GOMOS Level 2 Product Quality Readme File

Field:	Contents:	Filled
		by:
Title	GOMOS Level 2 processor version GOMOS/6.01 Readme	SPPA
		Engineer
Reference	ENVI-GSOP-EOGD-QD-12-0117, issue 1.0	SPPA
	Date: 17 December 2012	Manager
Affected	This readme file applies to all GOM_NL_2P processed with processor version	SPPA
data sets	IPF 6.01	Engineer
Product	Algorithm Theoretical Baseline Document (ATBD)	SPPA
specification	Product Specification:	Engineer
references	PO-RS-MDA-GS-2009, Issue 3, Revision K, 15/10/2012	
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	1. GOMOS processor Version 6.01 (Level 2)	
	The GOMOS/ENVISAT processor version 6.01 introduces a number of	
	upgrades that are detailed in this document and those referenced.	
	1.1 Processor upgrades	
	The following Level 2 processor upgrades have been implemented in	
	GOMOS version 6.01:	
	Accurate characterization of modeling errors in the full covariance	
	matrix inversion: impact on error estimates and X^2	
	New HRTP (High Resolution Temperature Profile) algorithm:	
	improves the High Resolution Temperature profiles	
	New coding of the error bar (logarithmic encoding of absolute error	
	value) instead of relative error	
	,	
	1.1.1 Note on re-processed Version 6.01 Level 2 products	
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A reprocessing campaign has been performed using IPF version 6.01 of the GOMOS processor covering the period from 15th April 2002 to 8th April 2012 (end of mission). The Level 2 re-processed products are identified by the following flags reported in the MPH and in the product name:

MPH Field	Value
Processing stage flag	R
Processing center	FINPAC ('FIN' in the
Trocessing center	product name)
Software version	GOMOS/6.01
REE DOC	PO-RS-MDA-GS-
REF_DOC	2009 issue 3/K

2. GOMOS Version 6.01 Level 2 products

2.1 Product data screening

To select highest quality data for scientific application, users should use the data in the GOM_NL__2P products that satisfies the following criteria:

- GEOLOCATION(30km)/SUN_ZENITH_TANGENT > 105°
- non-flagged density values: LOCAL_SPECIES_DENSITY/PCD set to "0".

Occultations from bright limb data, and to a lesser extent from twilight limb data, are of reduced quality and should be used with care.

Note that for Ozone profiles a more detailed data screening is suggested. This can be found in section 2.5.1.

2.2 Flagging and negative data

- The product flags for the retrieved quantities are processor flags, indicating if the retrieval has been successful or not.
- There is no systematic flagging of negative values of column and local densities. Negative values of line densities after spectral inversion are kept, and non-flagged negative values may be included in the vertical profiles. It is recommended to keep those data when calculating averages. Ignoring such data may introduce a risk of creating an artificial positive bias in averages of data.
- The XYZ_std fields (where XYZ is a species name) represent uncertainty estimates assuming the Gaussian error statistics. As of version IPF 6, these values represent the random error, which includes an estimate of the modeling error (mainly from incomplete scintillation correction, see Sofieva et al., 2010). Values are expressed in absolute value (cm⁻³) of the local density and correspond to 1o. In order to obtain the value of the standard deviation in cm⁻³ it is necessary to use a coding equation, reported in the Product Specification. This allows to convert the integer number written inside the product. The maximum value of the integer number for the error estimate is set to 6553.5.

2.3 Vertical resolution

All GOMOS occultations have the same vertical resolution, because a "target resolution" Tikhonov regularization is applied in the GOMOS inversion. The vertical resolution for different retrieved species is given in

