

Comparison of two Surface Reflectivity Models and Validation with Measurements

Outline:

1. Introduction
2. Models
3. Experimental Data
4. Results
5. Conclusions & Summary

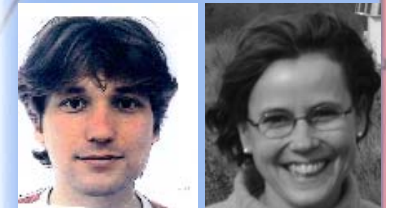
M. Schwank & I. Völksch



J.P. Wigneron & J.H. Kerr



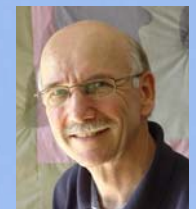
A. Mialon & P. de Rosnay



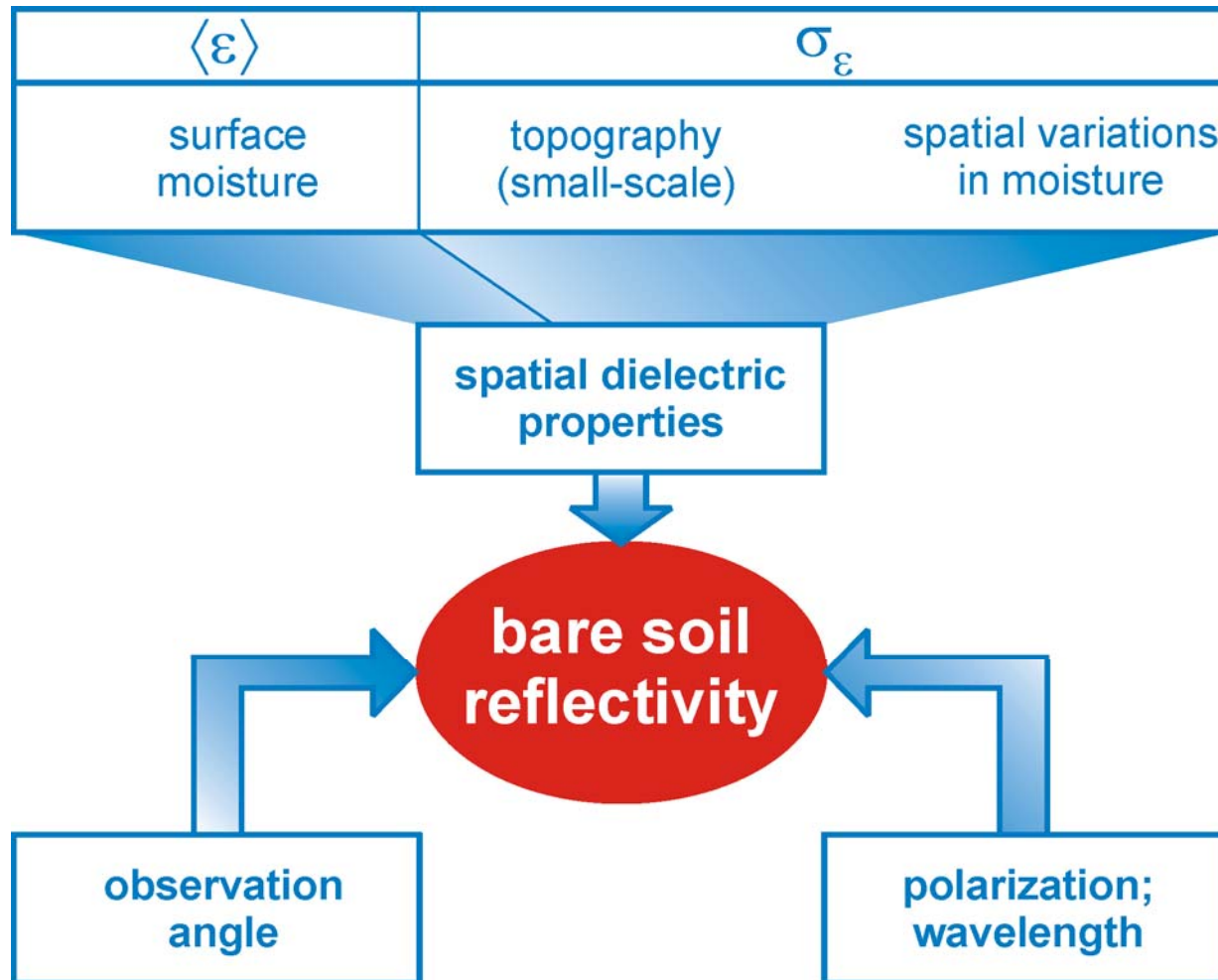
C. Mätzler



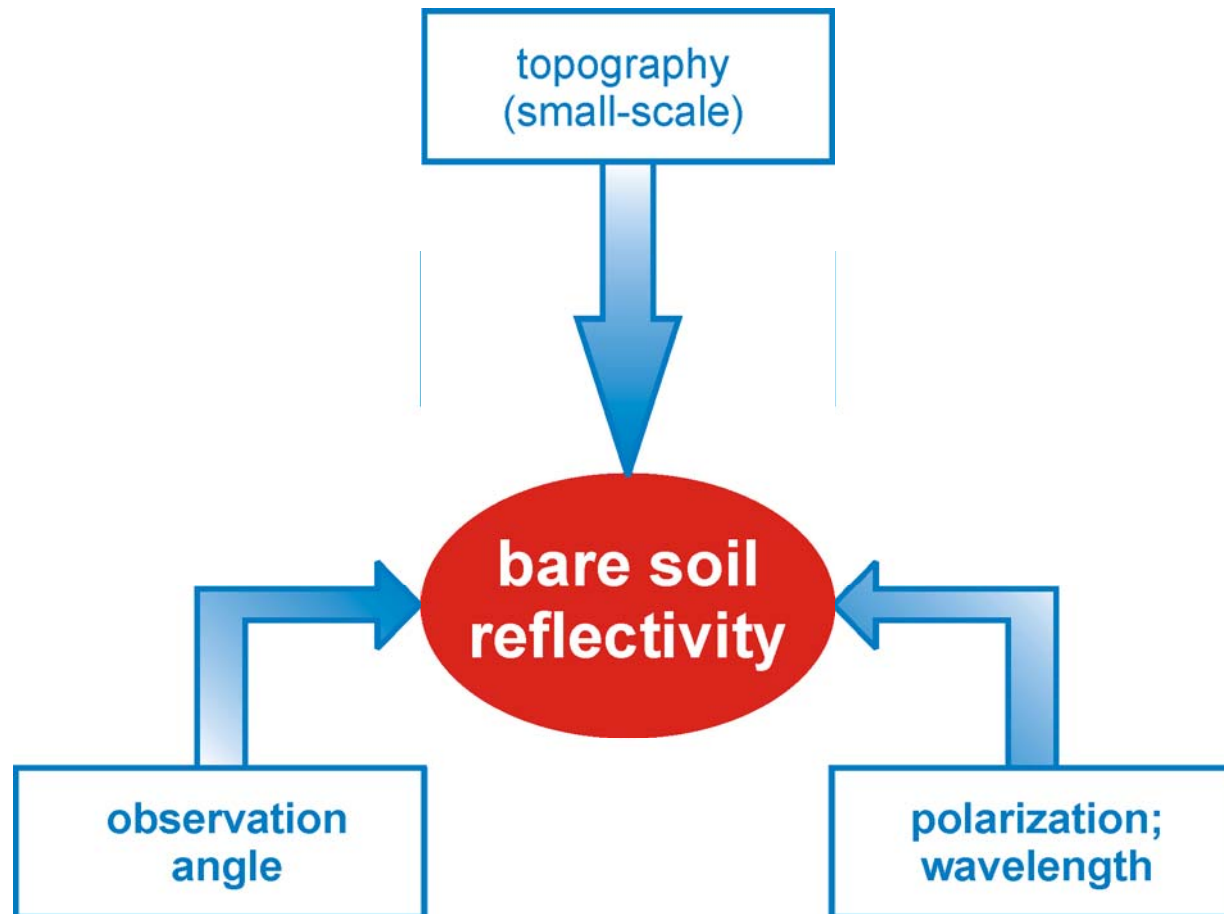
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L-band reflectivity of a bare soil is affected by...



