

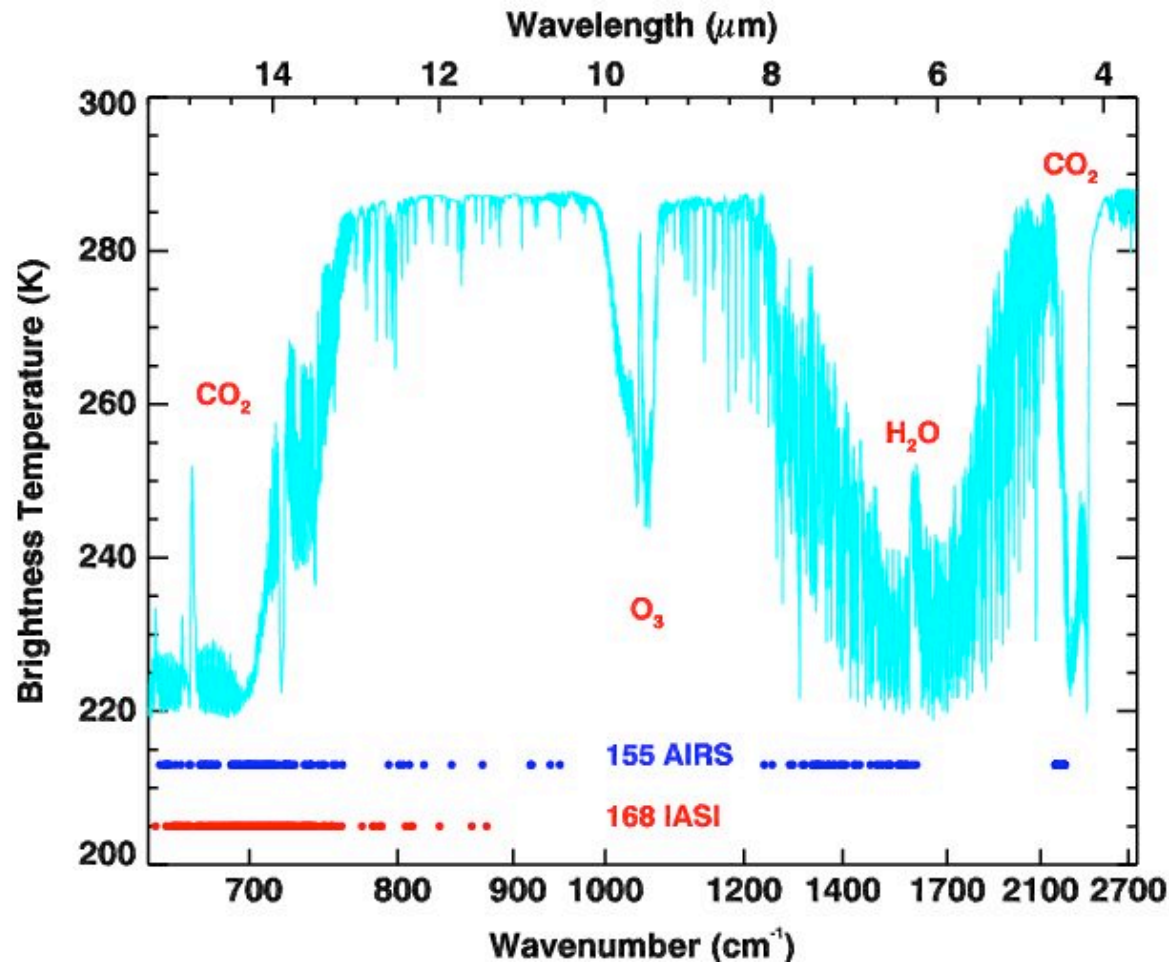
# **New Technologies, Applications and Reanalysis**

## **lecture 3**

**Tony McNally**  
**ECMWF**

# **Advanced IR sounders**

# Advanced infrared sounders: AIRS and IASI



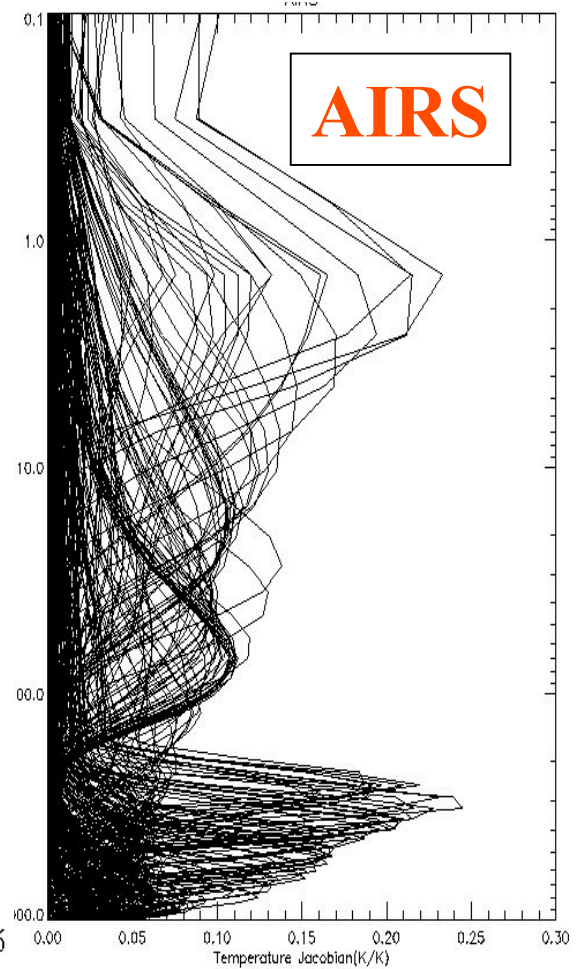
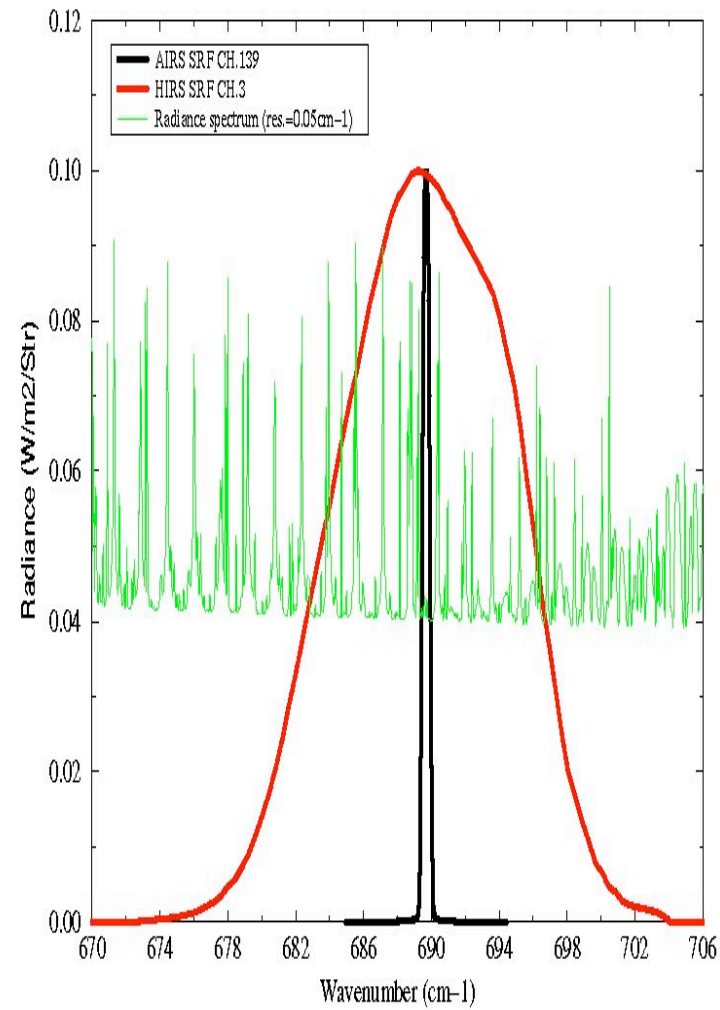
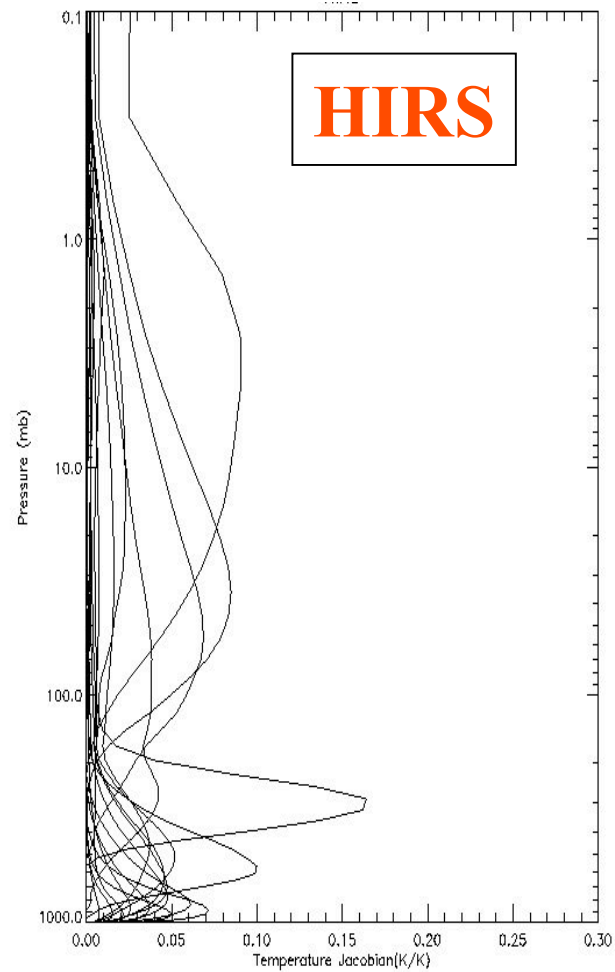
## AIRS

- | Operational at ECMWF since October 2003.
- | 324 channels received in NRT.
- | One FOV in nine used.
- | Up to 155 channels may be assimilated (CO<sub>2</sub> and H<sub>2</sub>O bands).

## IASI

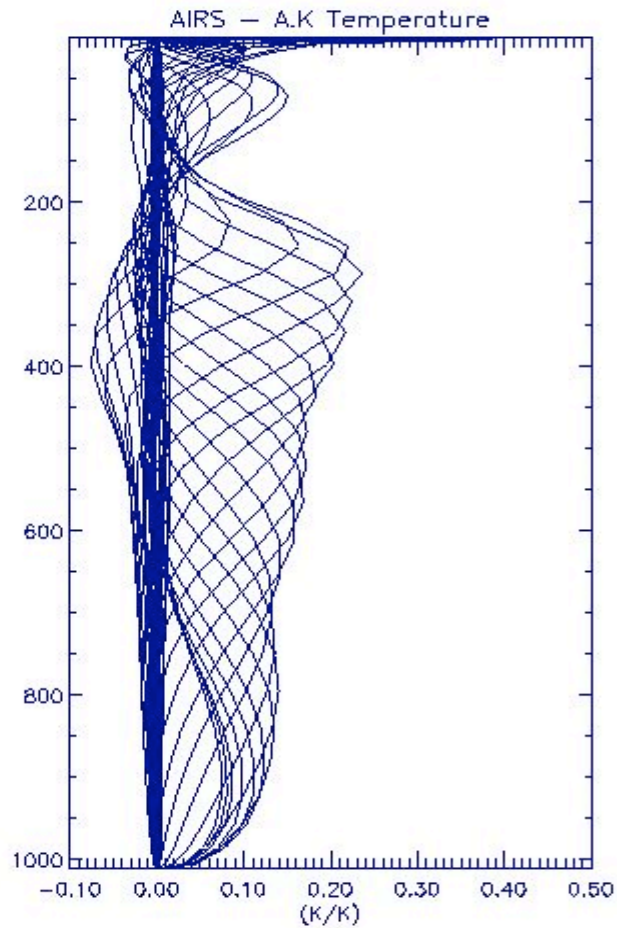
- | Operational at ECMWF since June 2007.
- | 8461 channels received in NRT.
- | All FOVS received; only 1-in-4 used.
- | 366 Channels routinely monitored.
- | Up to 168 channels may be assimilated (CO<sub>2</sub> band only).

