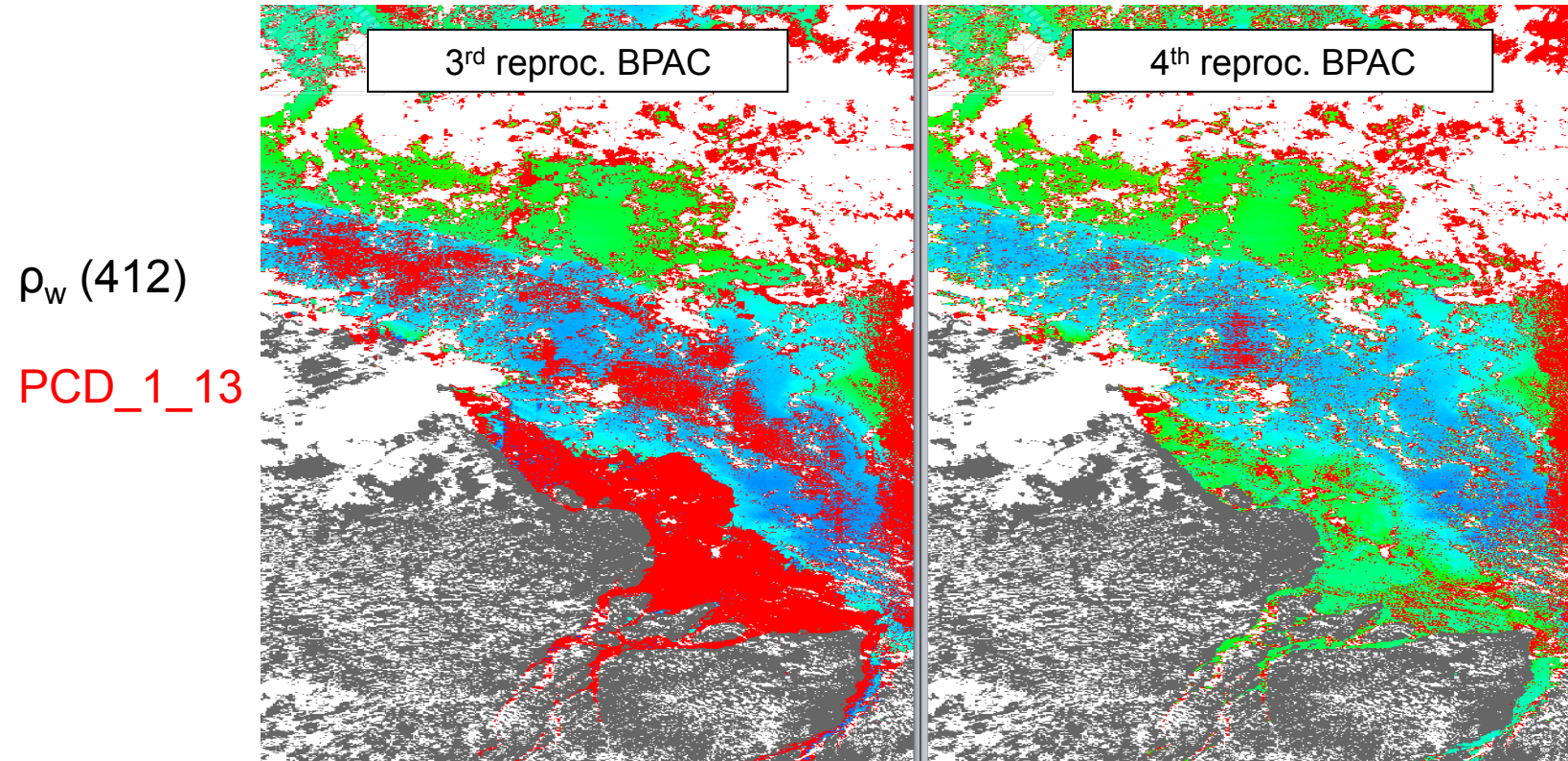
→ Does not impact χ^2

2. Suitability of the marine model with respect to a_p

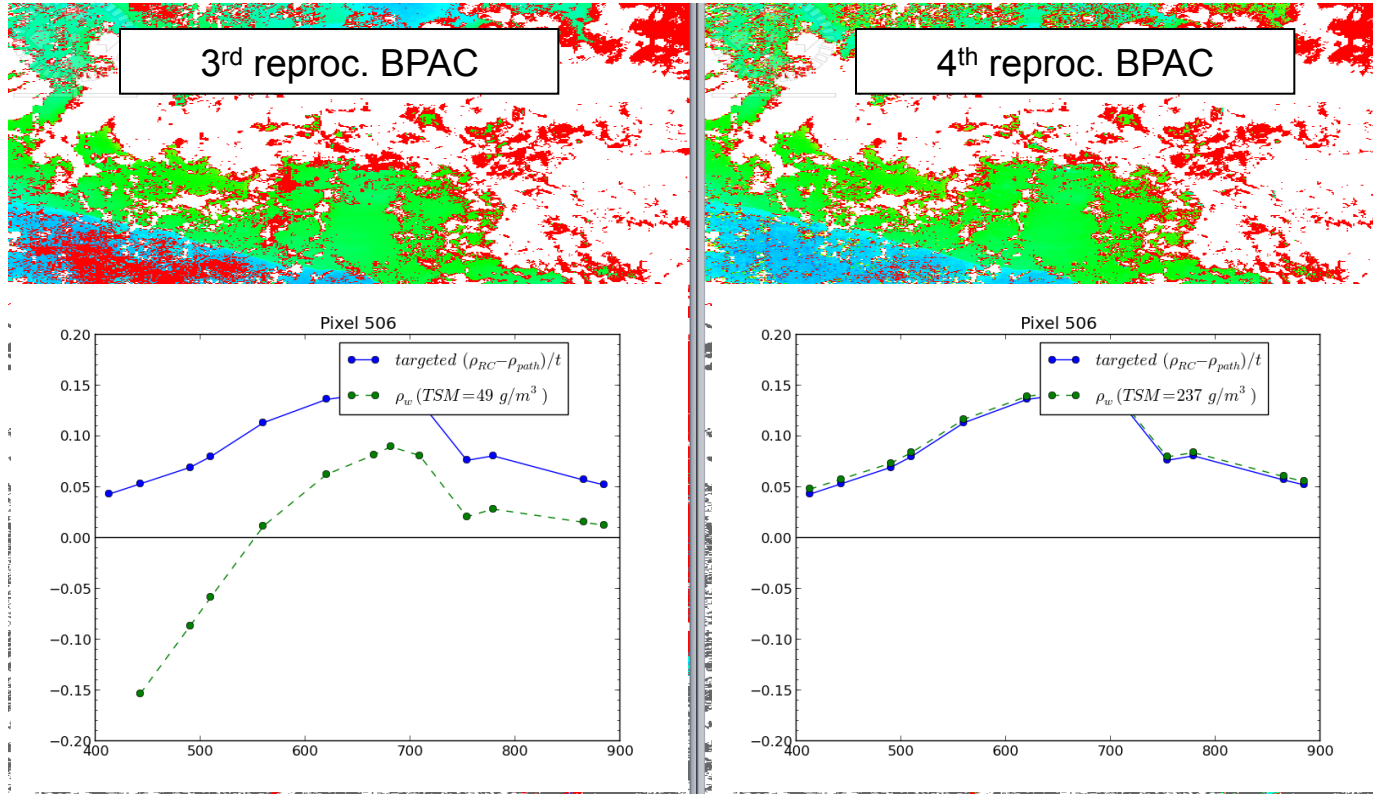
→ Sediment absorption weakly impacts χ^2 at high turbidity

Other conclusion for MERIS: NIR vicarious calibration (inter band) is no more strictly required