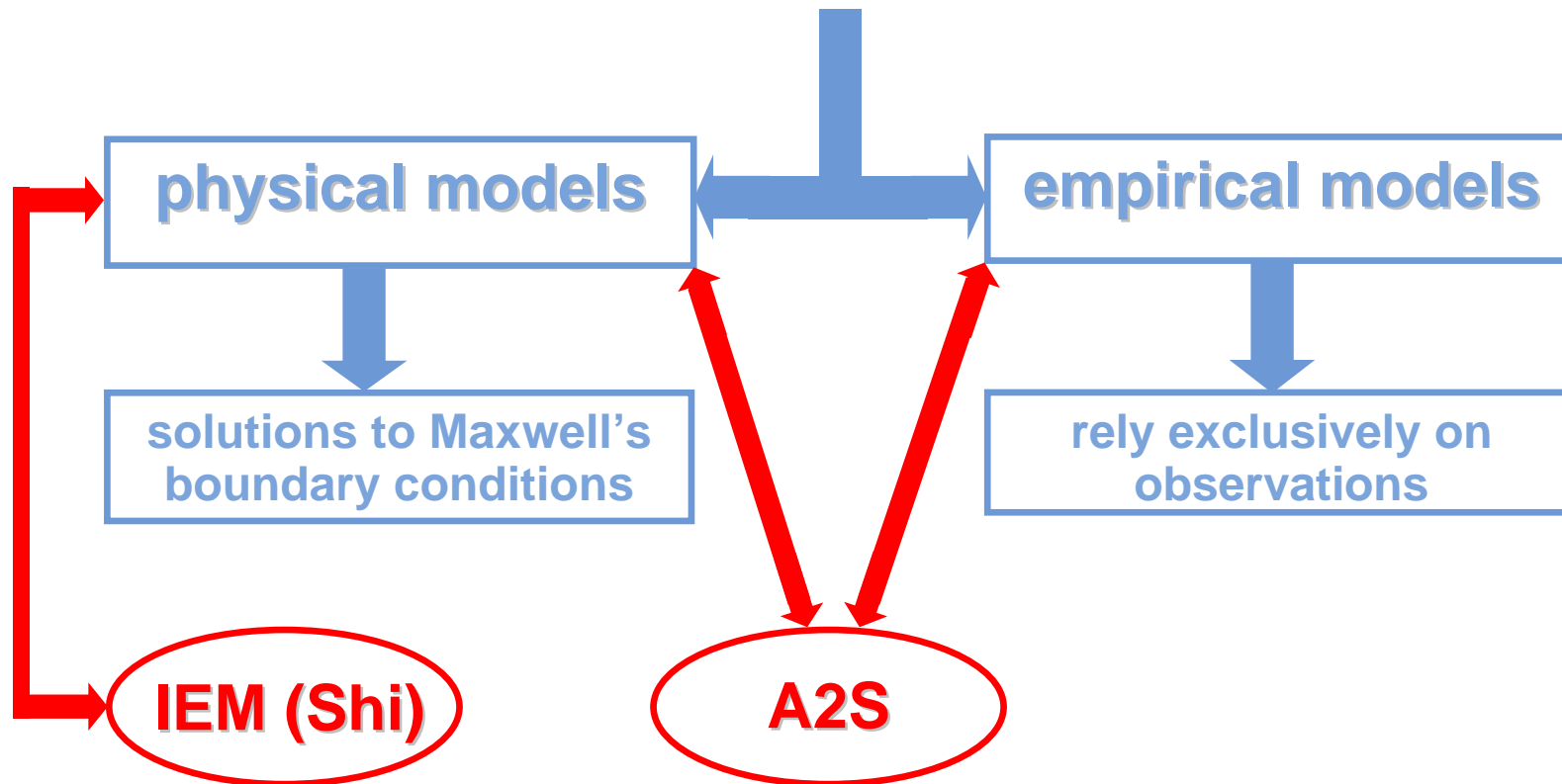
L-band reflectivity of a bare soil is affected by...