Dzore 2 km below 30 km; 3 km above 40 km NO2 4 km Aerosol 4 km H2O 2 km below 20 km; 4 km above 30 km O2 3 km below 30 km; 5 km above 40 km Table 1: Target vertical resolution by retrieved species. able 2 presents a summary of the altitude range of validity for the Level 2 roducts. More details for each product are given in Section 2.5. becies Validity/altitude range availd at all altitudes for hot stars, in the range ~12-100 km; for cold and weak stars (main) T < 7000K and visual magnitude > 1.9), data above 40 km should not be used D2 valid between 20 km and 50 km and winter polar regions up to 65 km; data at other altitudes should be considered with caution D3 valid between 25 km and 45 km, but noisy retrieved values within this altitude range; data at other altitudes should be considered with caution 20 valid between 25 km and 45 km, but noisy retrieved values with a caution 30 retrieved for 8 stars (Star_ids 1, 2, 3, 13, 14, 16, 26, 63) and up to 50 km; 20-30 km for oblique occultations; 20-30 km for vertical occultations; 20-30 km for oblique occultations; 20-30 km for vertical cocultations; 20-30 km for oblique occultations; 20-30 km for vertical cocultations; 20-30 km for oblique occultations; 20-30 km for oblique occultations; 20-30 km for oblique oc		s Vertical resolution	
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	The suggested data screening removes outliers in data (see section	
	measurements from occultations of dim and cool stars (see section	
	"Dim and cool star problem").	
	Known problems and features	
	• Tangent line density standard deviation problem in IPF 6.01	
	The error bars of local and tangent line densities have been changed from relative errors in % (IPF version 5.01) to absolute values (IPF version 6) without changing the field format (unsigned integer field). Therefore, a coding equation is needed (please refer to the Product Specification). Values are expressed in absolute value (cm ⁻³) of the local density and correspond to 1 σ . The maximum value of the error estimate is set to	
	However, it is recommended to not use the standard deviation of the tangent line density since it is not possible to retrieve the correct value of the tangent line densities error , in spite of the use of the conversion formula.	
	Note, that the IPF 6.01 products contain correct values for the standard deviation of the local densities .	
	Dim and cool star problem	
	Ozone retrieval in the mesosphere and upper stratosphere (above 40 km) is based solely on transmission values at UV wavelengths. Weak and cool stars have low S/N-ratios in this part of the spectrum. It has been found (see Kyrölä et al., 2006) that measurements (during 2003) with stars of a visual magnitude >1.9 and temperature <7000 K are not providing reliable ozone retrievals above 40 km.	
	With the ageing of the GOMOS instrument the dark charge problem of GOMOS CCDs has increased and consequently the S/N-ratio has become lower for all stars further since 2003. By comparing statistical averages of ozone from different stars (with solar zenith angle larger than 105 deg.) 58 dim and cool stars were identified that do not support ozone retrievals above 40-50 km. Their star numbers according to the GOMOS catalogue are:	
	In period 2002-2004: 3, 13, 14, 17, 26, 37, 43, 48, 50, 51, 52, 53, 54, 61, 63, 65, 75, 84, 92, 93, 94, 102, 106, 114, 116, 117, 120, 126, 137, 138, 139, 141, 148, 151, 161, 162, 165, 167, 169, 170, 171, 178; in 2005-2008: In addition to the above list, stars 16, 40, 101, 113, 132, 135, 142, 143, 154, 155, 157;	
	in 2009-2012: In addition to lists for 2002-2008, stars 122, 134, 159, 164, 173.	
	The ozone retrievals from these stars also show negative bias at lower altitudes (the reason for this is presently under investigation). The "bad" (dim and cool) and "good" stars on the star temperature-magnitude plane are shown in Figure 1.	
L		1







NDACC/GAW network of ozonesondes. Top: median percentage relative difference; bottom: half 68% inter-percentile range of the relative difference (equivalent to standard deviation if the relative differences follow a normal distribution, but less sensitive to outliers than the standard deviation). Results are calculated in 5° latitude bins. GOMOS data were screened according to the recommendations in Section 2.5.1, and collocate within 500 km and up to 12h from the ozonesonde measurement.

Mid-Latitudes

The best agreement is obtained at mid-latitudes. In these regions, the median relative difference (GOMOS-validation instruments) is between -3% and +1% for altitudes between 18 km and 42 km (Van Gijsel, October 2012). Below 15 km, the comparisons to sondes indicate an increasingly positive bias.

Polar Regions

For the polar regions, a negative median difference (GOMOS-validation instruments) is found above ~15 km. Both lidar and ozonesonde comparisons indicate that the negative bias increases towards -(7-10)% at 25 km, and decreases again to 0% at 33 km. Above 33 km, the negative bias with respect to lidars increases again towards -11% at 40 km. In the UTLS and the troposphere the bias is positive, about +10% at 10 km and higher at lower altitudes. During Antarctic ozone hole conditions GOMOS ozone is 15% higher than that measured by sondes.

Tropics

In the tropics, increasing, positive (GOMOS-validation instruments) differences are found below the ozone maximum. But between 23km and 45 km, differences are again within a few percent around 0. The bias with respect to sondes is slightly positive, but remains below +3%. Lidar comparisons between 20 and 45 km show a bias between -2% and +1%.

