Remote Sensing of Sea Ice

Peter Lemke

Alfred Wegener Institute for Polar and Marine Research Bremerhaven

> Institute for Environmental Physics University of Bremen

Contents

- 1. Ice and climate
- 2. Sea Ice Properties
- 3. Remote Sensing Techniques

Contents

- 1. Ice and climate
- 2. Sea Ice Properties
- 3. Remote Sensing Techniques

The Cryosphere



Climate System



Role of Ice in Climate

- V Impact on surface energy balance (global sink)
 - \Diamond Atmospheric circulation
 - \Diamond Oceanic circulation
 - \Diamond Polar amplification
- V Impact on gas exchange between the atmosphere and Earth's surface
- V Impact on water cycle, water supply
- V Impact on sea level (ice mass imbalance)
- **v** Defines boundary conditions for ecosystems

Role of Ice in Climate

V Polar amplification in CO₂ warming scenarios

<u>(surface energy balance; temperature – ice albedo feedback)</u>



Role of Ice in Climate



Warming from CO₂ doubling with fixed albedo (FA) and with surface albedo feedback included (VA) (Hall, 2004) Warming at CO₂ doubling (years 61-80) (Winton, 2006)

Contents

- 1. Ice and climate
- 2. Sea Ice Properties
- 3. Remote Sensing Techniques

Pancake Ice



Pancake Ice



Old Pancake Ice



First Year Ice

Pressure Ridges

Pressure Ridges

New Ice Formation in Leads

New Ice Formation in Leads

New Ice Formation in Leads

Melt Ponds on Arctic Sea Ice

Melt ponds in the Arctic

Sediment Patch on Arctic Sea Ice

Sea Ice Formation

Deformation: Pressure Ridges

Sea Ice Reduces Heat Exchange

Deep and Bottom Water Formation

Sea Ice Types

Class	Description	Thickness		
New ice	ice which began to grow a few hours or days ago	0 – 10 cm	 h	nigl
Young ice	transition between new and first-year ice	10 - 30 cm		₹
First-year ice	ice of no more than one winter's growth	30 – 200 cm		alini
Old ice	ice that has survived at least one summer's melt; most topographic features are smoother than on first-year ice	> 200 cm] ↓ ∣	ю Iow

Vertikal Structure

Vertikal Structure

Salinity Profiles

Vertical Salinity Profiles (Eicken, 1991)

Salinity vs. Thickness (Age)

Average salinity of sea ice as a function of ice thickness for cold sea ice sampled during the growth season (Cox and Weeks, 1974).

Salinity determines thermal conductivity: $\lambda_{si} = \lambda_i + \frac{\beta S_i}{T}$ with $\lambda_i = 9.82e^{-0.0057T}$ salinity determines electromagnetic properties:

i.e. dielectric constant

Contents

- 1. Ice and climate
- 2. Sea Ice Properties
- 3. Remote Sensing Techniques

Sea Ice: Measurement of Properties

Sea Ice Parameters

Fundamental

- albedo
- extent
- concentration
- drift
- type distribution
- thickness

Measurements

satellite imaging (radar, passive µ-wave, optical and IR) local: aircraft measurements+reconnaissance, ship reports Specific

- characteristics of leads and polynyas
- ice deformation and roughness
- melt onset and freeze-up
- ice floe size distribution
- thin ice type separation, frost flower detection
- snow properties

 (thickness, grain size...)

radar and laser altimetry, sonar, EM induction sounding, drilling

Remote Sensing of the Cryosphere

What Is Directly Measured ?