Surface Reflectivity Models



2. Models



the two models investigated are:

IEM model:

Shi's parametrization
(fast model)

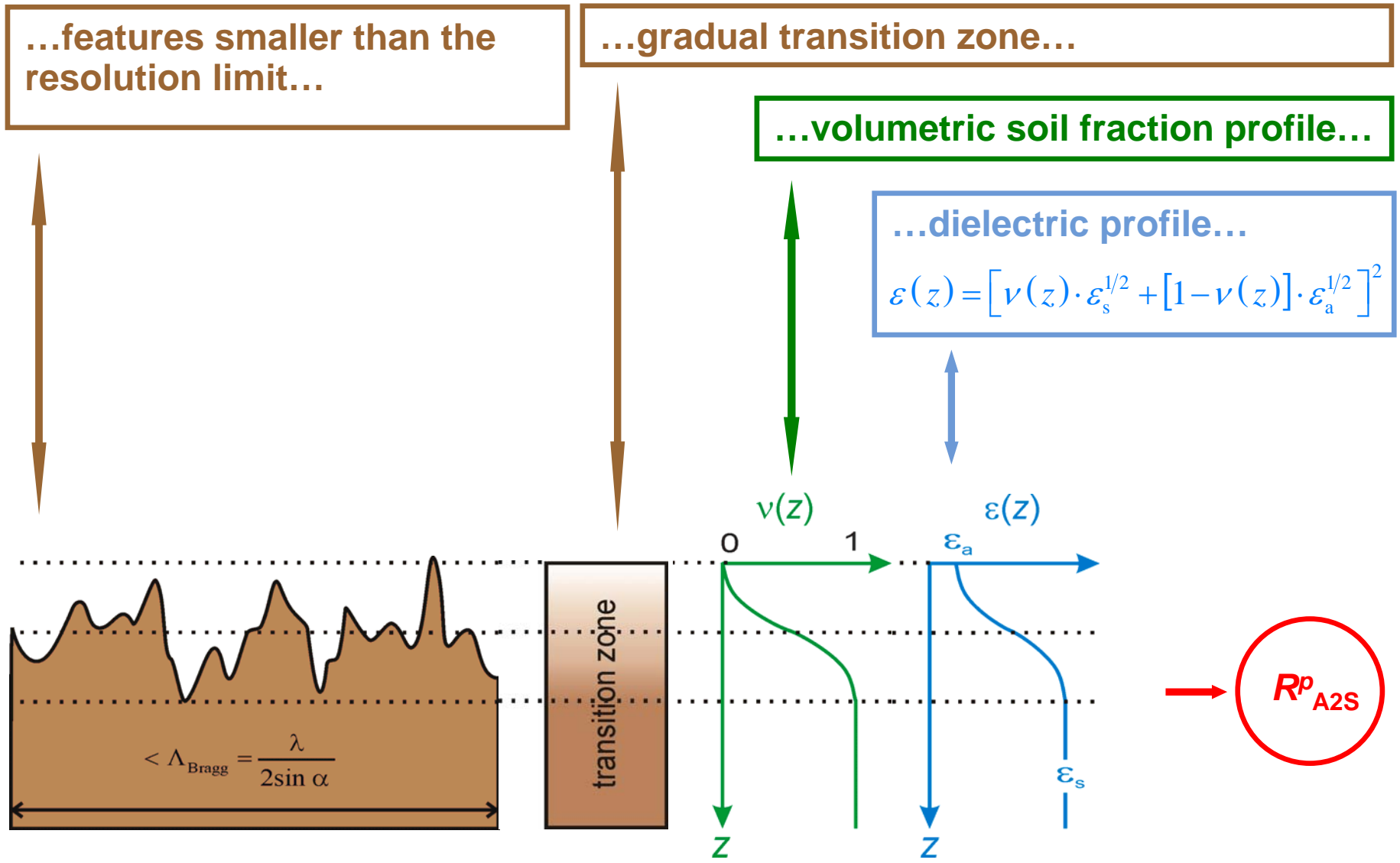
A2S transition model:

physical: Impedance matching.
empirical: Bragg limit.
Refractive mixing.

Concept of the A2S model



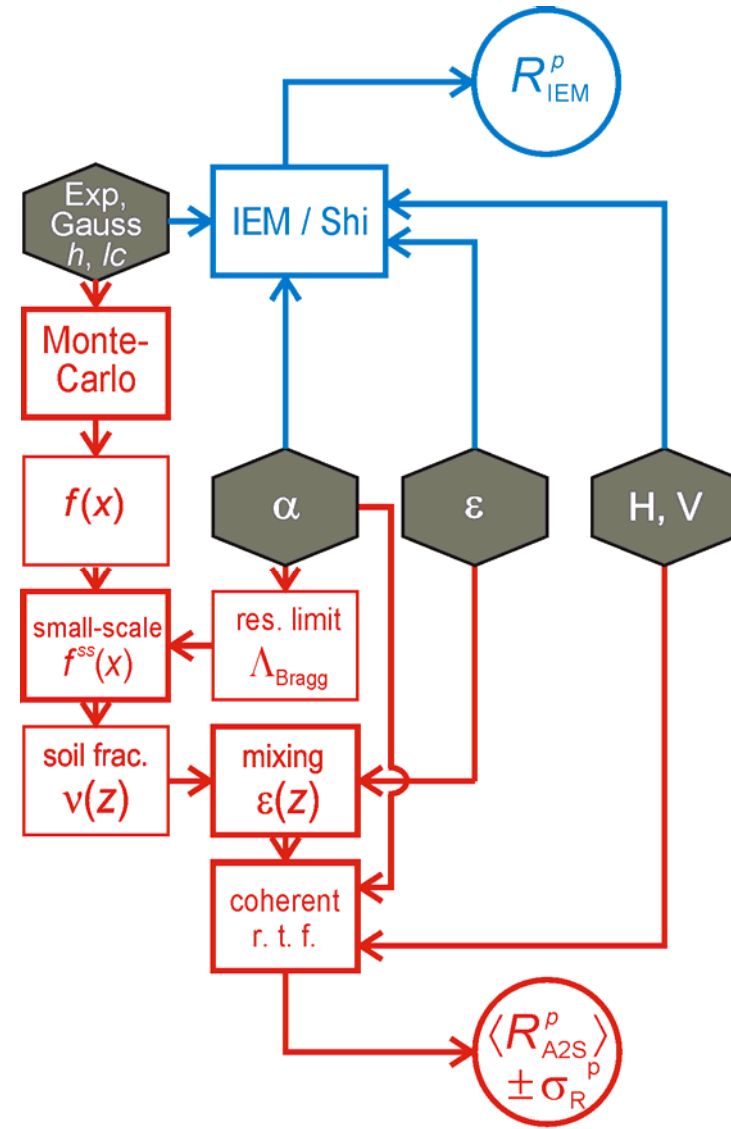
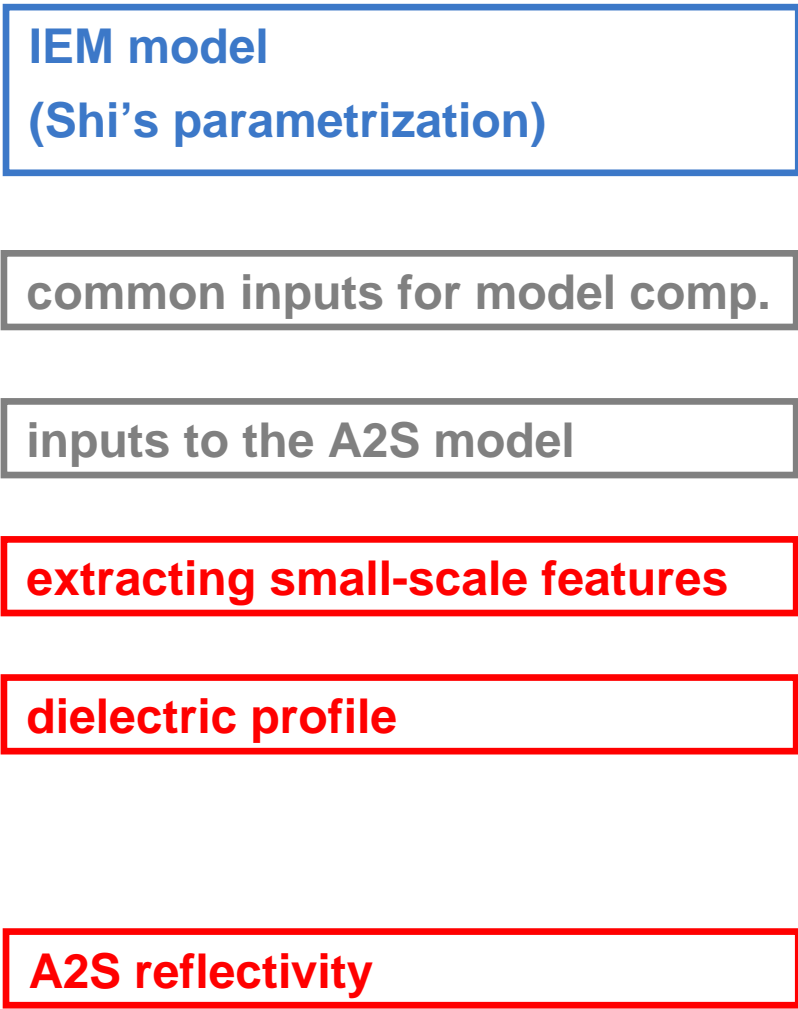
2. Models