Results: the Amazone plume

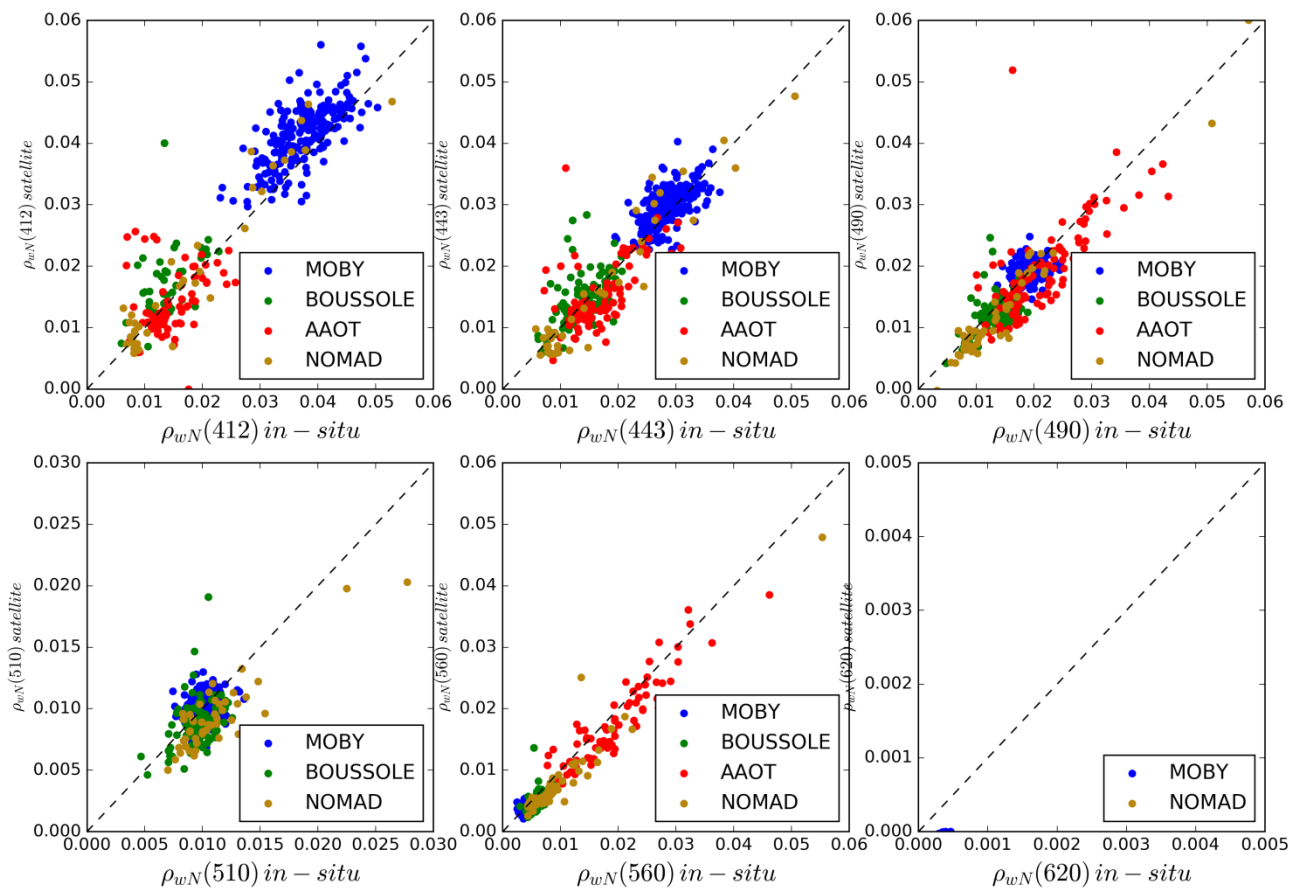


Results: the Amazone plume



Validation: MERMAID matchups

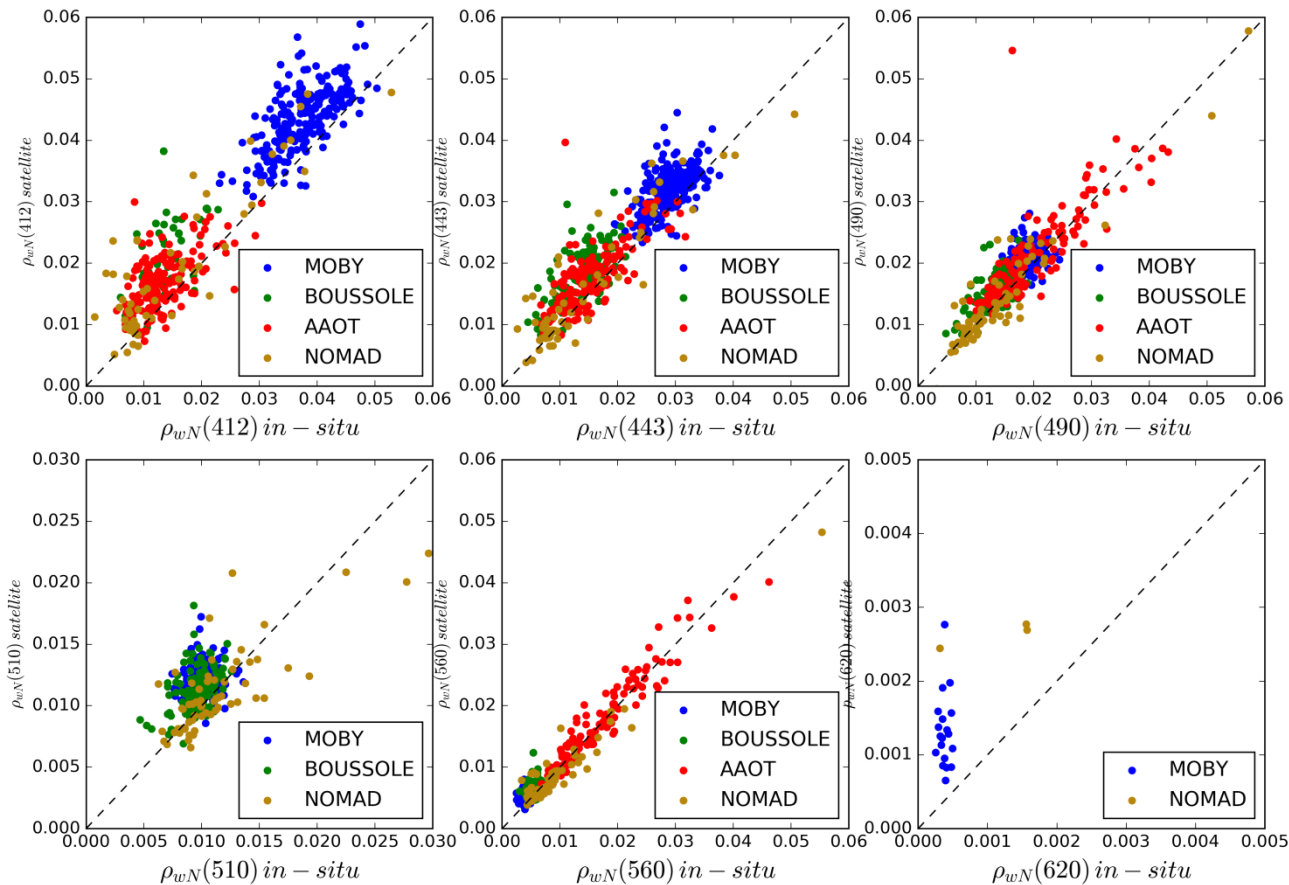
3RP BPAC (before VIS vicarious cal.)



Acknowledgements to ACRI & ESA for MERMAID and to PIs: G. Zibordi, JRC (AAOT), D. Antoine, LOV (BOUSSOLE), K. Voss, Univ. Miami (MOBY), Werdell and Bailey 2005(NOMAD)

Validation: MERMAID matchups

4RP BPAC (before VIS vicarious cal.)



More consistent biases between different water types \rightarrow there is a chance to have a global positive impact of VIS vicarious cal.