# Advanced infrared sounders: AIRS and IASI

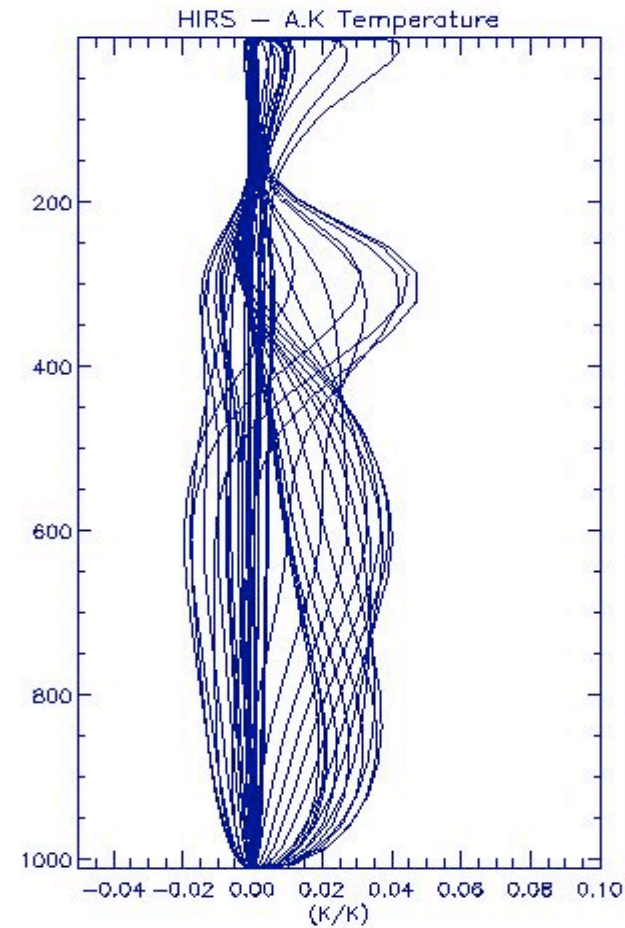


# Better measure of improved resolution is provided by the averaging kernels

**AIRS**



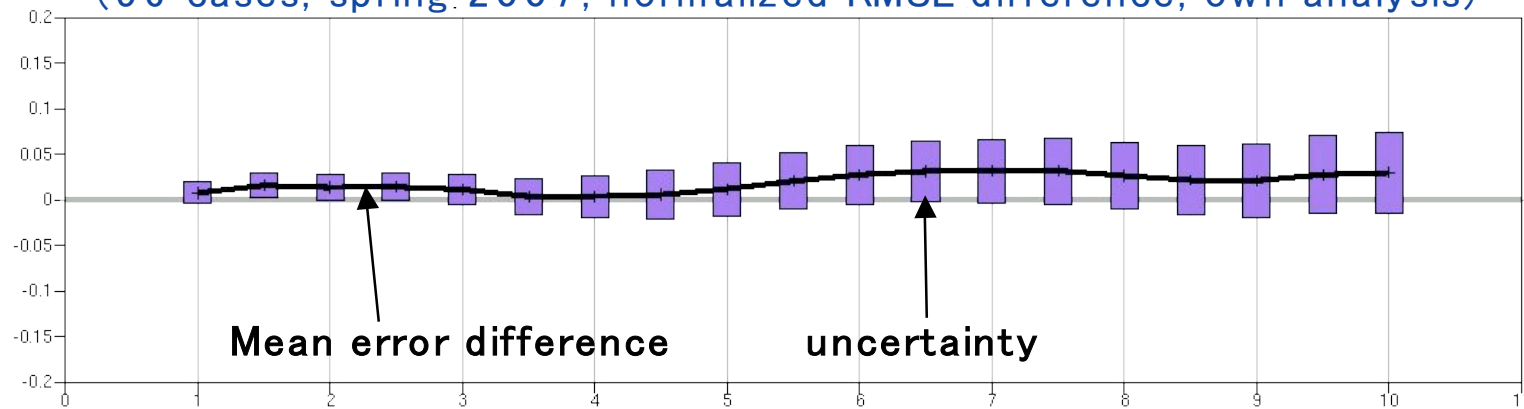
**HIRS**



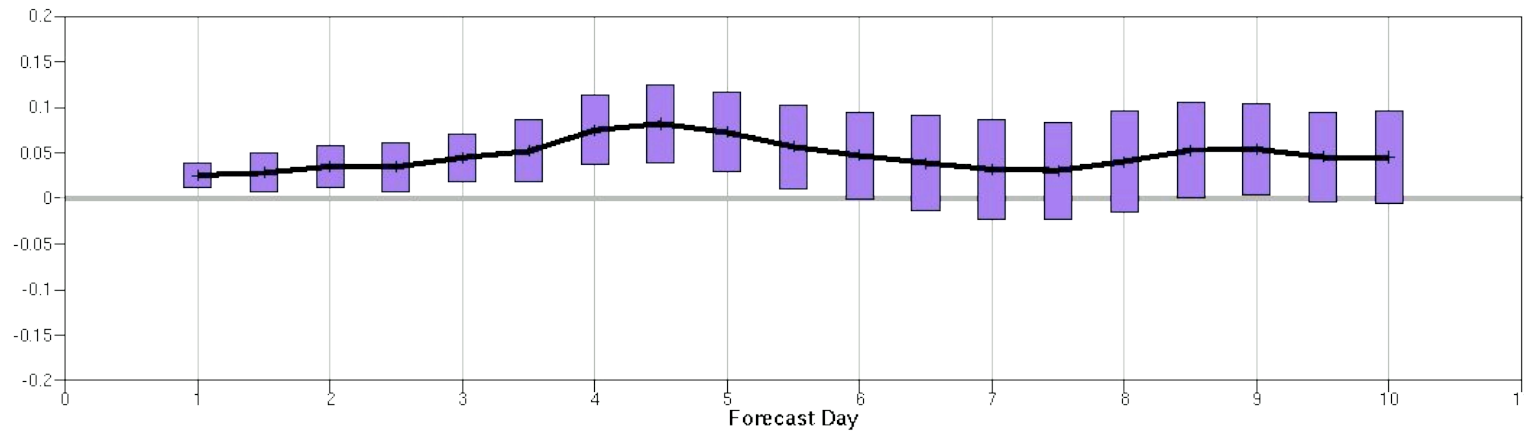
# IASI forecast impact

500 hPa geopotential anomaly correlation  
(56 cases, spring 2007, normalized RMSE difference, own analysis)

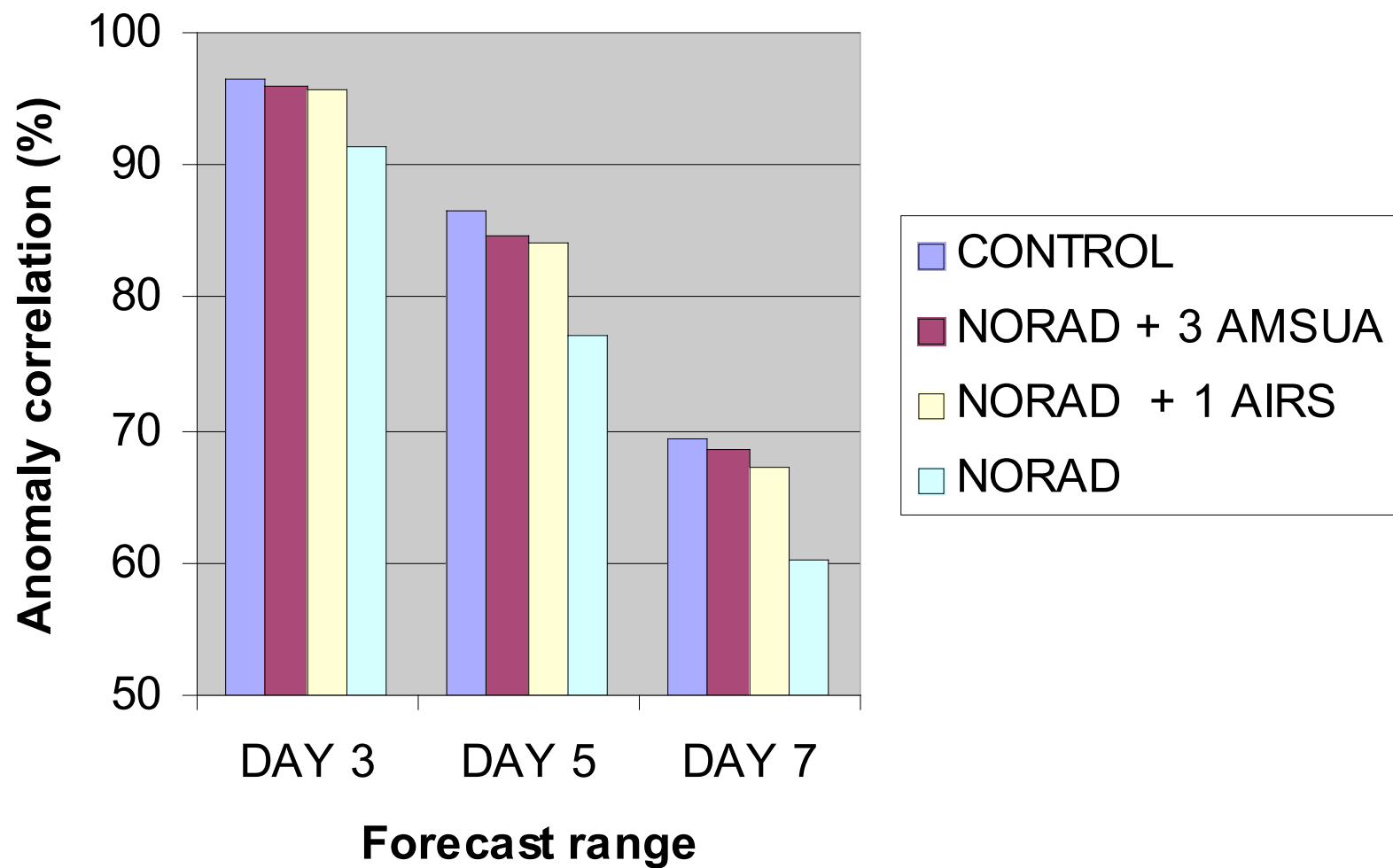
NH



SH



## Southern Hemisphere scores



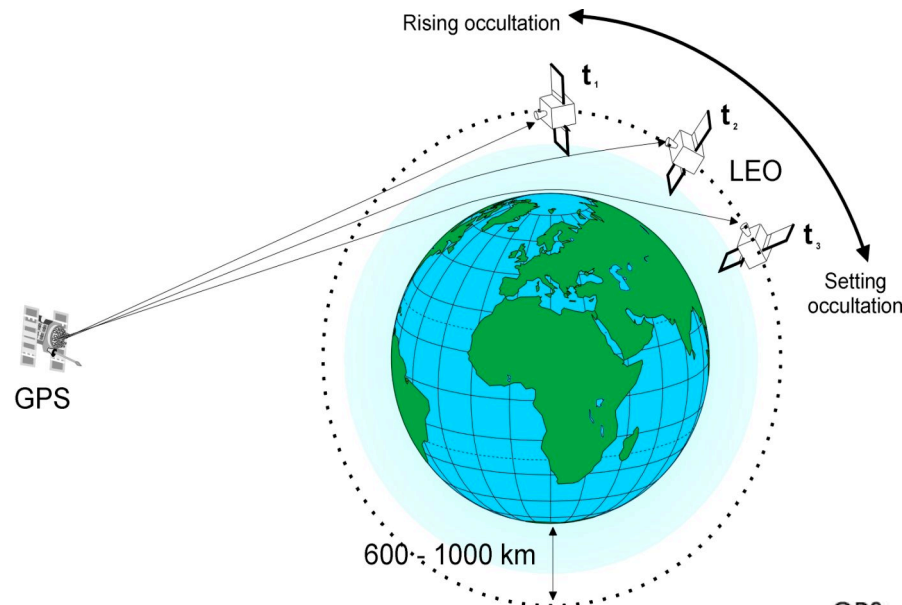
# **GPS Radio Occultation**



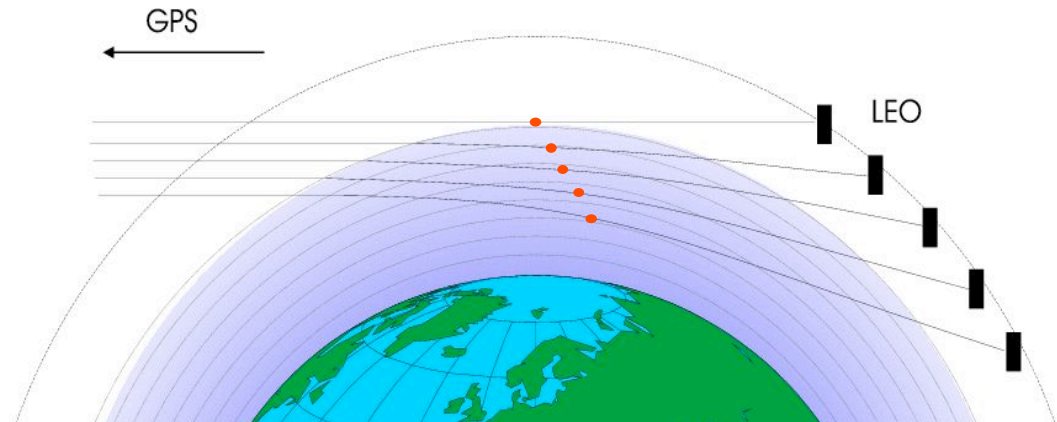
# Radio occultation geometry

- CHAMP, COSMIC, METOP-GRAS

- The impact of the atmosphere on the signal propagation depends on the refractivity => the vertical profile of the refractivity (and further down temperature, humidity and pressure) at the location of the ray perigee can be inverted from the observation