A2S- and IEM- Modeled Reflectivities



2. Models



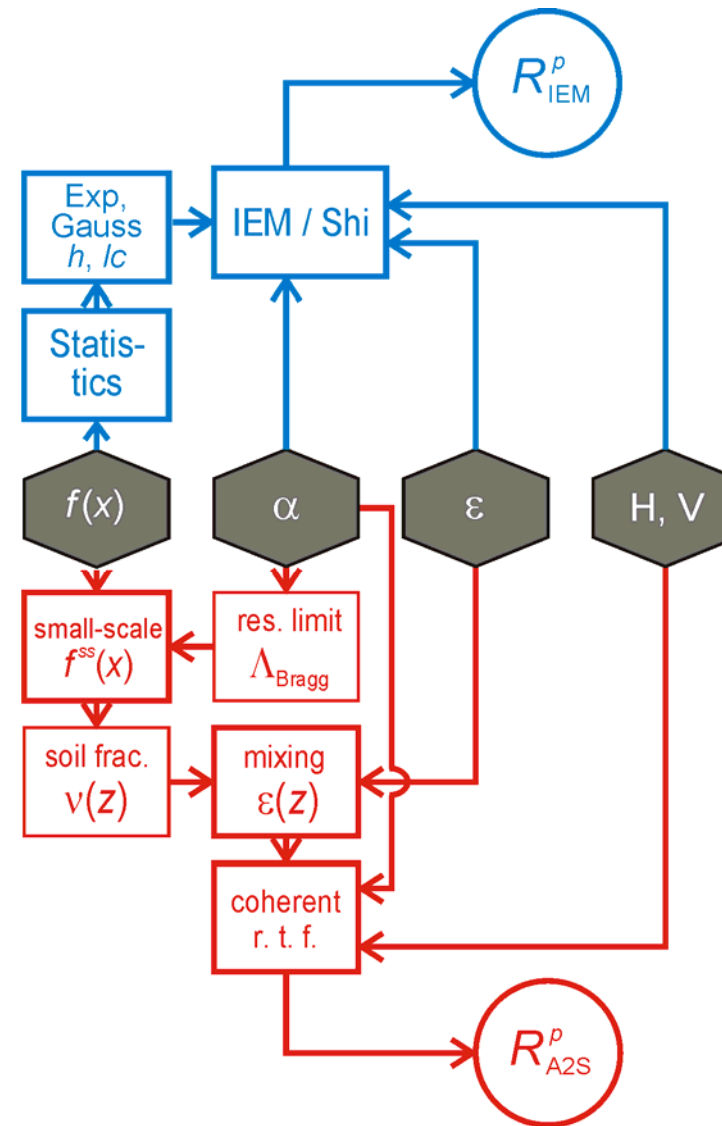
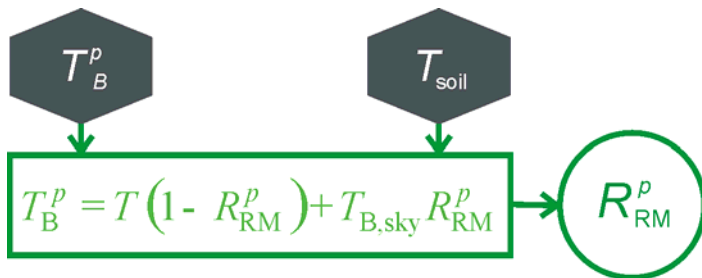
Modeled and Measured Reflectivities