Acknowledgements to ACRI & ESA for MERMAID and to PIs: G. Zibordi, JRC (AAOT), D. Antoine, LOV (BOUSSOLE), K. Voss, Univ. Miami (MOBY), Werdell and Bailey 2005(NOMAD)

Validation: very turbid waters

In collaboration with Pierre Gernez, Univ. Nantes

- Bourgneuf Bay
- In situ TSM from Dutertre et al 2009

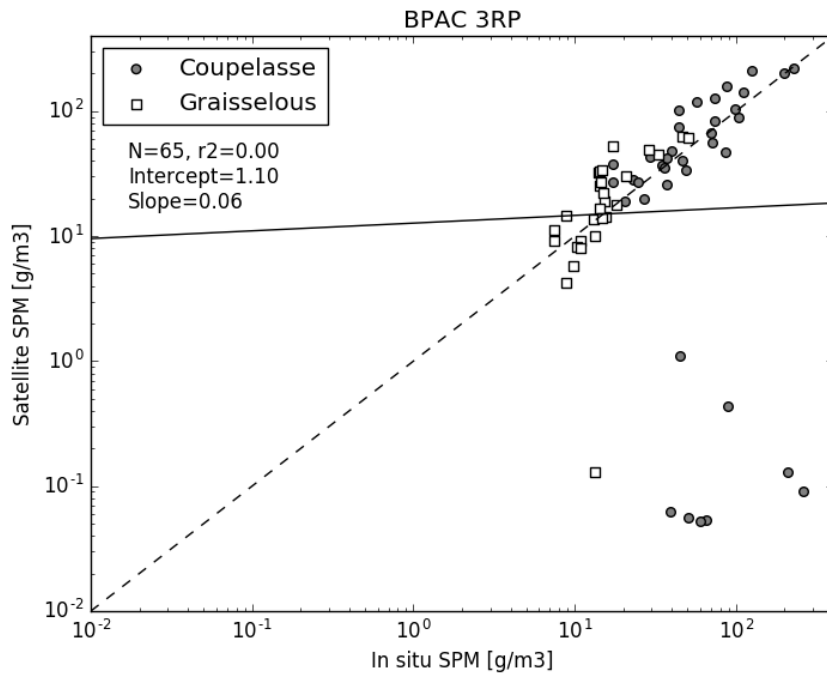


- Data & method:
 - MERIS Full Resolution data (MERCIS)
 - Match-up extraction (BEAM)
 - Pixel selection (closest underwater pixel to the station)
 - Time interpolation of in situ TSM to MERIS time
 - Processing (ODEAS)
 - Regional TSM model from marine reflectance (Gernel et al. 2014)