Sensitivity to aerosol model

The concentration of ozone around the tropopause level is very sensitive to the aerosol model. For some cases, the ratio between ozone densities retrieved with two different aerosol models (2nd order and 3rd order polynomial for instance) is at least 1.5 below 20 km. Following intensive studies a second order polynomial was applied. This has to be considered for investigations regarding ozone at tropopause levels.

2.5.2 NO₂

The impact of the residual scintillation and of the choice of the aerosol model is reduced thanks to the DOAS inversion and regularisation.

Known problems and features

• Tangent line density standard deviation problem in IPF 6.01

It is recommended to not use the standard deviation of the tangent line density since it is **not** possible to retrieve the correct value of the **tangent line densities error**.

Note, that the IPF 6.01 products contain correct values for the standard deviation of the **local densities**.

Initial Validation/Verification Results

Validation studies have been based on balloon-borne stellar and moon occultation spectrometers (Renard et al., 2004) and satellite solar

occultation from HALOE (NO+NO₂ at sunset). The methodology is described in more detail in (Hauchecorne et al., 2005). The validity range is 20 km - 50 km. In this altitude range, the random error is estimated to be about 20 %. It depends mainly on the star magnitude. The systematic error, due to the uncertainties and temperature dependence in cross-sections, is no more than 10%. At other altitude ranges, data should be considered with caution, as the uncertainties and the temperature dependence in cross-sections may have a significant contribution to the error budget. 2.5.3 NO₃ The impact of the choice of the aerosol model is reduced thanks to the DOAS inversion. Known problems and features Tangent line density standard deviation problem in IPF 6.01 It is recommended to not use the standard deviation of the tangent line density since it is not possible to retrieve the correct value of the tangent line densities error. Note, that the IPF 6.01 products contain correct values for the standard deviation of the local densities. Initial Validation/Verification Results Validation studies have been based on balloon-borne stellar and moon occultation spectrometers (Renard et al., 2004). The methodology is described in more detail in (Hauchecorne et al., 2005). The validity range is 25 km-45 km. In this altitude range, the random error is estimated to be about 30%. It depends mainly on the star magnitude although cold stars are slightly better. The systematic error, due to the uncertainties and temperature dependence in cross-sections, is no more than 15%. At other altitude ranges, data should be considered with caution, as the uncertainties and the temperature dependence in cross-sections may have a significant contribution on the error budget. Retrieval is still noisy within the validity range. 2.5.4 Aerosols Aerosol spectral dependency is expressed as a 2nd order polynomial. The spectral parameters given for the aerosol extinction profile are those computed during the spectral inversion and correspond to the spectral dependence of the slant aerosol optical thickness, i.e. averaged along the line-of-sight. Aerosol spectral dependence is very sensitive to residual scintillation. Known problems and features As the current atmosphere is extremely transparent, the capacity to retrieve individual profiles is often at the limit of the instrument sensitivity. Very small particles are not spectrally distinguishable from Rayleigh

scattering.

Polar stratospheric clouds can be detected.

Data should be considered with caution above 35 km (low aerosol extinction) and below 10 km (small signal strength). Furthermore, due to the current implementation of the spectral law, only extinction profiles at the reference wavelength of 500 nm receive Tikhonov regularization during spatial inversion. This often results in rapidly oscillating profiles at other wavelengths.

It is therefore recommended to use only 500 nm aerosol extinction profiles.

• Local and Tangent line density standard deviation in IPF 6.01

Note, that the IPF 6.01 products the standard deviations of the **local and tangent line densities of the Aerosol** are still stored in percentage as it was for the IPF 5.01.

Initial Validation/Verification Results

No systematic validation has been performed for aerosol extinction spectral behaviour. However, preliminary studies indicate that GOMOS aerosol extinction profiles at 500 nm are in agreement within 50 % with data from the ACE/MAESTRO mission, in the altitude range from 10 km to 35 km.

2.5.5 H₂O

Known problems and features

• Tangent line density standard deviation problem in IPF 6.01

It is recommended to not use the standard deviation of the tangent line density since it is **not** possible to retrieve the correct value of the **tangent line densities error.**

Note, that the IPF 6.01 products contain correct values for the standard deviation of the **local densities**.