2. Models

common inputs for evaluating the models using measured topographies

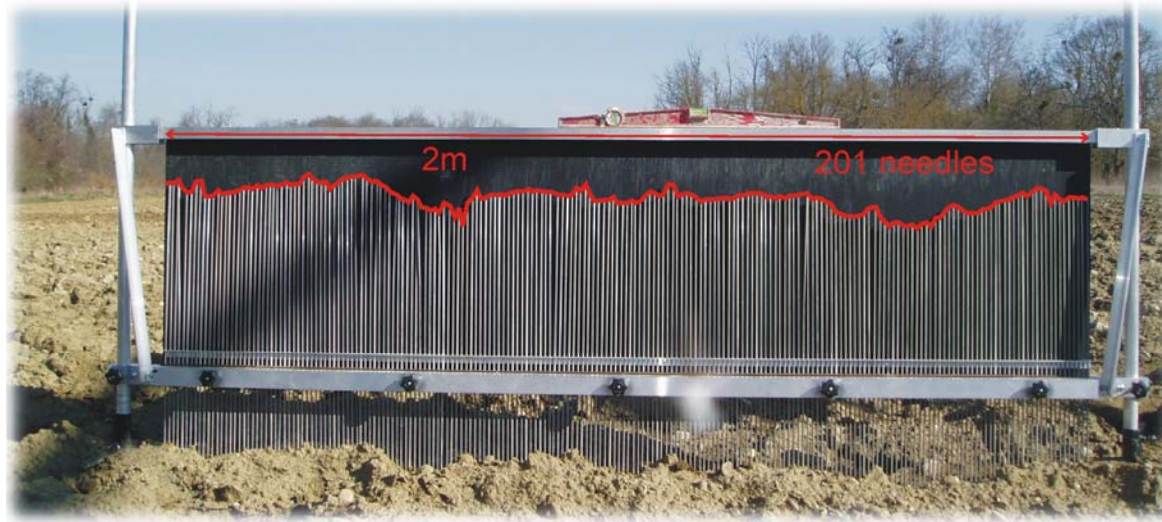
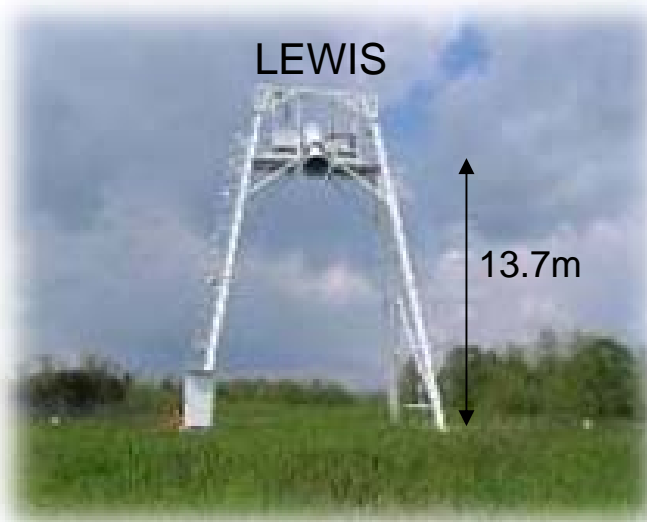
Reflectivities derived from Radiometer measurements



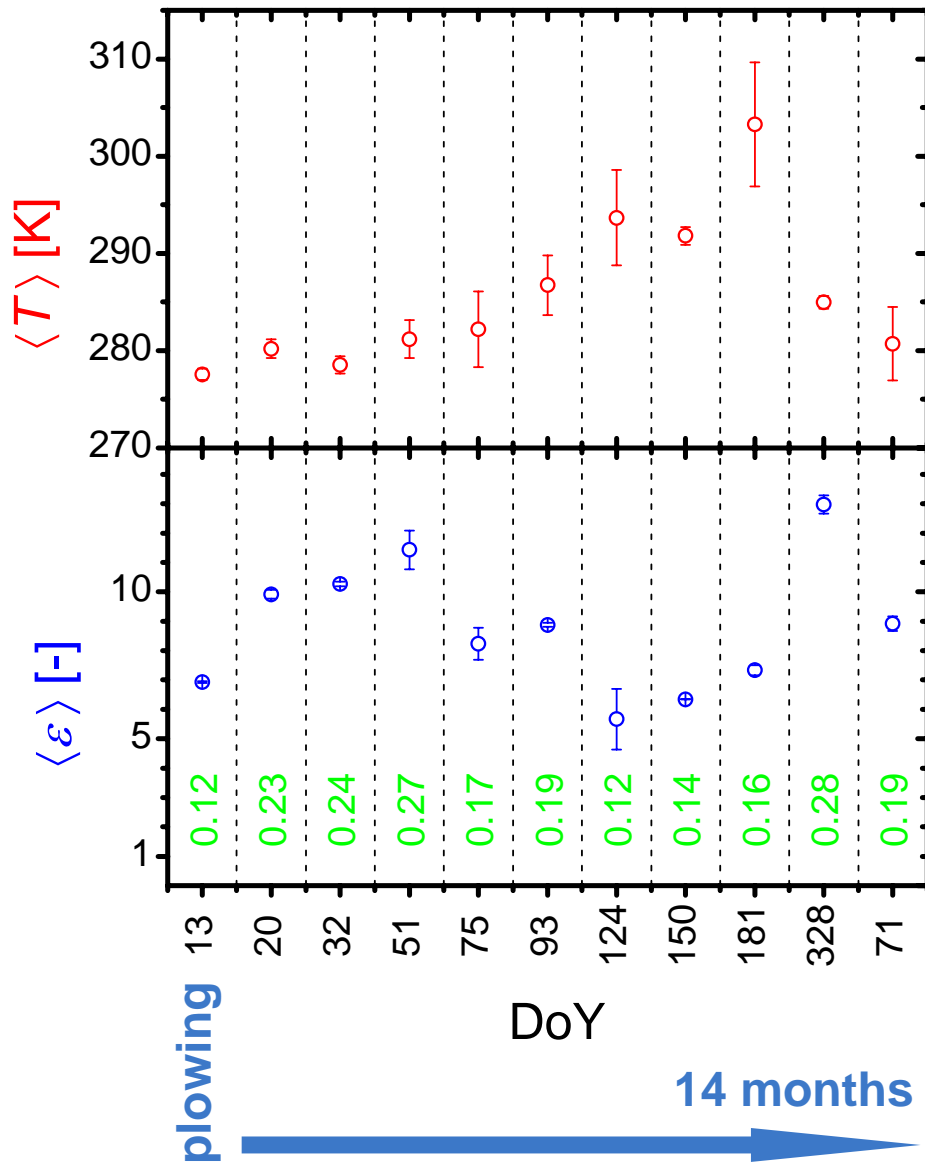
SMOSREX Data (near Toulouse, France)



- **11 days:**
DoY = 13, 20, 32, 51, 75, 93, 124, 150, 181, 328 in 2006 and
DoY = 71 in 2007
- T_B^H, T_B^V over bare soil:
Lewis; $\alpha = 20^\circ - 60^\circ$; $\Delta t = 3$ h
- In-situ T and ε :
thermistors & Theta-probes
- topography profiles:
needle board; 11-16 per day