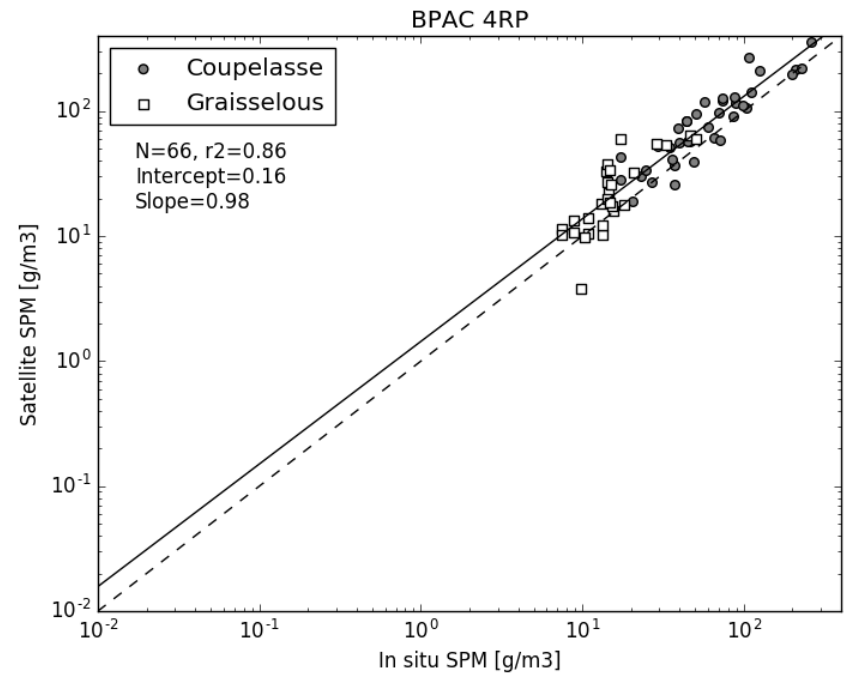
Validation: very turbid waters

In collaboration with Pierre Gernez, Univ. Nantes

3RP BPAC

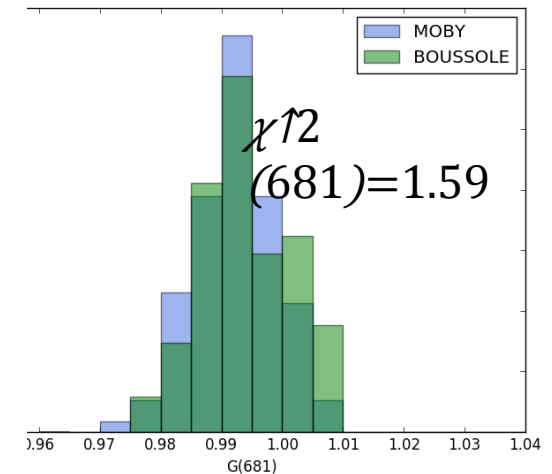
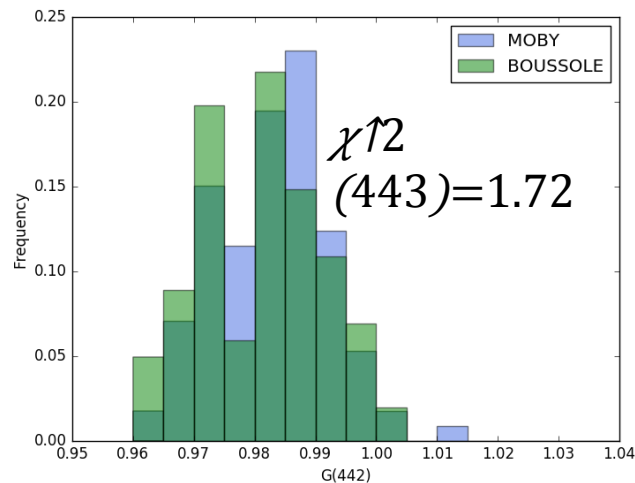
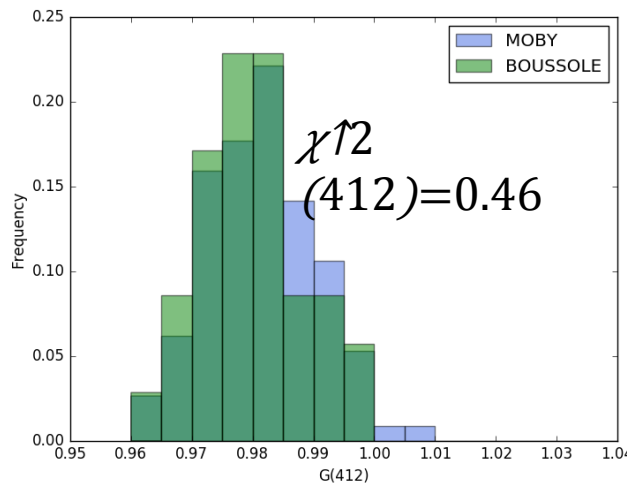