Initial Validation/Verification Results

H₂O densities are not retrieved above 50 km.

 H_2O profiles are of acceptable quality provided only by the 8 brightest stars in the near IR in dark limb (cold bright stars or very bright stars). The star numbers are: 1, 2, 3, 13, 14, 16, 26, 63. The retrieved densities from the other stars are of poor quality, related to too small SNR.

2.5.6 Neutral density

Neutral density is not retrieved with this version. The neutral density profile comes from ECMWF data completed by MSIS90 climatology above 1hPa pressure level. To avoid any confusion, products related to neutral density have been set to 0: the MDS for local density, the error bar, the vertical resolution, and the additional error. The covariance matrix terms related to neutral density have also been set to 0.

2.5.7 O₂

Known problems and features

second component of Star Therefore, the occultation during this period."	August 25 th to Octob S0034 appears in th is of this star can	per 28 th of each year, the upper background ban give degraded O ₂ profile	he id. es
Tangent line dens	sity standard deviat	ion problem in IPF 6.01	
It is recommended to not density since it is not poss line densities error. Note, that the IPF 6.01 pr deviation of the local dens	use the standard de sible to retrieve the co roducts contain corre sities.	eviation of the tangent lin orrect value of the tange ect values for the standa	ne ent ird
Initial Validation/Verification	n Results		
The relative difference be assessed from hundreds as (O ₂ GOMOS-O2 ECM computed assuming an a altitude. The mean bias is and 45 km, as indicated in Some profiles contain outl somewhat vertically during	etween O_2 by GOMO of dark limb occultat MWF)/ O_2 ECMWF. air mixing ratio of O_2 negative below 14 km Table 3. liers in the line densit g the vertical inversion	S and O_2 by ECMWF wa ions of star Sirius in 200 The O_2 from ECMWF of 0.20946 constant wi m, and positive between ty profiles, which propaga n.	as)3, is ith 16 tte
Altitude range (km)	Mean bias (%)	Dispersion (%)	
12 - 14 km	-3	6	
14 - 16 km	0	2	
16 - 30 km	+2.5	2	
30 - 35 km	+5	4	
30 - 35 km 35 - 40 km	+5 +10	4	
30 - 35 km 35 - 40 km 40 - 45 km	+5 +10 +15	4 10 20	
30 - 35 km 35 - 40 km 40 - 45 km 45 – 55 km	+5 +10 +15 0	4 10 20 30	
30 - 35 km 35 - 40 km 40 - 45 km 45 - 55 km Table 3. Mean bias and GOMOS and the ECMWF 2.5.8 OCIO	+5 +10 +15 0 dispersion of the r taken as a reference	4 10 20 30 elative difference betwee	en
30 - 35 km 35 - 40 km 40 - 45 km 45 - 55 km <i>Table 3. Mean bias and</i> <i>GOMOS and the ECMWF</i> 2.5.8 OCIO Not retrieved in the current	+5 +10 +15 0 dispersion of the r taken as a reference	4 10 20 30 elative difference betwee	en
30 - 35 km 35 - 40 km 40 - 45 km 45 - 55 km <i>Table 3. Mean bias and</i> <i>GOMOS and the ECMWF</i> 2.5.8 OCIO Not retrieved in the current	+5 +10 +15 0 dispersion of the r taken as a reference	4 10 20 30 elative difference between	en
30 - 35 km 35 - 40 km 40 - 45 km 45 - 55 km Table 3. Mean bias and GOMOS and the ECMWF 2.5.8 OCIO Not retrieved in the current 2.5.9 GAP	+5 +10 +15 0 dispersion of the r taken as a reference	4 10 20 30 elative difference between g.	en
30 - 35 km35 - 40 km40 - 45 km45 - 55 kmTable 3. Mean bias and GOMOS and the ECMWF2.5.8 OCIONot retrieved in the current2.5.9 GAPThe GAP (GOMOS Atmos) provided in the geolocation replaced by 0 and the error	+5 +10 +15 0 dispersion of the r taken as a reference eversion of processin pheric Profile) informa h ADS products. The r is set to 65535.	4 10 20 30 elative difference betweeners g. g. ation is not currently product values are	en
30 - 35 km 35 - 40 km 40 - 45 km 45 - 55 km <i>Table 3. Mean bias and</i> <i>GOMOS and the ECMWF</i> 2.5.8 OCIO Not retrieved in the current 2.5.9 GAP The GAP (GOMOS Atmosprovided in the geolocation replaced by 0 and the error 2.5.10 HRTP	+5 +10 +15 0 dispersion of the r taken as a reference eversion of processin pheric Profile) informa ADS products. The r is set to 65535.	4 10 20 30 elative difference betweeners g. g. ation is not currently product values are	en