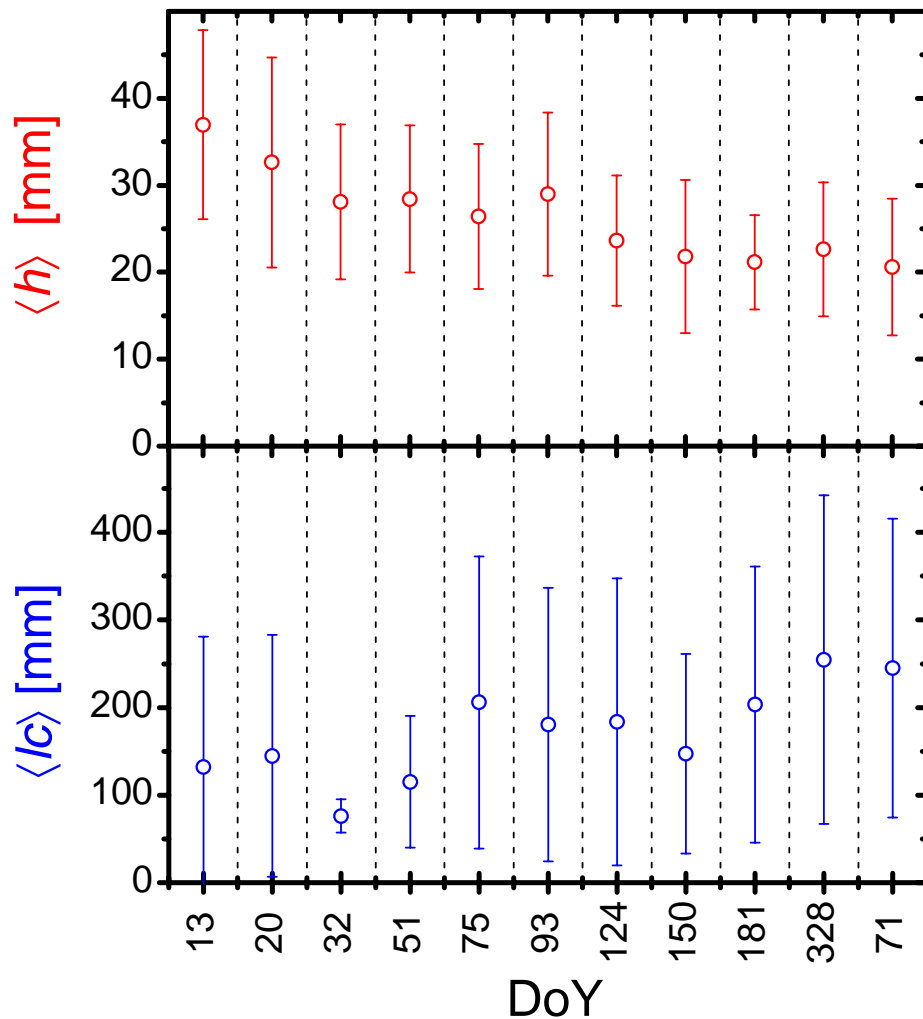
Daily Mean Soil Temperature and Permittivity



- Several probes installed within the topmost 6 cm.
- One measurement per 30 minutes

- Volumetric moisture [m³m⁻³] (Topp et al., 1980)

Daily Mean RMS-height and Correlation length



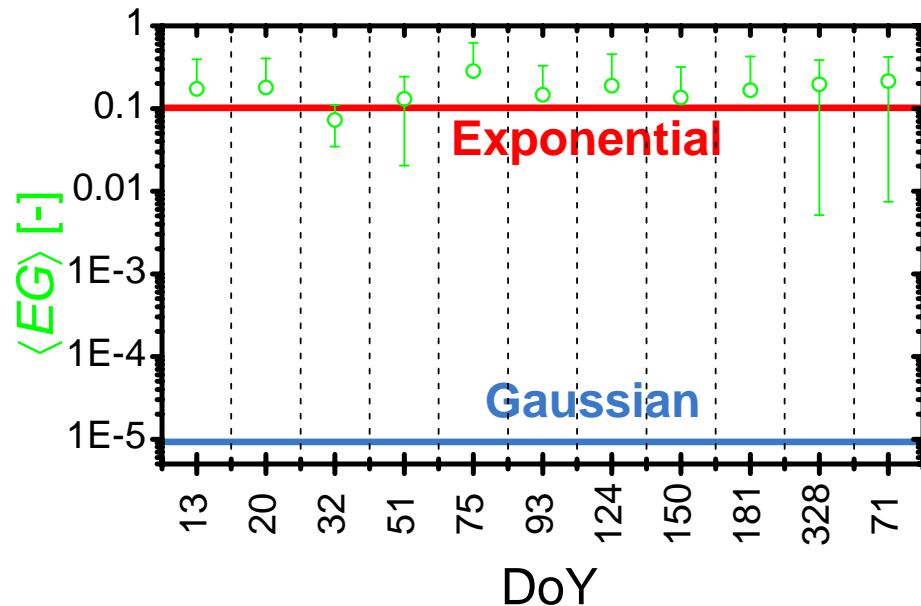
▪ Surface RMS-height decreases from $\langle h \rangle \approx 39$ mm to $\langle h \rangle \approx 20$ mm

▪ Surface correlation length $\langle l_c \rangle$ does not increase significantly with time.

Standard deviations σ_h and σ_{l_c} do not decrease with time.



Daily Mean Spectral Indicator



$$EG = \frac{\text{spatial wavelenghts} \leq \text{correlation length}}{\text{total power spectrum}}$$

$$EG = \frac{\sum_{k \geq 2\pi/lc} |c_k|^2}{\sum_k |c_k|^2}$$

$$\langle EG \rangle \approx 0.17$$

$$EG_S \equiv \frac{\int_{k_{lc}}^{\infty} W_S(k) dk}{\int_0^{\infty} W_S(k) dk}$$

spectral power densities:

$$W_G(k) = \frac{h^2 lc}{2\sqrt{\pi}} \exp\left(-\frac{k^2 \cdot lc^2}{4}\right)$$