4RP BPAC



MERIS SVC for 4RP

- SVC gain definition: $g(\lambda) = (t(\lambda)\rho \downarrow w \uparrow t(\lambda) + \rho \downarrow path \uparrow(\lambda)) / \rho \downarrow gc \uparrow(\lambda)$
- Questions: uncertainties, choice of site(s) and consistency
- MOBY and BOUSSOLE gains agreement: Chi2 test of homogeneity:
if $|g \downarrow M - g \downarrow B| / \sqrt{\sigma \downarrow M \uparrow 2 / N \downarrow M + \sigma \downarrow B \uparrow 2 / N \downarrow B} < 1.96$, there is 95% probability that both sets of gains belong to the same distribution



SVC for OLCI

- Only few matchups: can we use a L3 climatology for $\rho \downarrow w \uparrow t$?
 - Limiting factor is satellite, not in-situ (daily measurements exist at MOBY): several sites must be considered → consistency
 - Uncertainty due to using past $\rho \downarrow w \uparrow t$ data):

$$u(\rho \downarrow w \uparrow t) \approx \text{std}(\rho \downarrow w \uparrow t (\text{past}))$$

→ Stability of $\rho \downarrow w \uparrow t$ per L3 temporal bin needs to be within 5% for each site

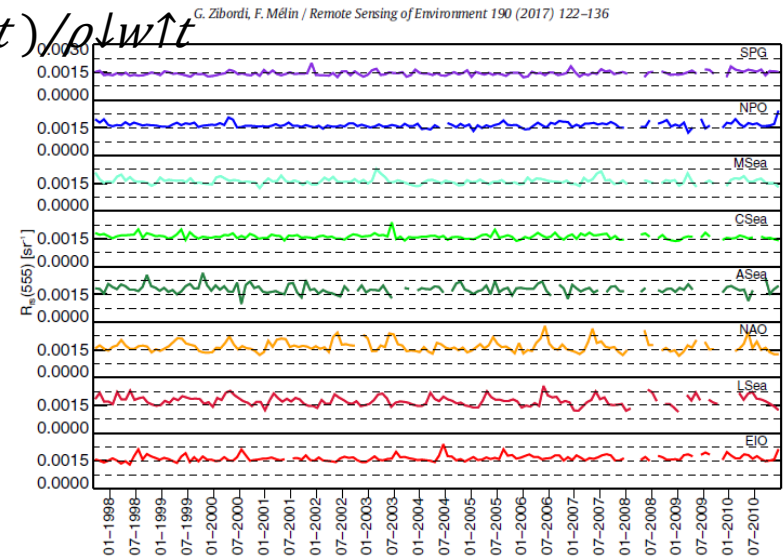


Fig. 3. Time-series of monthly values of $R_{\rho}(555)$ for the various regions of interest.

Conclusion: messages for SSPA

- 1. OCR needs an integrated algorithmic approach**
 - Vicarious calibration depends on AC
 - AC depends on BPAC – even over clear waters
 - BPAC depends on upstream steps (glint, gas, NIR cal.)
- 2. BPAC has been improved by working on the numerical inversion, while keeping same model**
- 3. Check consistency of AC with AERONET data (for SVC)**
 - See Francis Zagolski's presentation
- 4. Validation requires more match-ups and tools (data extraction).
Plan for OLCI?**