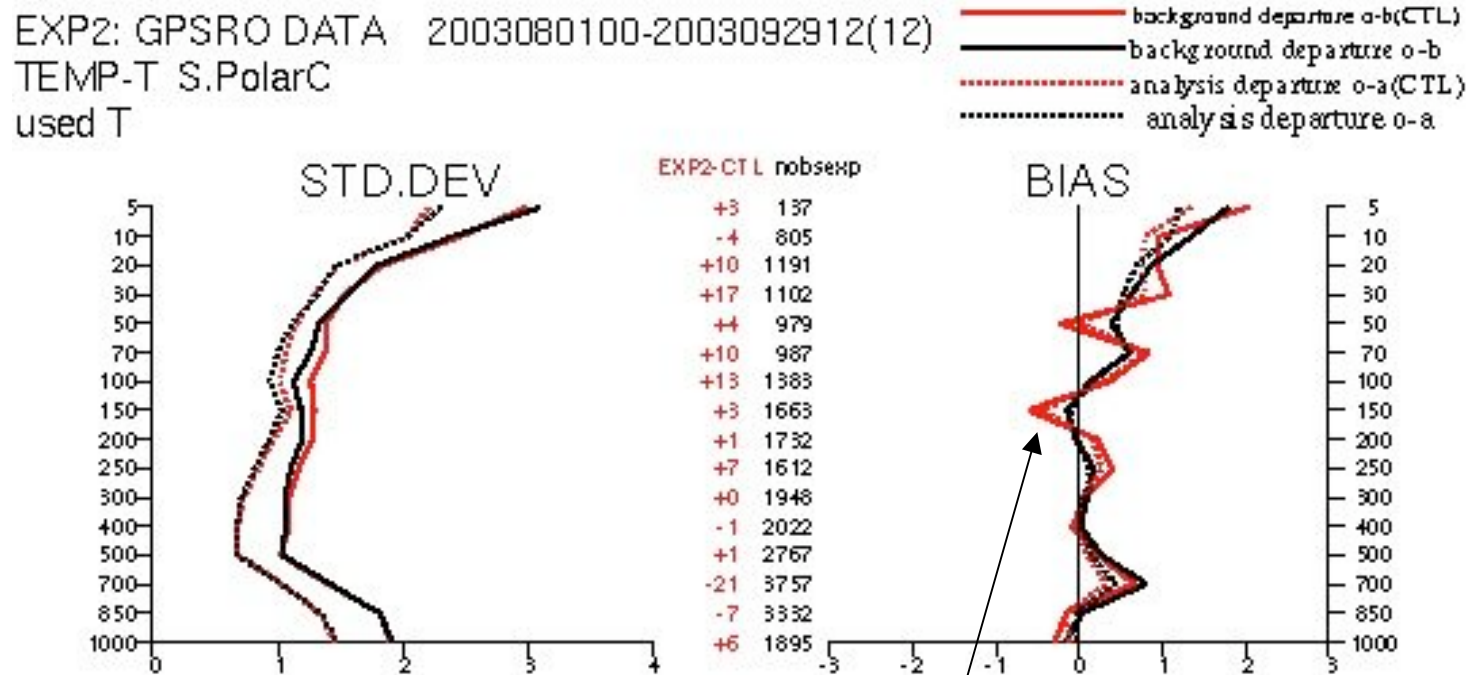


- = the path of the ray perigee through the atmosphere



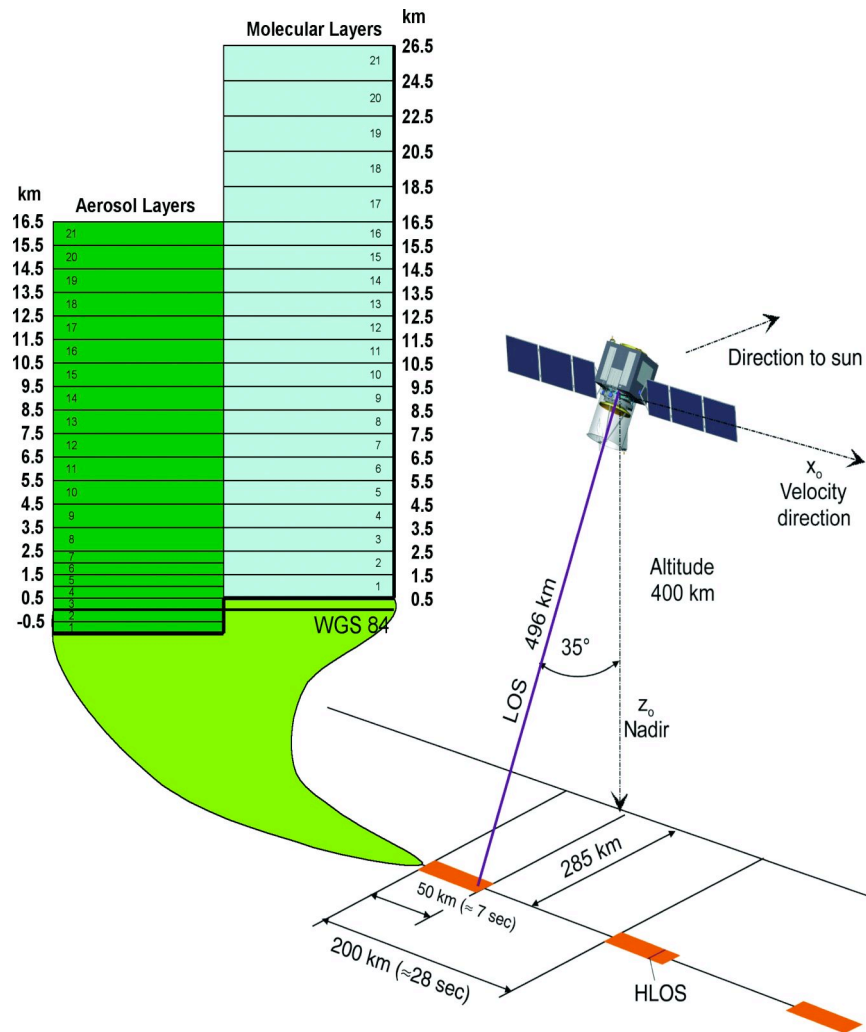
# Radiosonde comparisons for Antarctica

## 12h forecasts



# **Doppler Wind LIDAR**

# ADM-Aeolus

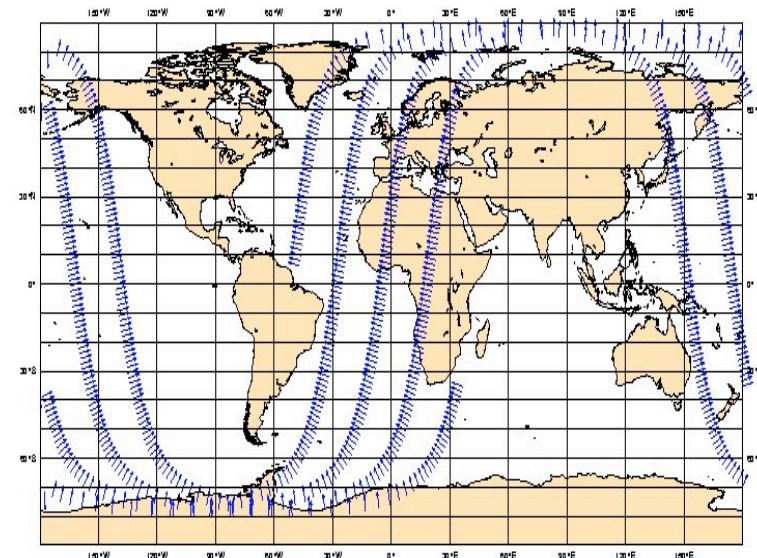


UV lidar (355 nm) with **two** receivers

- Mie (aerosol), Rayleigh (molecules)
- both use direct detection

Wind profiles from surface to 30 km with resolution varying from 0.5 to 2 km

- vertical bins configurable in flight
- HLOS component only
- direction 7° from zonal at equator
- 6 hour coverage shown



# **Environmental monitoring**

# ECMWF role in GMES

## Global Monitoring for Environment and Security

### | Reactive gases

- Couple IFS with global CTMs, carrying O<sub>3</sub>, CO, NO<sub>2</sub>, SO<sub>2</sub> and HCHO in the IFS, and develop data assimilation

### | Aerosols

- Implement in IFS, based on externally-produced parameterizations, and develop data assimilation

### | Greenhouse gases

- Introduce CO<sub>2</sub> and CH<sub>4</sub> into global ECMWF weather forecast model (IFS) and develop data assimilation

### | Acquisition of global data, and provision of support for regional air-quality forecasting

### | Near-real-time and retrospective (2003-2007) running of integrated global system

# Three types of product from ECMWF real-time system

## | Model simulation

- Extended sequence of 12-hour forecasts
- Meteorological variables are reset to analysed values from ECMWF archives
- GEMS composition variables are carried over from preceding 12-hour forecasts
- Used as basis for first set of near-real-time forecasts

## | Control

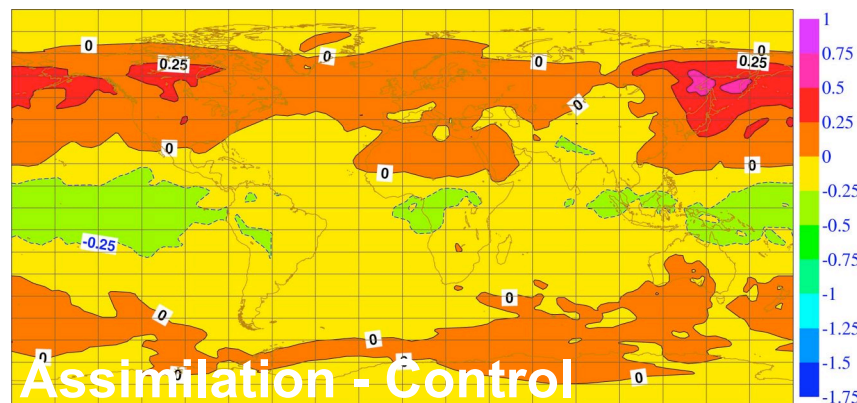
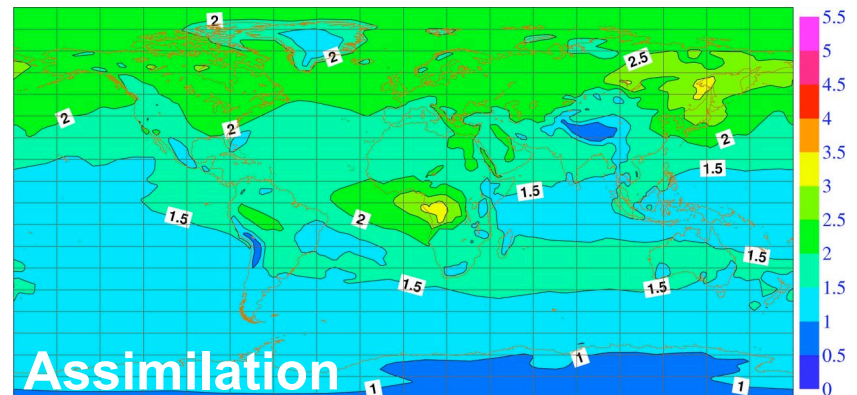
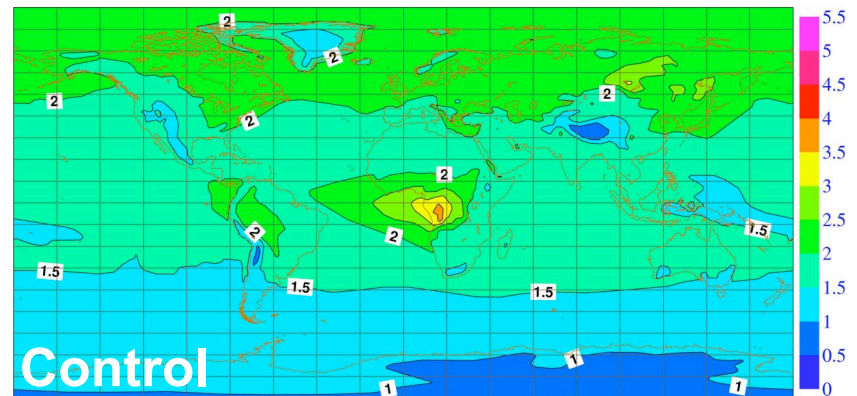
- Data assimilation for meteorological variables, with 6h or 12h cycling
- No data assimilated for GEMS composition variables
- Similar to model simulation, with differences due to different meteorological analyses and possibly cycling period

## | Assimilation

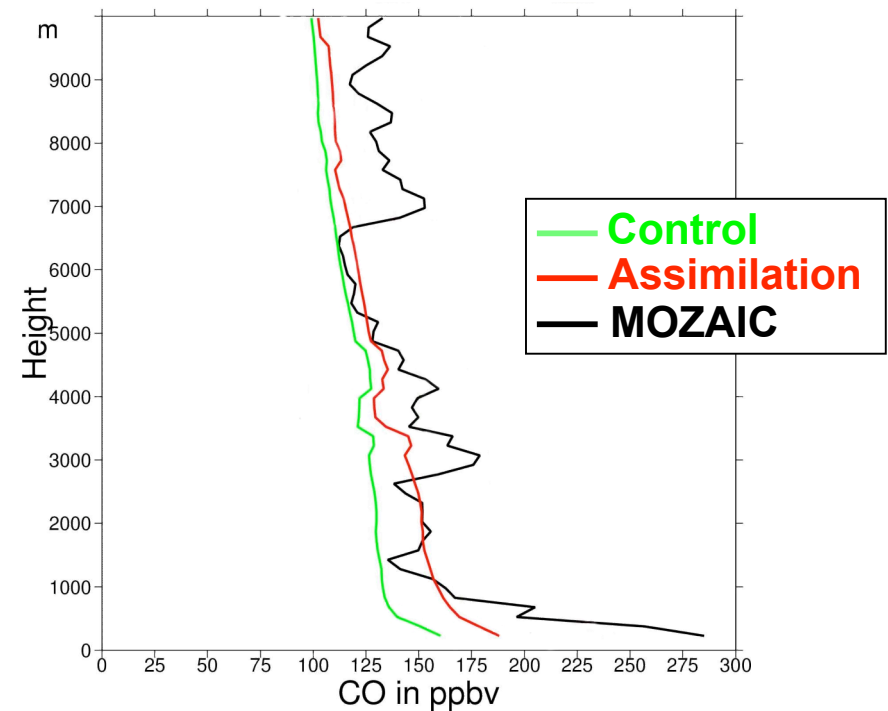
- Data assimilation for meteorological variables and (some) GEMS composition variables, with 6h or 12h cycling



# Mean CO from 15 to 30 July 2003 from assimilation of MOPITT total-column data



## Comparison with MOZAIC aircraft data over Osaka



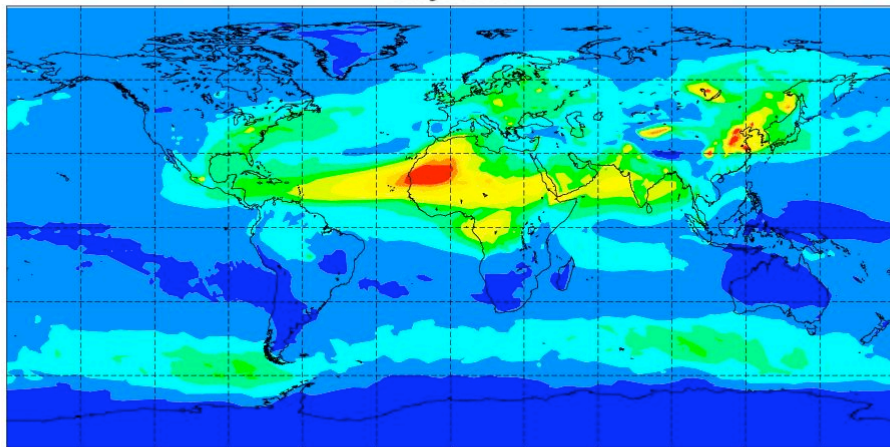
Carlos Ordóñez

Unit:  $10^{18}$  molec/cm<sup>2</sup>

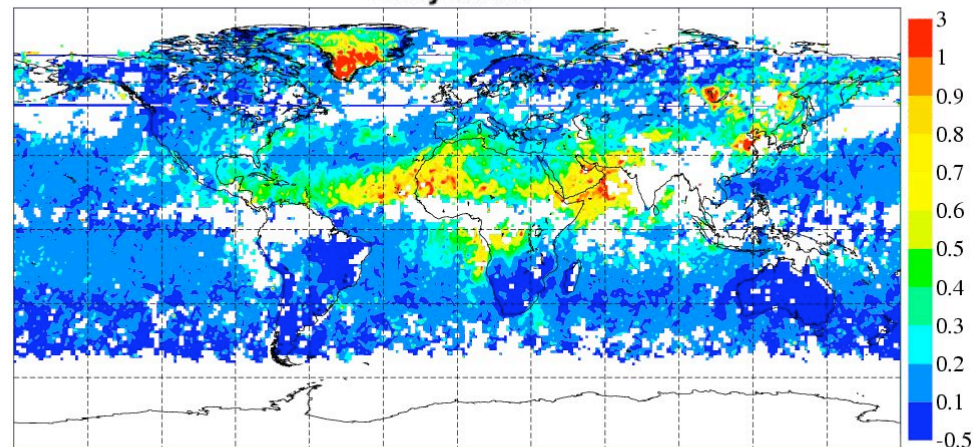


# Comparison of GEMS simulated and analysed aerosol optical depth with MODIS and MISR for July 2003

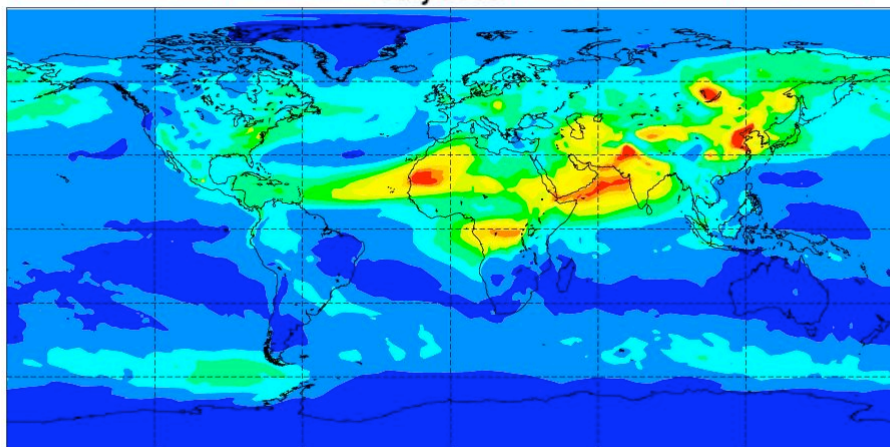
Aerosol Optical Depth at 550 nm from Unconstrained Model Run  
July 2003