$$W_E(k) = \frac{h^2 lc}{\pi(1+k^2 lc^2)}$$

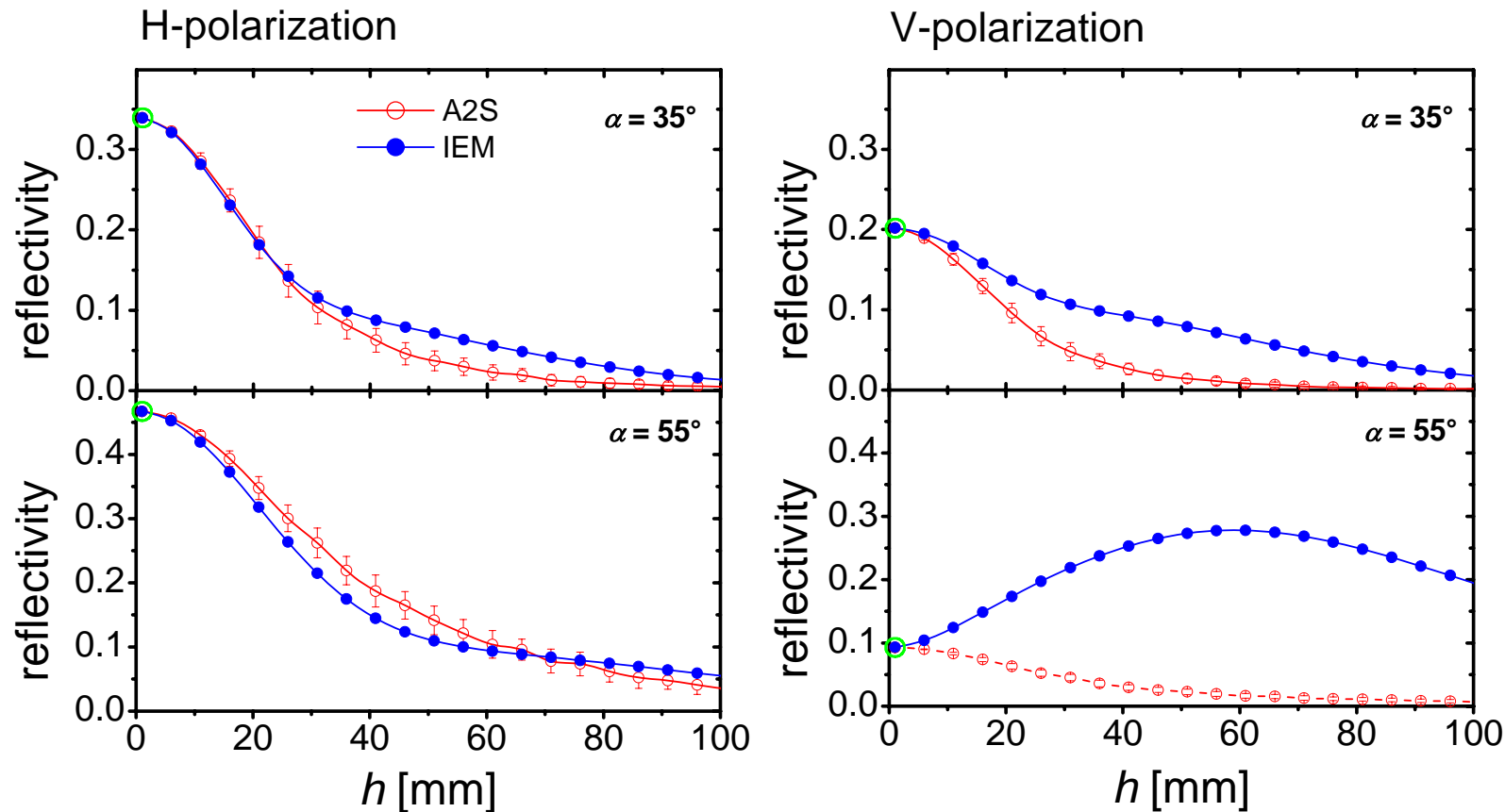
⇒ The correlation function of a naturally weathered bare soil surface is Exponential