Surface Reflectivity (VIS and NIR)

Brightness Temperature (IR, μ-wave) μ-wave: Τ_b=εT

Radar Backscattering Coefficient

+ spatial variations of these parameters

Remote Sensing of Sea Ice Parameters

ice variable	VIS-	IR	MR	SAR	ALT	SCAT	LASER	SONAR
	NIR							
extent	++	++	++	+	++	+	++	++
coverage	++	++	+	++	(+)		++	++
floe size		_						
distribution	++	+		++	(+)			(+)
ice type			+	(+)	(+)			
ice thickness					(+)		+	++
ice				(+)	++	(+)	++	++
roughness albedo	++							
drift	++	++	++	++		+		(+)
horiz.								
resolution	>10m	1 km	30 km	30 m	1 km	50 km	1m	5 m

 V "Passive": The sensor consists of a receiver and detects the natural radiation.

V The sensor measures the thermal radiation of the Earth's surface in the range of 1-100GHz (microwave region).

 Advantages: measurements do not depend on light conditions and are almost independent of weather conditions. Data are obtained with good spatial and temporal sampling (global coverage of sea ice covered area every day), stable instrument calibration.

 V Disadvantage: Spatial resolution is only poor (5-80 km, dependent on frequency and sensor), mixed pixel problem, coastal contamination effects large variability

Planck Spectrum

All substances emit electromagnetic radiation. This thermal radiation is a function of temperature. For perfect emitters, i.e. *blackbodies*, the spectral radiance L_{ν} is related to the absolute temperature through Planck's^{II} Law (Fig. 6.5). L_{ν}^{\star} is the energy flux density per solid angle (sr) and frequency interval with the unit $[Wm^{-2}Hz^{-1}sr^{-1}]$.

Thermal emission in MW Domain

In the microwave domain, where $h\nu \ll kT$, the Rayleigh-Jeans Law applies, which represents an approximation of the Planck Law (6.9). With $\exp[h\nu/(kT)] \simeq 1 + h\nu/(kT)$ one finds

$$L_{\nu} \simeq \frac{2\nu^2 k}{c^2} T$$

Therefore, the spectral radiance in the microwave domain is directly proportional to the absolute temperature, and a definition of the so-called *radiative temperature* T_B in terms of the observed radiance L_{ν} is possible.

radiance L_{ν} is possible. brightness temperature $T_B = \frac{c^2}{2\nu^2 k} L_{\nu}$ for a black body

for natural bodies

$$T_{i}(\theta) = \rho_{i}(\theta)T_{s} + e_{i}(\theta)T_{g}$$

$$T_{i}(\theta) = T_{g} + \rho_{i}(\theta)(T_{s} - T_{g})$$

$$T_{i}(\theta) = T_{s} + e_{i}(\theta)(T_{g} - T_{s})$$

$$emissivity$$

$$e_{i} = 1 - \rho_{i}$$
reflectivity

 $-\rho_i$

temporal evolution of sea ice:

Øaging processes Øsalinity reduction Øreflectivity change $\Delta \rho = \rho \frac{2\Delta \epsilon}{5(m+1)}$

$$\Delta \rho = \rho \frac{2\Delta \epsilon}{\sqrt{\epsilon} \left(\epsilon - 1\right)}$$

 $\in=3$ and $\Delta\in=0.6$ (20% change)

$$\frac{\Delta \rho}{\rho} = 0.34 \quad \text{and} \quad \Delta \rho = 0.34 \times 0.07 = 0.024$$
$$\Delta T = 5.4 \text{ K} \text{ i.e. ice type also easily detectable}$$

Sea Ice Concentration Algorithm

Mixed pixel: sea ice and water (C = ice concentration)

$$T_B = (1 - C)e_w T_w + Ce_i T_i$$

$$C = \frac{T_B - e_w T_w}{e_i T_i - e_w T_w}$$

Mixed pixel: three surface types: First-year sea ice, multi-year sea ice, and water $T_B = (1 - C_f - C_m)e_wT_w + C_fe_fT_f + C_me_mT_m$

Measurements at more than one frequency or polarization required

sea ice with dry snow cover

sea ice with wet snow cover

Microwave Signatures as a Function of Frequency and Polarizations

Fig. 4-18. Polarization and spectral characteristics of open water, first-year and multiyear sea ice as observed with the DMSP SSM/I on January 17, 1988.