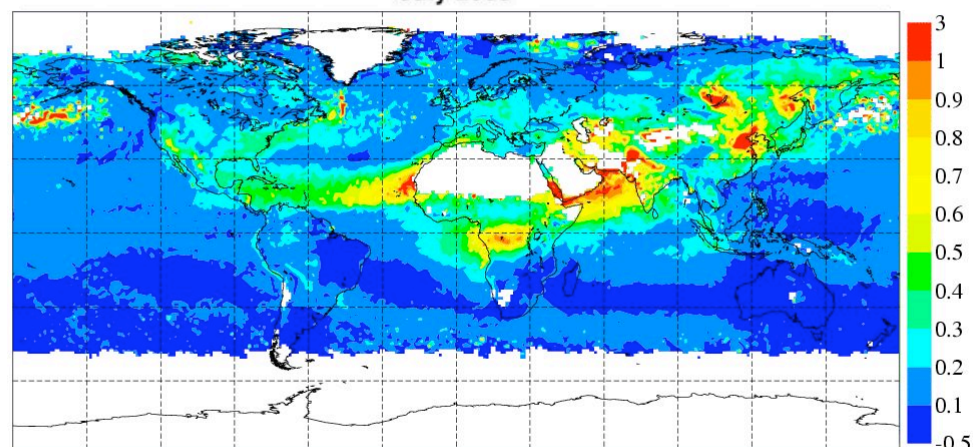
MISR Terra Aerosol Optical Depth at 557.5 nm [unitless]  
July 2003



Aerosol Optical Depth at 550 nm for Reanalysis using MODIS AOD  
July 2003

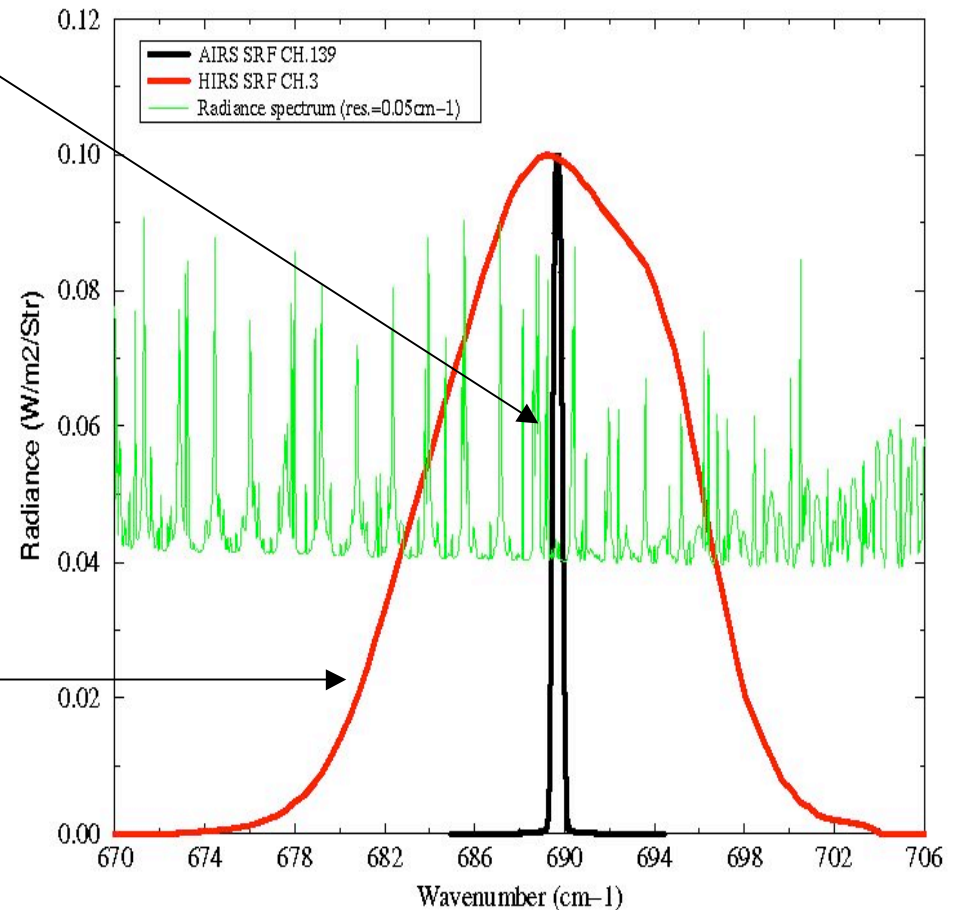


MODIS Terra MOD08-M3.005 Aerosol Optical Depth at 550 nm [unitless]  
July 2003



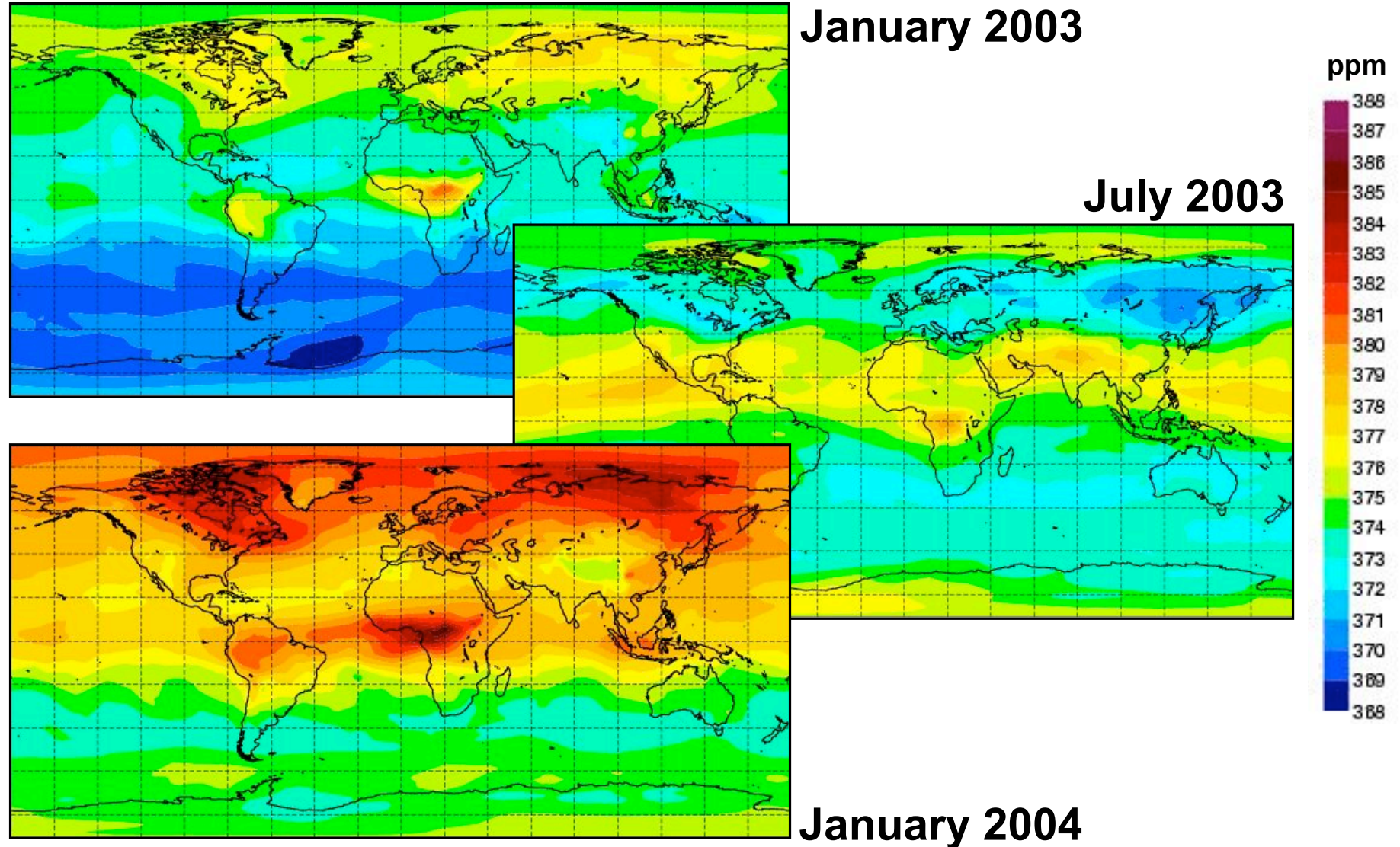
# Radiance assimilation in GEMS (AIRS/IASI)

- By sampling the IR spectrum at very high resolution ( $R=1200$ ) we can measure radiation that is only dependent on temperature and the atmospheric  $\text{CO}_2$  concentration (small groups of **pure lines**)
- If we have **accurate temperature information** (from the ECMWF analysis driven by AMSUA data) we can separate out the  $\text{CO}_2$  signal.
- Instruments with coarse spectral resolution (e.g. HIRS) sample radiation that is a mixture of absorbing species (e.g.  $\text{CO}_2$  /  $\text{N}_2\text{O}$  /  $\text{O}_3$  and  $\text{H}_2\text{O}$ ) and cannot resolve the  $\text{CO}_2$



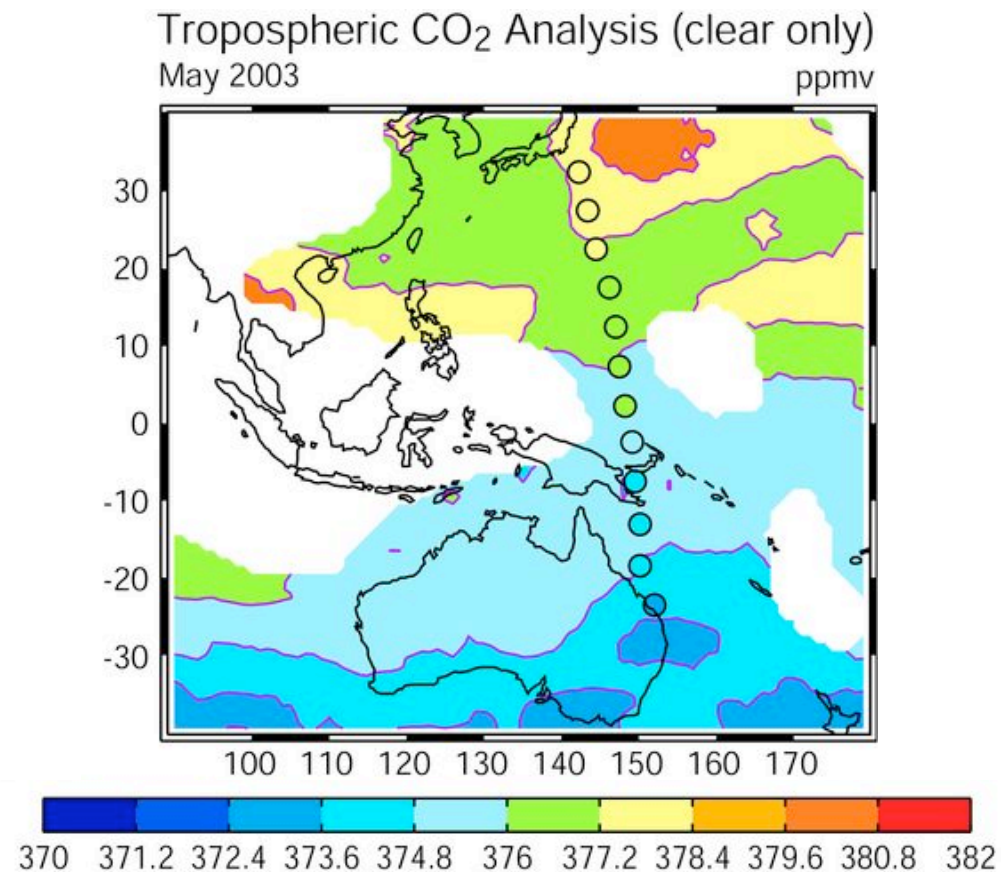


# Mean column CO<sub>2</sub> from assimilation of AIRS radiances



<http://gems.ecmwf.int>

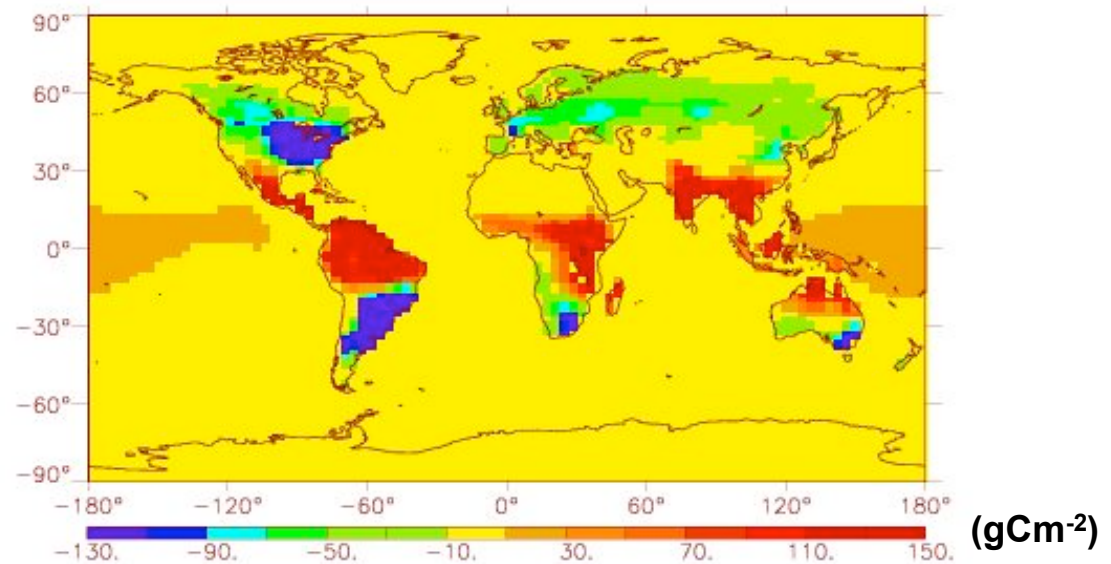
# Mean column CO<sub>2</sub> from assimilation of AIRS radiances: Validation with aircraft



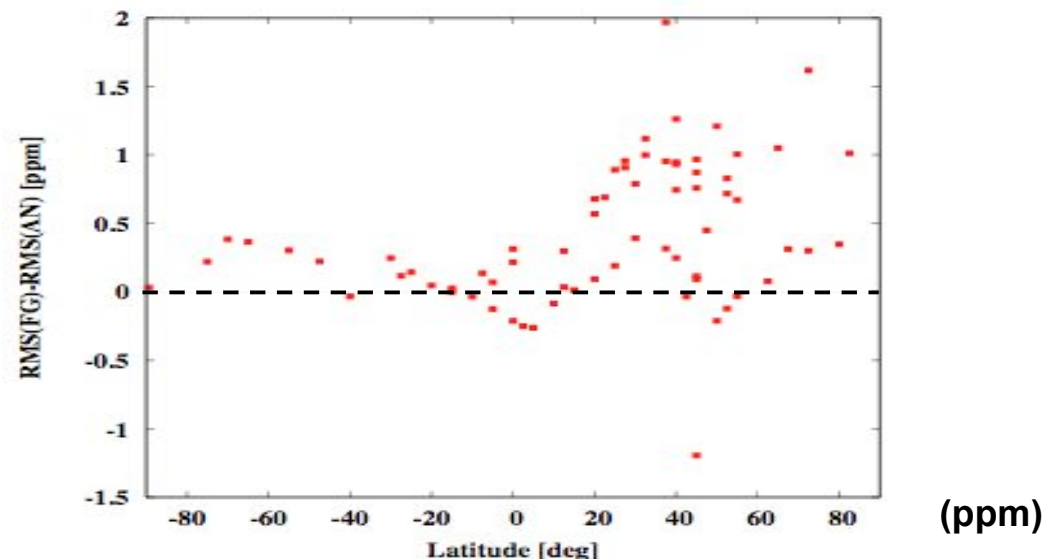
# CO<sub>2</sub> flux inversion from AIRS data assimilation

Cumulated surface flux increments derived from first CO<sub>2</sub> analyses for January-November 2003

Positive values imply flux into atmosphere.