depends on the obliquity of the occultation. The best precision is achieved for vertical occultations. It is estimated to be about 1-2 K at altitudes between 18 km and 30 km, for vertical occultations and for close to vertical occultations. Known problems and features In case of a bad correlation between the photometer signals at a given altitude, the weighted mean of the time delay estimate computed using ECMWF air density and the experimental value are used. The weights are inversely proportional to uncertainties in experimental and ECMWF-model time delays. The HRTP values at this altitude are flagged (the corresponding error estimates for density and temperature are set to 6500%). Initial Validation/Verification Results First validation results for GOMOS version 6.01 show a very good agreement in comparison to sonde data up to 30 km (not significantly different from 0). In comparison to lidar data, the highest altitudes show an underestimation (increasing with altitude) of the temperature by GOMOS. This could be related to the upper-limit initialization used for GOMOS HRTP retrievals. 2.6 References **GOMOS Algorithm theoretical Basis Document** E. Kvrölä, L. Blanot, J. Tamminen, V. Sofieva, J. L. Bertaux, A. Hauchecorne, F. Dalaudier, D. Fussen, F. Vanhellemont, O. Fanton d'Andon, G. Barrot ENVISAT-GOMOS Level 1b to 2 Processing, ATBD (GOM-FMI-TN-040, issue 3.0), December 2012 Validation of GOMOS version 6.01 ozone profiles using ground-based lidar observations Anne van Gijsel, October 2012 http://earth.eo.esa.int/pcs/envisat/gomos/documentation First climatology of polar mesospheric clouds from **GOMOS/ENVISAT** stellar occultation instrument K. Pérot, A. Hauchecorne, F. Montmessin, J.-L. Bertaux, L. Blanot, F. Dalaudier, D. Fussen, and E. Kyrölä Atmos. Chem. Phys., 10, 2723-2735, 2010 GOMOS O₃, NO₂, and NO₃ observations in 2002–2008 E. Kyrölä, J. Tamminen, V. Sofieva, J. L. Bertaux, A. Hauchecorne, F. Dalaudier, D. Fussen, F. Vanhellemont, O. Fanton d'Andon, G. Barrot, M. Guirlet, T. Fehr, and L. Saavedra de Miguel Atmos. Chem. Phys., 10, 7723-7738, 2010 Optical extinction by upper tropospheric/stratospheric aerosols and clouds: GOMOS observations for the period 2002-2008 F. Vanhellemont, D. Fussen, N. Mateshvili, C. Tétard, C. Bingen, E. Dekemper, N. Loodts, E. Kvrölä, V. Sofieva, J. Tamminen, A. Hauchecorne, J.-L. Bertaux, F. Dalaudier, L. Blanot, O. Fanton d'Andon, G. Barrot, M. Guirlet, T. Fehr, and L. Saavedra de Miguel Atmos. Chem. Phys., 10, 7997-8009, 2010

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	2.7 Acronyms and Abbreviations	
	ACE/MAESTROAtmospheric Chemistry Experiment/Measurement of Aerosol Extinction in the Stratosphere and Troposphere Retrieved by OccultationADSAnnotation Data SetCCDCharge Coupled DeviceDOASDifferential Optical Absorption SpectroscopyECMWFEuropean Centre for Medium-term Weather ForecastGAWWMO's Global Atmosphere WatchGOMOSGlobal Ozone Monitoring by Occultation of StarsHALOEHALogen Occultation ExperimentHRTPHigh Resolution Temperature ProfileIPFInstrument Processing FacilityLUTLook-Up TableMDSMeasurement Data SetMPHMain Product HeaderMSIS90Mass-Spectrometer-Incoherent-Scatter-1990 atmosphere modelNDACCNetwork for the Detection of Atmospheric Composition ChangePCDProduct Confidence DataPRNUPixel Response Non Uniformity SATUSATUStar Acquisition and Tracking Unit	
	2.8 Acknowledgement	
	Input for this product quality readme file came from the GOMOS Quality Working Group, independent validation teams, and the scientific community which is responsible for the processor development.	
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