Separation of ice and water at lower frequencies, using different polarizations

 $PR = \frac{T_B[18V] - T_B[18H]}{T_B[18V] + T_B[18H]}$ (Polarization Ratio)

Separation of FY and MY ice by combination of lower and higher frequencies

$$GR = \frac{T_B[37V] - T_B[18V]}{T_B[37V] + T_B[18V]}$$

(Gradient Ratio)

 \Rightarrow Why are ratios used?

In order to minimize effects of physical temperature.

Fractions of first- and multi-year ice:

$$C_{f} = \frac{a_{0} + a_{1}PR + a_{2}GR + a_{3}PR \cdot GR}{D}$$

$$C_{m} = \frac{b_{0} + b_{1}PR + b_{2}GR + b_{3}PR \cdot GR}{D}$$

$$D = c_{0} + c_{1}PR + c_{2}GR + c_{3}PR \cdot GR$$
Open water fraction?
$$C_{w} = 1 - C_{f} - C_{m}$$

The coefficients a_i , b_i , c_i are derived from "tie points", i.e. measurments over 100% FY ice, MY ice, and OW. For the NASA Team Algorithm, this is carried out at 18GHz V+H, and 37GHz V.

Scatterplots of Brightness Temperatures

Separation of ice and water at lower frequencies, using different polarizations Separation of FY and MY ice by combination of lower and higher frequencies

Sea ice extent

Sea ice drift

Polar Science Center Applied Physics Laboratory University of Washington

Annual mean of ice motion in the Arctic based on 1979 - 1990 buoy data. The superimposed lines indicate the number of years til the ice exits the arctic basin through the Fram Strait

Sea Ice in Fram Strait

AVHRR - VIS:

Drift Floe size distributiuon

Sea ice drift

Sea Ice Drift from SSM/I Images

Winter 94/95

Sea ice drift and -deformation from SAR-imagery

RADARSAT Geophysical Processor System. Due to a large swath width (~400 km) in ScanSAR mode, an image is obtained at least every 7. day (by courtesy of Ron Kwok).

Sea Ice Thickness

Ice thickness measuring techniques

EMI on the ice surface

- EM31 (Geonics)
 - Coil seperation3.66m
 - Frequency 9.8 kHz
 - coils coplanar
 - measures conductivity and inphase ppms

EMI on the ice surface

Germany's smallest research vessel... (the yellow one)

A normal 12h station

EM Thickness Profiles

EMI ship-borne

The SIMS (Sea Ice Monitoring System)

EMI airborne

AWI's Helicopter EM system

2 Frequencies -3.68 kHz -112 kHz

3.4 m length 100 kg weight

Airborne EM ice thickness sounding

EM induction sea ice thickness sounding

Ice thickness variability in the Transpolar Drift: 1991, 1996, 1998, 2001 & 2004

- The travel time of a pulse between the sensor's position and the ice surface is measured (accuracy of a few centimeters, atmospheric corrections are required).
- The orbit of the satellite is also determined with centimeters accuracy.

Freeboard and Ice Thickness

Retrieval of Sea Ice Thickness

- Individual echos are separated into groups "FY- and MY-ice" and "Open water and thin ice"
- Freeboard is obtained by subtracting travel times over water from travel times over ice.
- Conversion of freeboard into thickness: assumption of hydrostatic equilibrium, using average ice and water densities, and snow ,,climatologies"

Ice thickness t_E at hydrostatic equilibrium and a snow load of mass m_S per unit area:

$$\boldsymbol{t}_{E} = \frac{\rho_{W}}{\rho_{W} - \rho_{E}} \boldsymbol{f}_{E} + \frac{1}{\rho_{W} - \rho_{E}} \boldsymbol{m}_{S}$$

Altimeter

Freeboard measurements by satellite is the only technique which can provide sea ice thickness data at the time and Length scales required by climate research.

Radar altimeter (ERS):

Accuracy of thickness

estimates: ~0.2m

Thanks for your attention