Change in RMS fit to GLOBALVIEW-CO<sub>2</sub> (2006) data from use of incremented fluxes from latest analysis in LMDZ transport-model simulations



Peter Rayner, Frédéric Chevallier

# Reanalysis

## What is Re-analysis ?

Analysis of past (historical) observational data using a fixed, tried-and-tested, operational data assimilation system.

## What does it produce ?

A comprehensive time series of global analyses (i.e. gridded fields of temperature, humidity, wind etc..) and a homogeneous organized / quality controlled data-set of observations.

## What is it used for ?

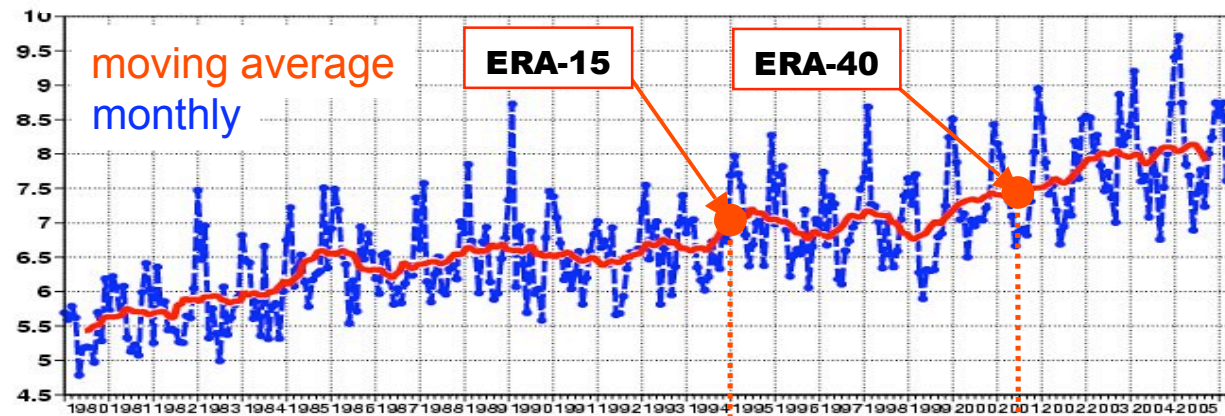
Meteorological research – into processes, composition, low-frequency variability, predictability, model development and general climate studies.



# Operational forecast performance 1980-2006

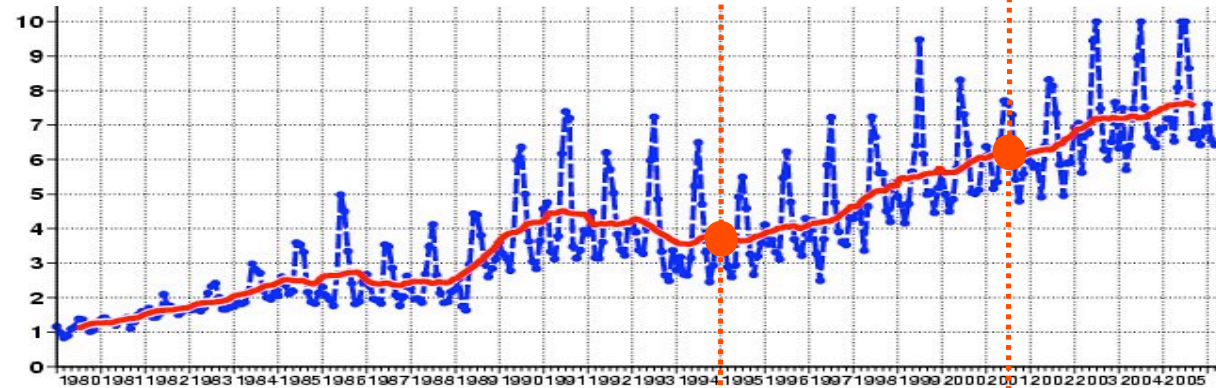
## Northern Hemisphere

500 hPa geopotential  
ANC reaching 60%



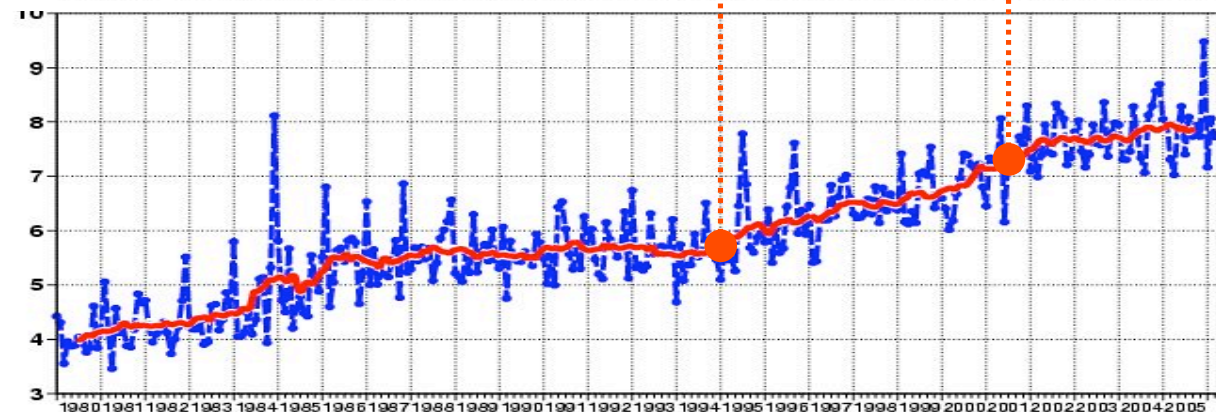
## Tropics

850 hPa wind vector  
ABC reaching 70%



## Southern Hemisphere

500 hPa geopotential  
ANC reaching 60%





# Some practical considerations ...

**The operational assimilation method and resolution is not affordable**

- We have to use a degraded, possibly less-tested configuration

**The operational system has known problems**

- We can implement recent fixes for some problems
- We may have to accept and try to account for other problems

**The operational system is tuned / based on current observing systems**

- We have to adapt the assimilation system to use older observing systems
- We have to handle changes in observation processing

**If errors are found during the course of production**

- We have to decide whether or not to stop / correct them / re-run

# International REANALYSIS activity

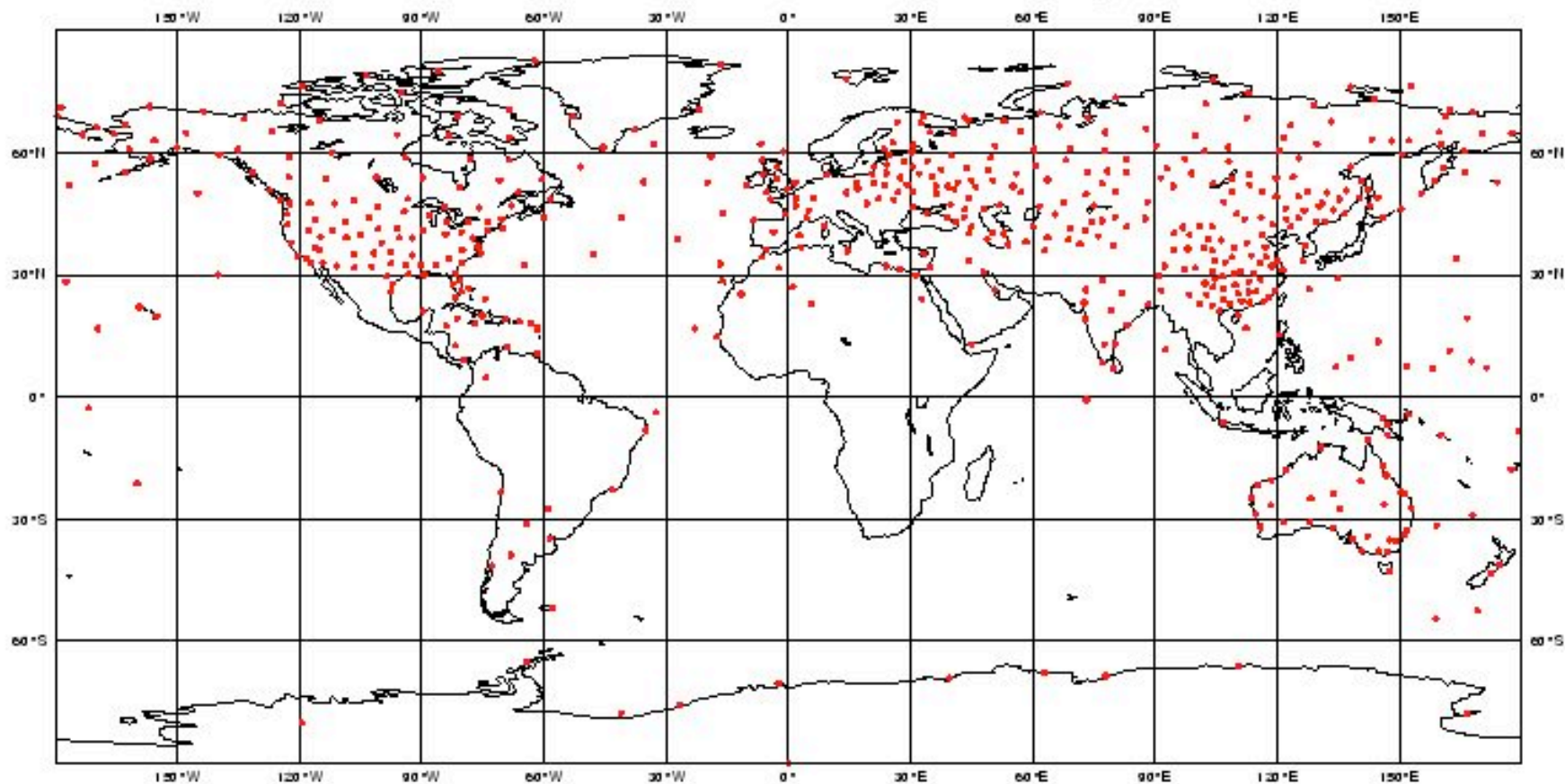
	NCEP/ NCAR	1948◇
	NASA/ DAO	1980-1995
	ECMWF, ERA-15	1979-1993
	ECMWF, ERA-40	1957-2002
	JMA/ JRA-25	1979-2004
	ECMWF, ERA-Interim	
	NASA GMAO	
	Polar Reanalysis	

## ***In-situ* (“conventional”) observations for ERA-40**

- **Radiosonde and pilot-balloon soundings** **1957 - 2002**
- **Surface data from land stations and ships** **1957 - 2002**
- **Flight-level data from commercial aircraft** **1973 - 2002**
- **Surface data from ocean buoys** **1979 - 2002**