Comparison between A2S- and IEM Model



4. Results

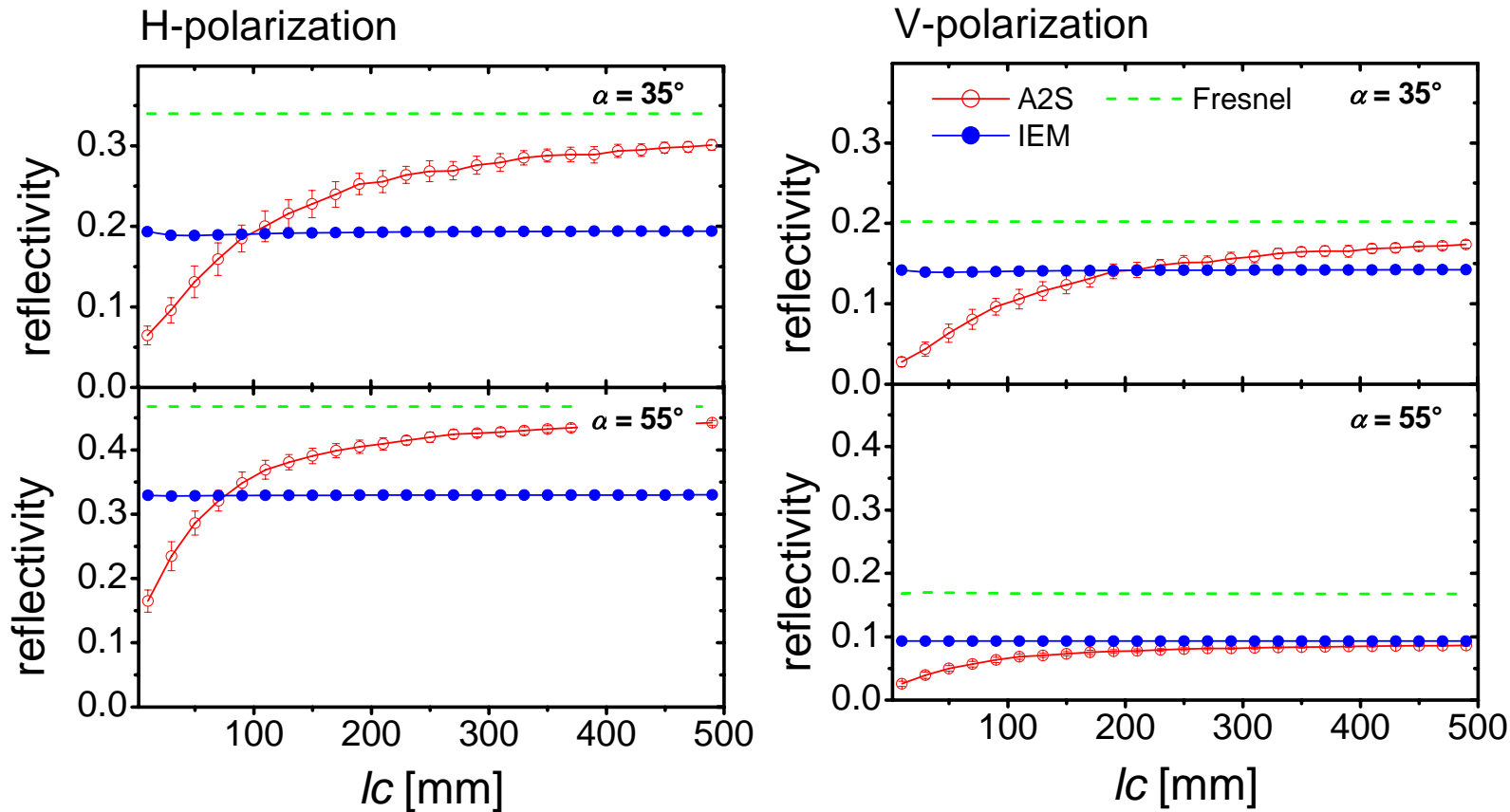


- Variation of h ; $\varepsilon = 10$
- Both models reproduce the **Fresnel reflectivity** for $h \rightarrow 0$
- Similar results for $p = H$ (reflectivity decreases with h , $\propto \text{Exp}(-h^2)$)
- IEM reflectivity: $p = V$, $\alpha = 55^\circ$ increase with h ;
polarization cross-talk, not observed

Comparison between A2S- and IEM Model



4. Results

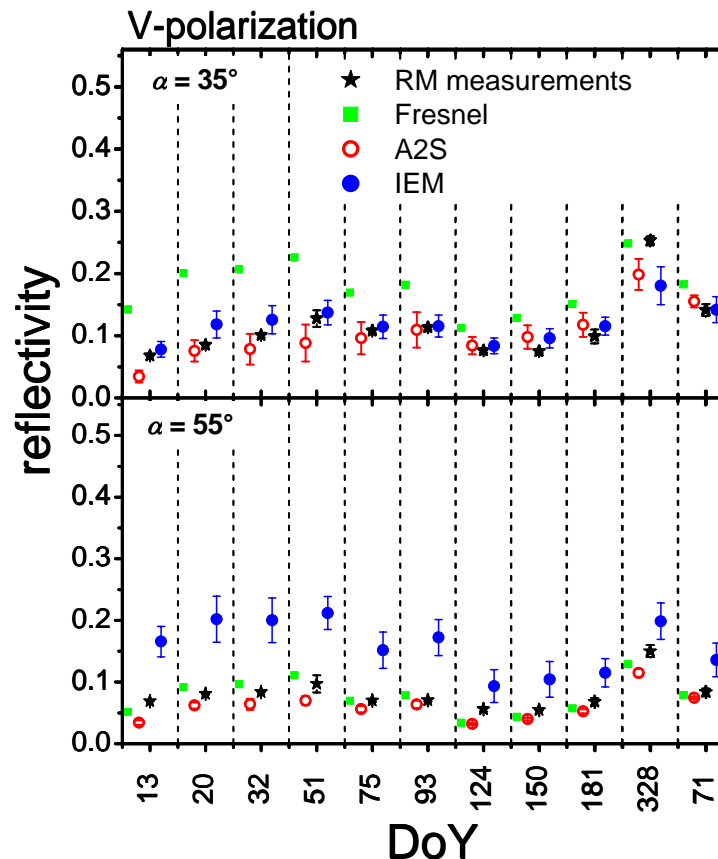
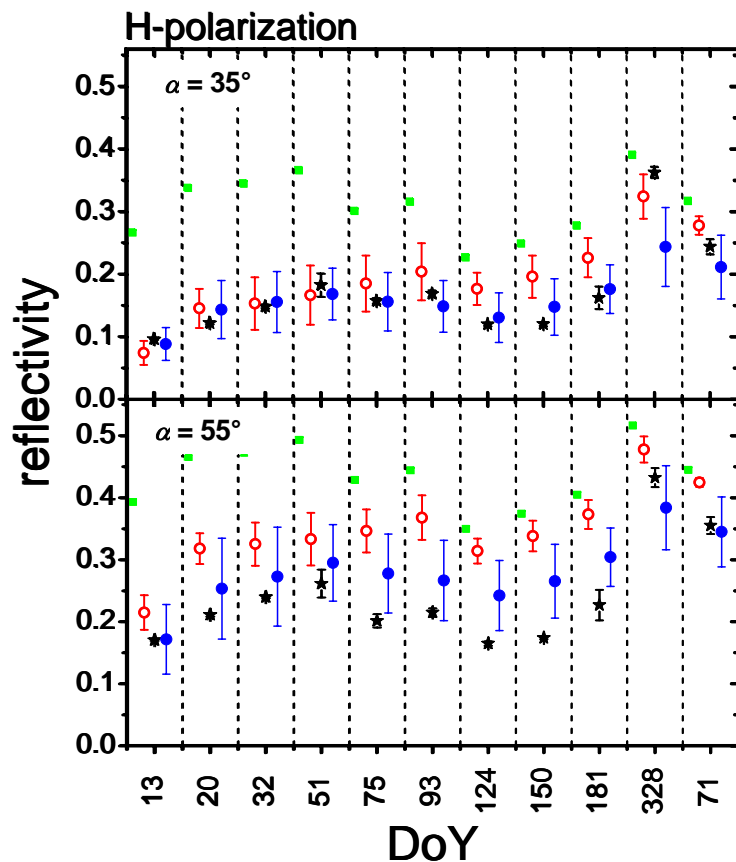


- Variation of lc ; $\varepsilon = 10$
- IEM reflectivities are not much sensitive with respect to lc
- A2S reflectivities increase with lc to approach values smaller than the Fresnel reflectivities.
- $lc \uparrow \Rightarrow$ proportion of the spatial small-scale features $\downarrow \Rightarrow R_{A2S} \uparrow$

Comparison with Measurements



4. Results



α [°]	p	OK_F	δ_F [%]
35	H	0	97
35	V	1	68
55	H	0	92
55	V	3	16

$$\delta_F = \frac{100}{n_{\text{DoY}}} \sum_{i=1}^{n_{\text{DoY}}} \frac{|\langle R_F^p \rangle_i - \langle R_{\text{RM}}^p \rangle_i|}{\langle R_{\text{RM}}^p \rangle_i}$$



- The impact of roughness on the soil reflectivity was analyzed (IEM-model \Leftrightarrow A2S-model \Leftrightarrow experimental data)
- Distinct differences between the models
- Depending on the polarization and the observation angle, either the IEM-or the A2S-model performs better.
- Polarization cross-talk was overestimated by the IEM-model
- Detailed information on the soil topography might be not sufficient for accurate predictions of soil reflectivity. The impact of dielectric heterogeneities and anisotropies of the bulk soil in the topmost centimeters can be dominant.

Thank you for your Attention

Mike Schwank