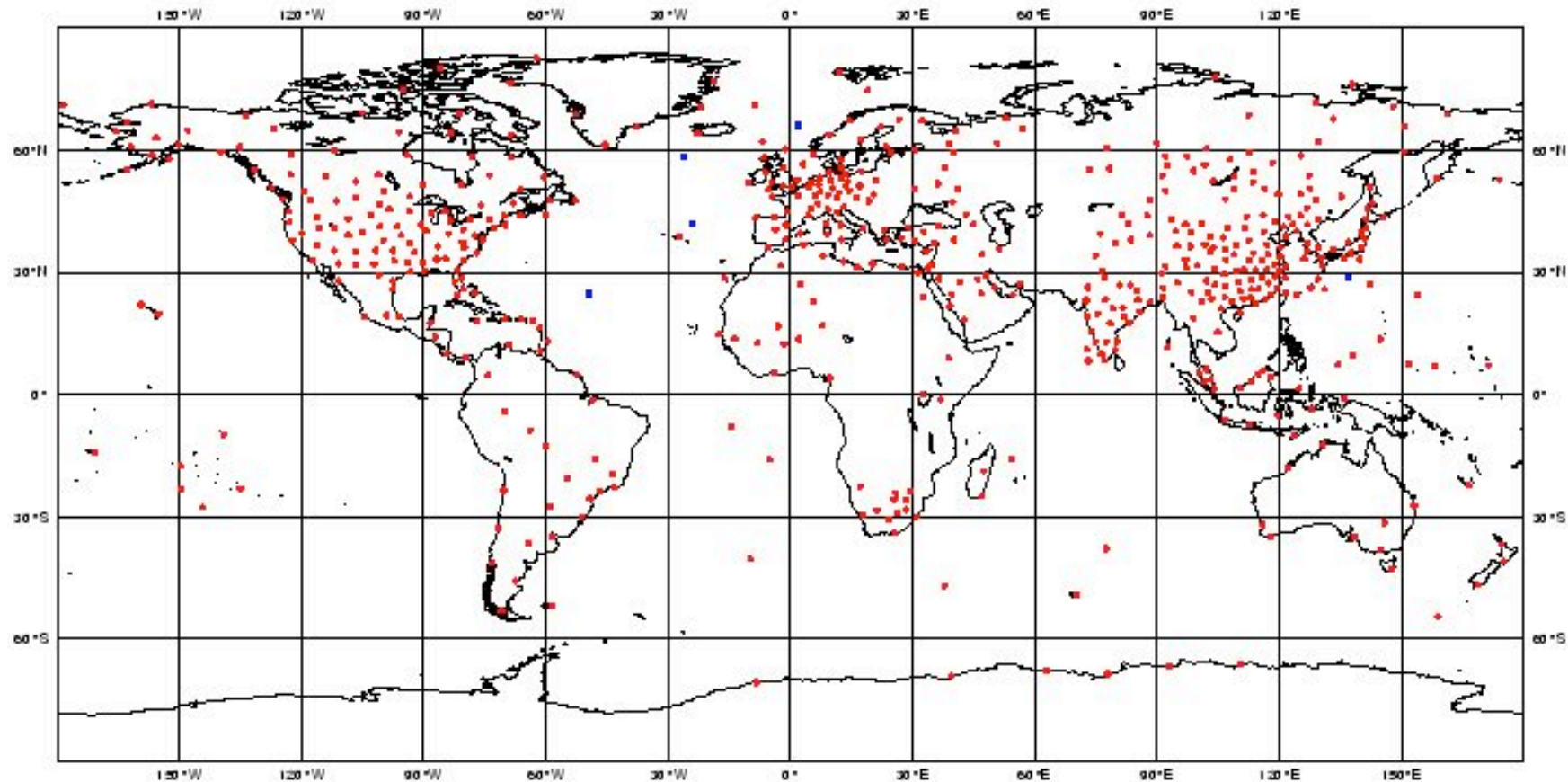
## Radiosonde coverage in October 1961

**Total number of obs = 669**



## Radiosonde coverage in May 1997

**Total number of obs = 564**

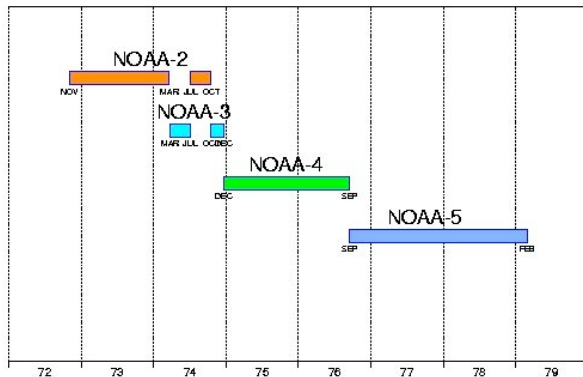


## Satellite data for ERA-40

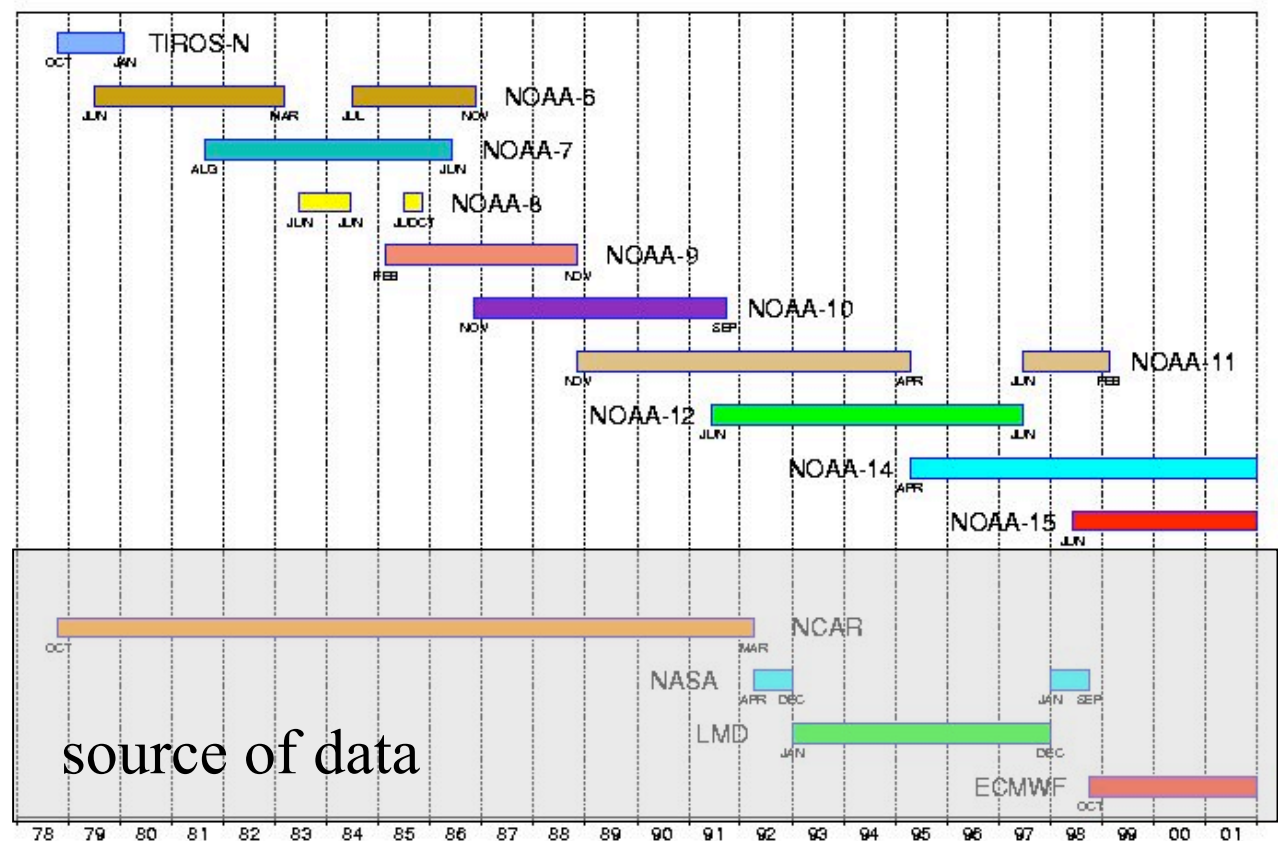
- NOAA VTPR radiances 1973 - 1978
- NOAA TOVS/ATOVS radiances 1979 - 2002
- Winds from geostationary orbit 1979 - 2002
- TOMS/SBUV ozone retrievals 1979 - 2002

# The ERA-40 Re-analysis used 41 satellite instruments carried by 15 different NOAA polar satellites

VTMR instrument 72-78



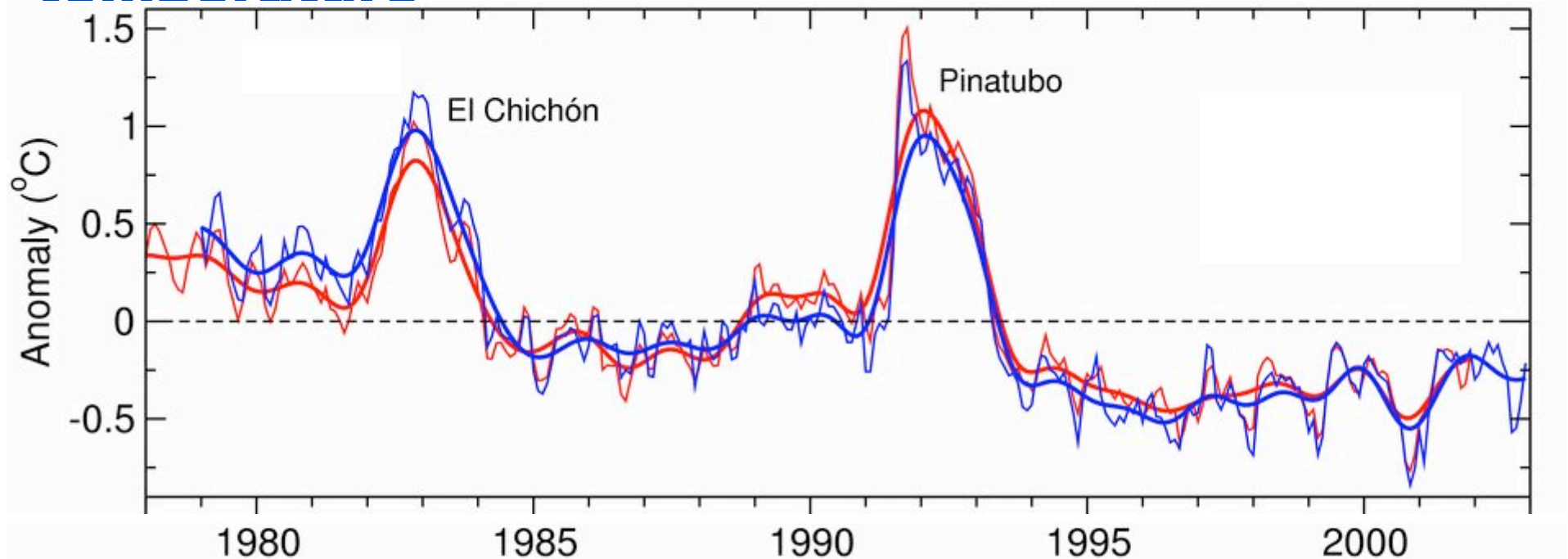
HIRS/MSU/SSU/AMSU instruments 78-02



**Some considerable  
reanalysis  
successes ...**



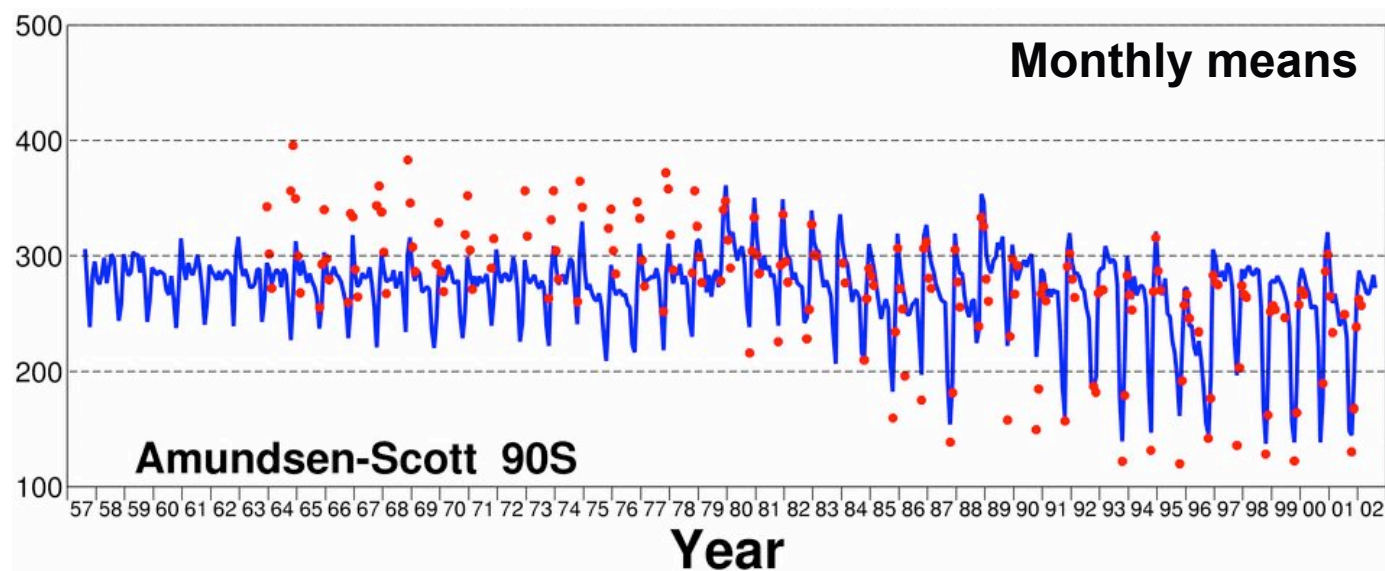
# Trend and variability in lower stratospheric temperature



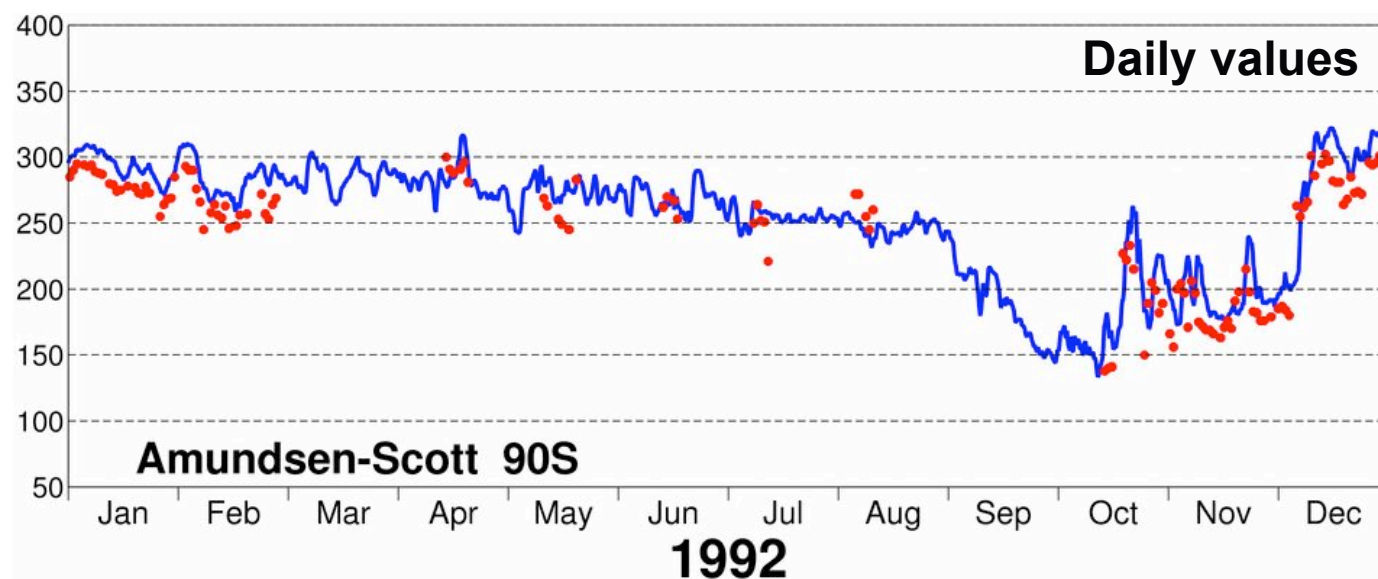
— MSU-4 data analyzed by Mears et al. (2003)  
— ERA-40 equivalent from Ben Santer

Linear trend:	MSU-4	- 0.39°C/decade
	ERA-40	- 0.30°C/decade
	NCEP	- 0.82°C/decade

# Total ozone (Dobson units)



**Blue: ERA-40  
(TOMS and SBUV  
data assimilated  
1979-1988 and  
1991-2002)**



**Red: Ground-  
based  
measurements  
(NOAA/CMDL)**

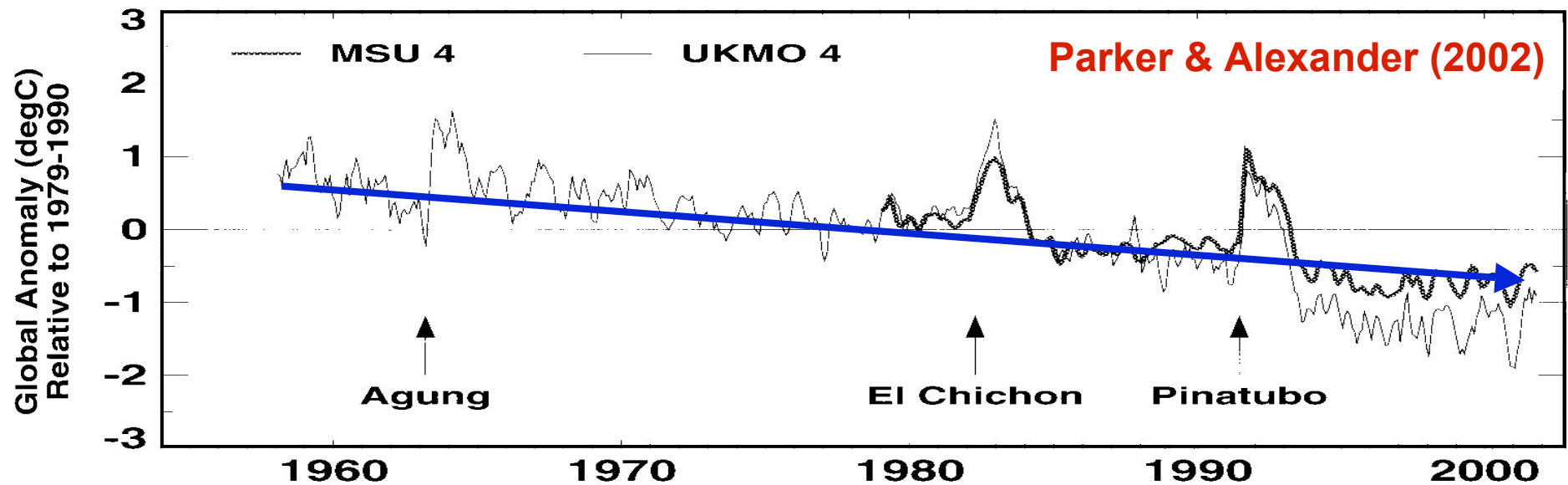
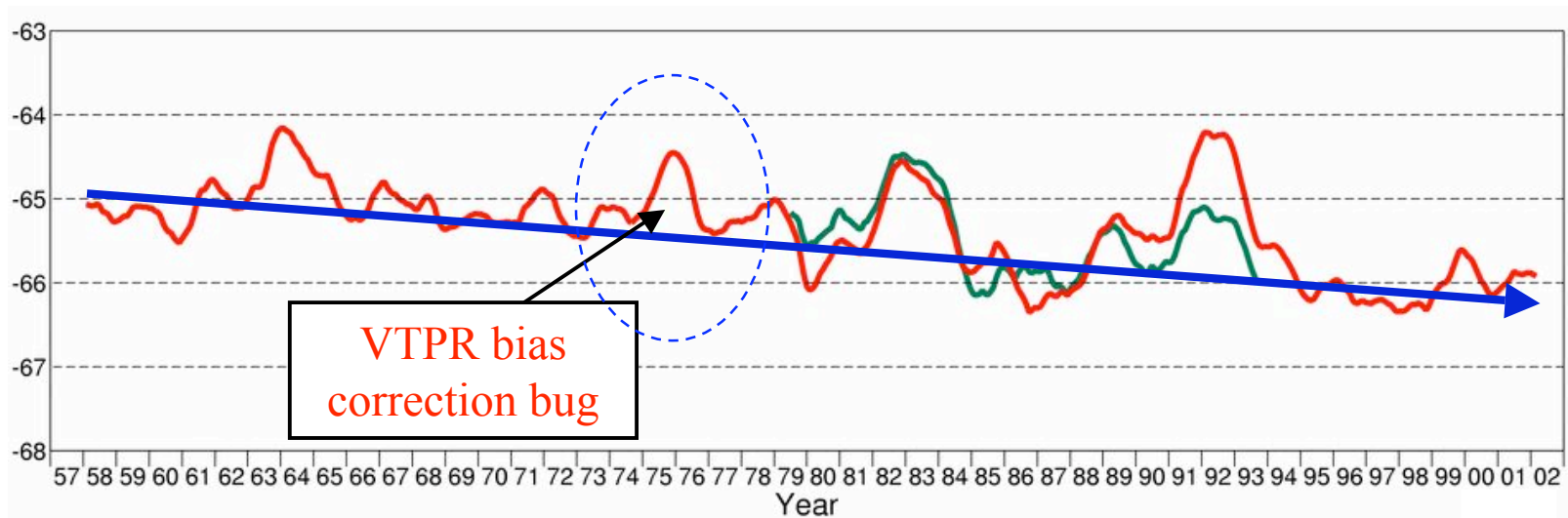
**But...**

**some considerable  
reanalysis issues ...**

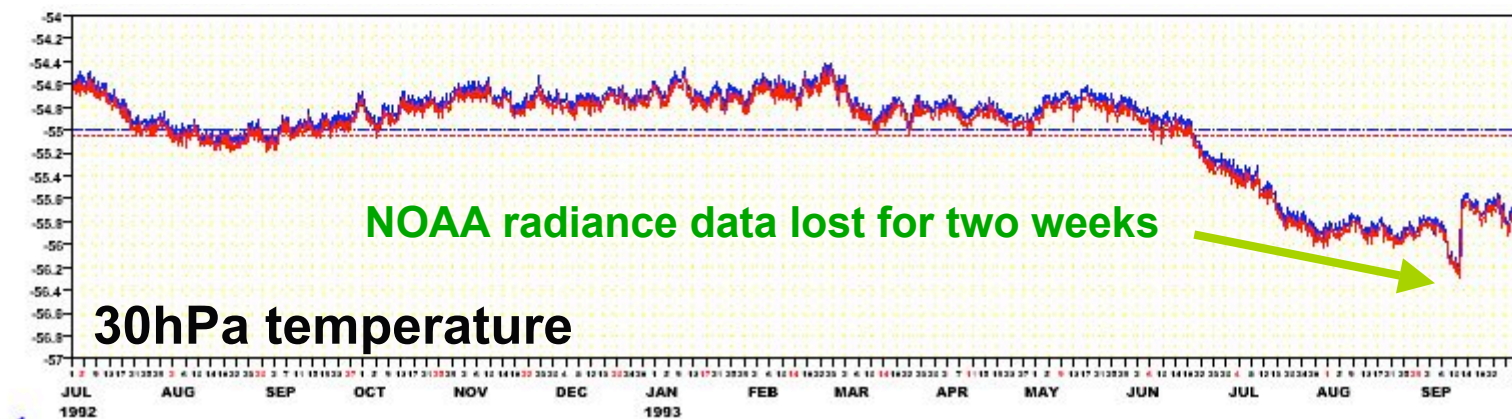
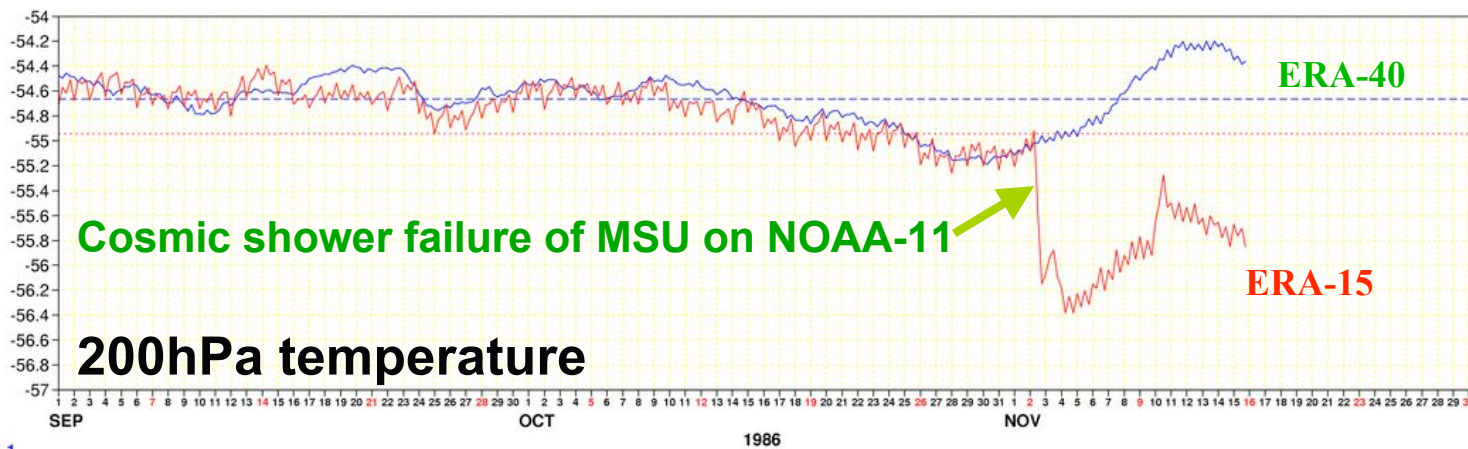
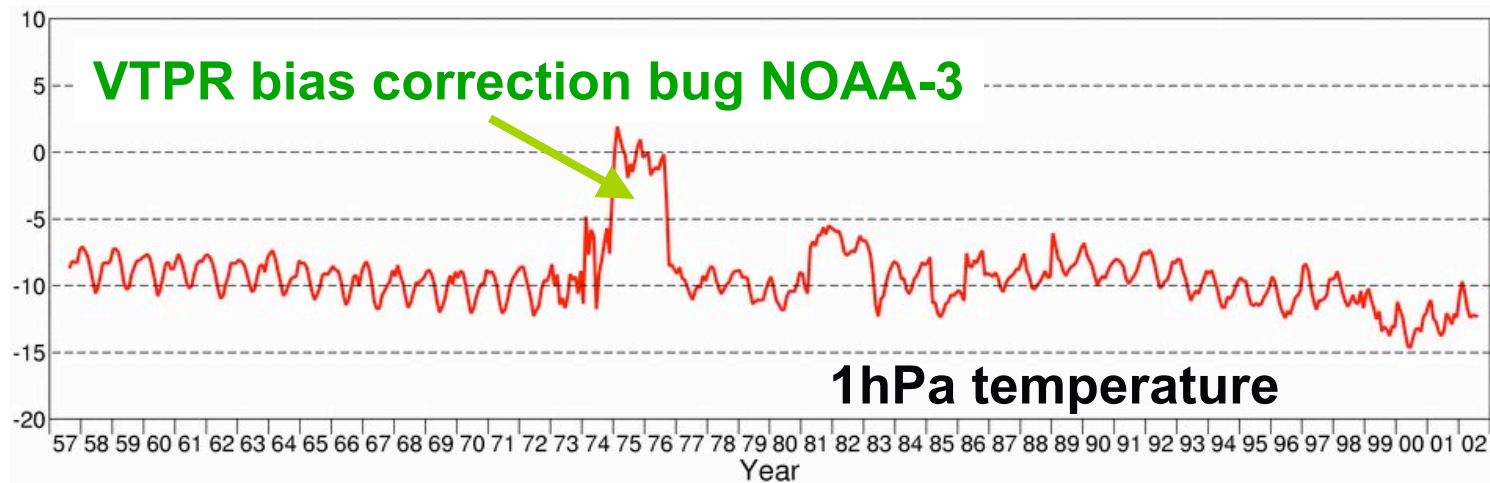
# Particular Issues related to the use of satellite data in re-analysis

- Instrument **drift/shift** over the lifetime of a satellite
- Absolute and **inter-satellite calibration** between different satellites
- Intermittent (sudden) **disruptions / contamination** of data (by nature)
- Changes to **channel / instrument** payload
- **Events are often difficult to detect / fix in time during production...**

## Trends in Global 70hPa temperature (real and artificial)



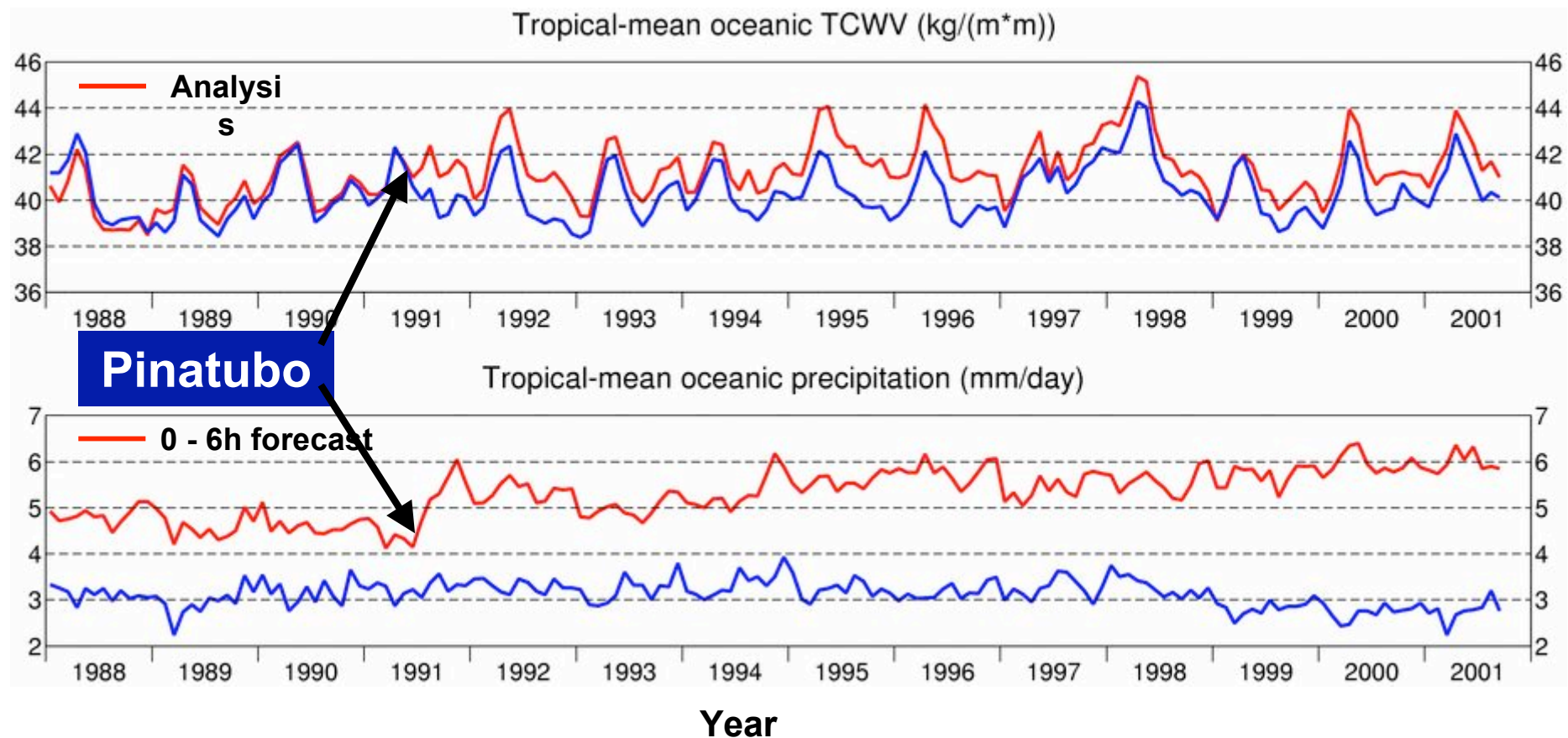




# Aspects of tropical humidity analysis

— ERA-40

— SSM/I retrievals from Remote Sensing Systems



**The next step...**

**ERA-interim**



**ERA-Interim 1989 ◇ to continue...**

**ERA-40 1957-2002**

- | **Data-assimilation system**

  - 12 hour 4D-Var**

- | **Satellite radiances**

  - Adaptive bias correction**

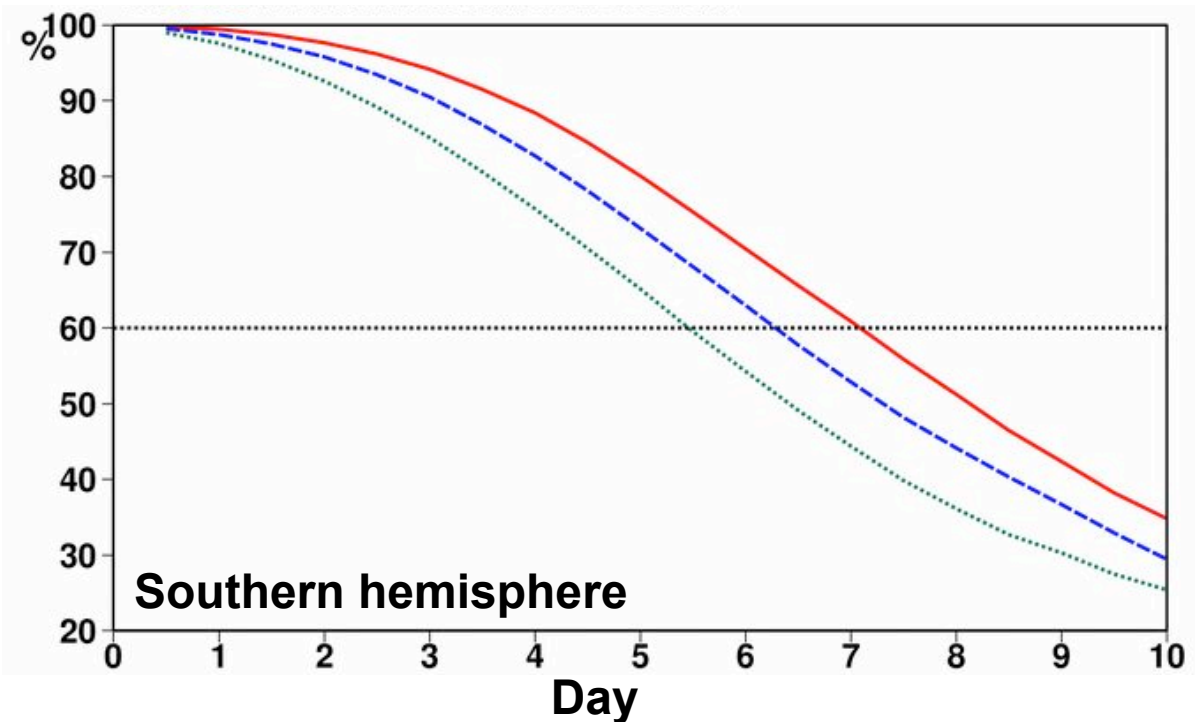
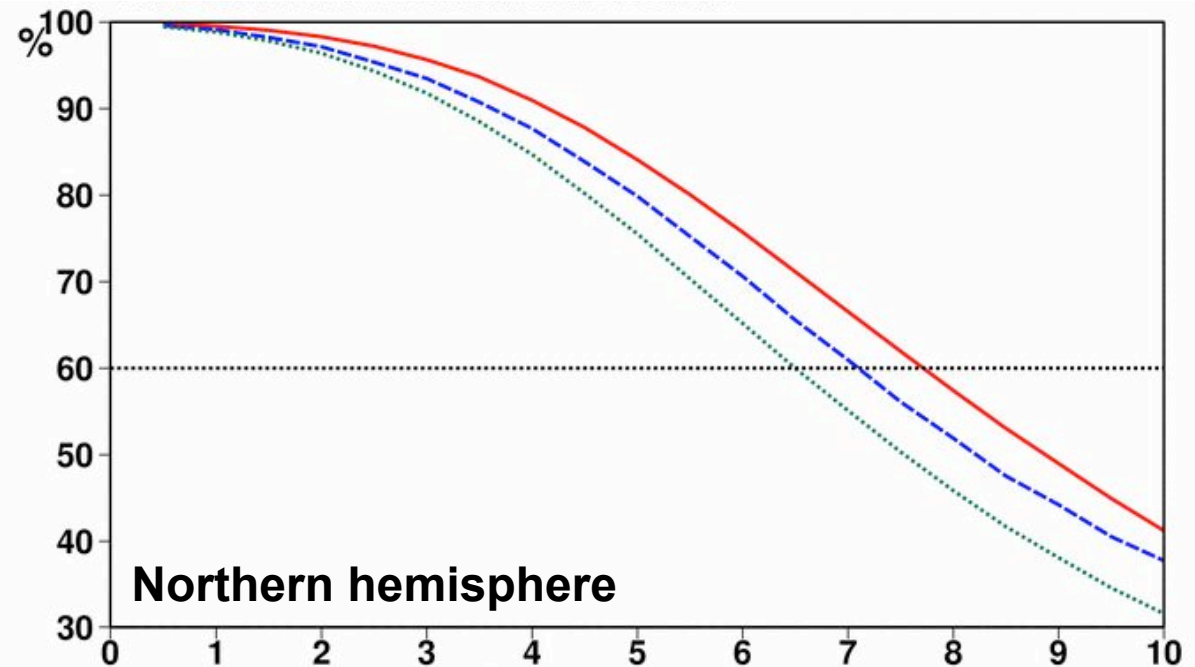
- | **Improved use of radiosondes**

  - Bias correction and homogenization  
based on ERA-40**

# Much improved 500hPa height forecasts with 4DVAR

Anomaly  
correlation of  
500hPa height,  
averaged for  
12UTC forecasts  
from 1 January to  
31 December 1989

— ERA-interim  
- - - ERA-  
... Operations 1989



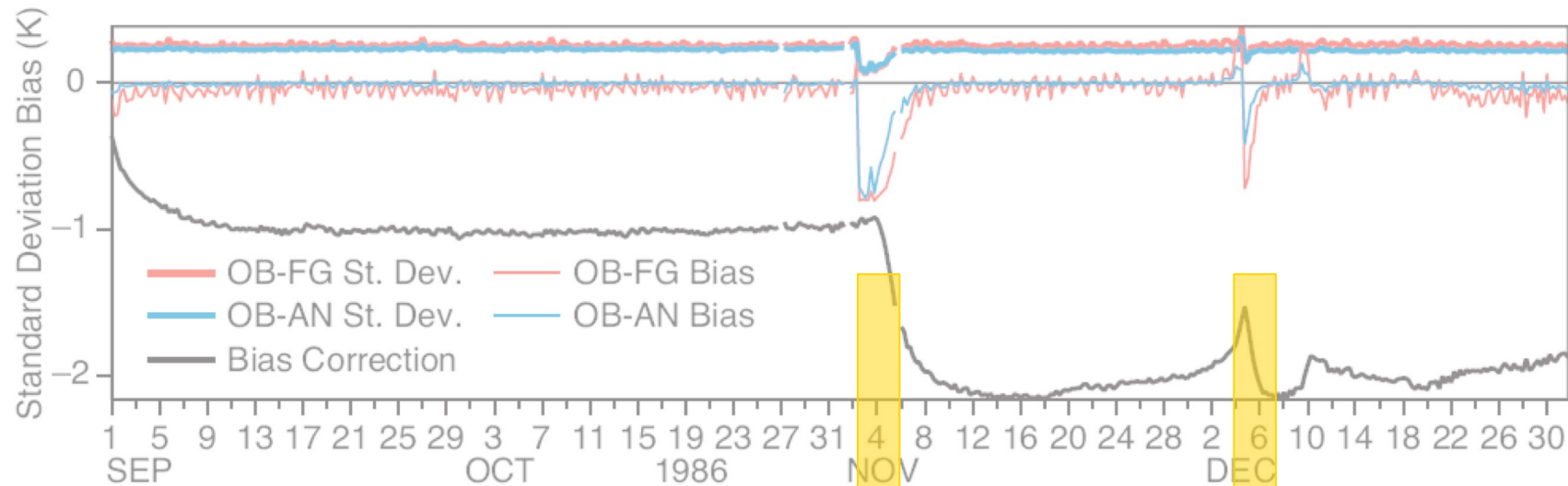
# Variational bias correction of radiance data

- Radiance bias correction expressed in terms of a small number of unknown parameters:
  - A constant offset
  - Predictors depending on instrument scan position (scan bias)
  - Predictors depending on model state  $x$  (air-mass dependent bias)
- Add the bias correction parameters to the control vector in the variational analysis

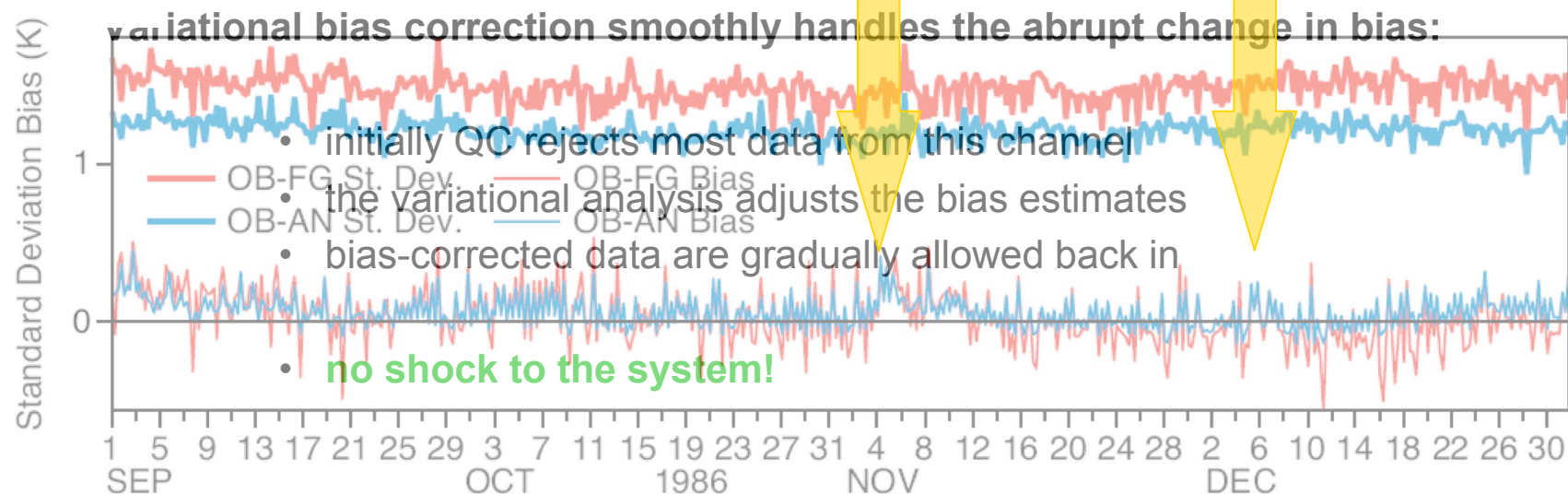
$$\begin{aligned} \mathbf{J}_b: \text{background constraint for } x & \quad \mathbf{J}_\beta: \text{background constraint for } \beta \\ \mathbf{J}(x, \hat{a}) = & \underbrace{(\mathbf{x}_b - x)^T \mathbf{B}_x^{-1} (\mathbf{x}_b - x)}_{\mathbf{J}_b} + \underbrace{(\hat{\mathbf{a}}_b - \hat{a})^T \mathbf{B}_a^{-1} (\hat{\mathbf{a}}_b - \hat{a})}_{\mathbf{J}_\beta} \\ & + \underbrace{\left[ y - \mathbf{b}_o(x, \hat{a}) - \mathbf{h}(x) \right]^T \mathbf{R}^{-1} \left[ y - \mathbf{b}_o(x, \hat{a}) - \mathbf{h}(x) \right]}_{\mathbf{J}_o: \text{bias-corrected observation constraint}} \end{aligned}$$

- The analysis then estimates bias parameters jointly with model state variables (Derber and Wu 1998)

## NOAA-9 MSU channel 3 bias corrections (cosmic storm)



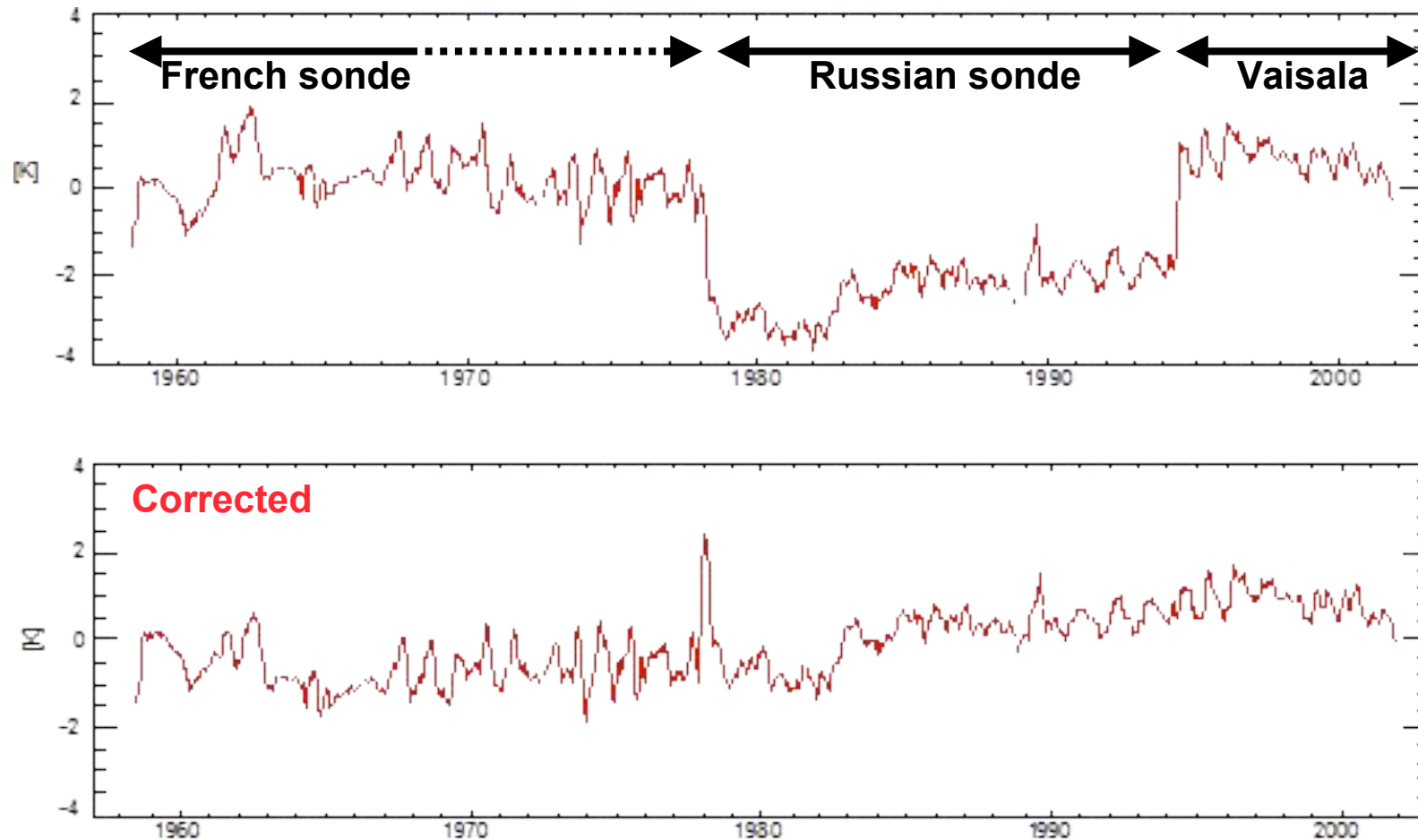
### 200 hPa temperature departures from radiosonde observations



# Need to homogenize radiosonde biases in time

(example: Haimberger, 2005 using ERA-40 feedback data)

SAIGON / TAN-SON-NHUT 00UTC 200hPa temperature (Background – Observation)



**THE END**